SIMULATION OF GROUND-WATER FLOW IN

THE MESILLA BASIN, DOÑA ANA COUNTY, NEW MEXICO,

AND EL PASO COUNTY, TEXAS

SUPPLEMENT TO OPEN-FILE REPORT 88-305

By Peter F. Frenzel

U.S. GEOLOGICAL SURVEY

Water-Resources Investigations Report 91-4155

Prepared in cooperation with the

U.S. SECTION, INTERNATIONAL BOUNDARY AND WATER COMMISSION-United States and Mexico



Albuquerque, New Mexico

U.S. DEPARTMENT OF THE INTERIOR

MANUEL LUJAN, JR., Secretary

U.S. GEOLOGICAL SURVEY

Dallas L. Peck, Director

For additional information write to:

District Chief U.S. Geological Survey Pinetree Corporate Centre Water Resources Division 4501 Indian School Rd. NE, Suite 200 Albuquerque, New Mexico 87110 Copies of this report can be purchased from:

U.S. Geological Survey Books and Open-File Reports Federal Center Box 25425 Denver, Colorado 80225

CONTENTS

	Page
Abstract	1
Introduction	2
Purpose and scope Previous investigations Well-numbering systems Acknowledgments	4 4 5 5
Description of the Mesilla Basin	8
Physical features	8 9 10 11 11
Hydrogeology of the basin-fill deposits	12 15 16
Net irrigation flux	18 21 21 22 24
Summary of geohydrology	24
Simulation of ground-water flow	25
Description of the model	25
Model gridAquifer characteristics	26 26
Hydraulic conductivity and transmissivity Leakance	29 35 36
Initial condition and time periods	36 37
No-flow boundary	37 38 41 41

CONTENTS - - Concluded

			Page
		Flow to and from the river and drains	41 42 42 42 44
Мс	odel	ad justments	44
	S	ystem properties used for comparison	44
		Surface-water depletions	45 47 47
	S	ystem properties adjusted	49
		Extinction depth	50 50 50 56 56
		s of existing withdrawals on flow in the Rio Grande	57 62
Summary	and	conclusions	64
Referen	ices.		66
		FIGURES	
Figure	1.	Map showing location of the study area	3
	2.	Diagram showing system of numbering wells in New Mexico	6
	3.	Diagram showing system of numbering wells in Texas	7
4	-6.	Graphs showing:	
		4. Average annual and April-to-October precipitation at selected stations in and near the Mesilla Valley	9
		5. Annual discharge of the Rio Grande at Leasburg and El Paso Narrows, and depletion of Rio Grande discharge	12
		6. Distribution of hydraulic conductivity with depth for the upper, middle, and lower hydrostratigraphic units of the Santa Fe Group	14

FIGURES -- Continued

			Page
Figure	7.	Diagram showing interactions between ground water and surface water in the Mesilla Valley	17
	8.	Graph showing estimate of net irrigation flux in the Mesilla Valley, 1915-85	19
	9.	Graph showing estimated net ground-water withdrawal rates for nonirrigation uses in the Mesilla Valley	23
	10.	Map showing model grid	27
:	11.	Diagrammatic sections along row 22 and column 17 showing layer thicknesses and relation between hydrostratigraphic units and layers	28
12-	16.	Maps showing:	
		12. Hydraulic conductivity assigned to model layer 1, the top layer	30
		13. Transmissivity assigned to model layer 2	31
		14. Transmissivity assigned to model layer 3	32
		15. Transmissivity assigned to model layer 4	33
		16. Recharge, underflow, and evapotranspiration boundaries for layer 1, steady-state simulation	39
	17.	Graphs showing major basinwide flow rates	40
	18.	Map showing model-derived evapotranspiration from nonirrigated lands for 1975 and area of net irrigation flux	43
	19.	Graph showing relation of annual surface-water depletions calculated from measured discharges to depletions calculated from model-derived discharges	45
2	20.	Graph showing relation of cumulative surface-water depletions calculated from measured discharges to cumulative depletions calculated from model-derived discharges	46
:	21.	Hydrographs showing measured and model-derived drain discharges	48
2	22.	Map showing measured hydraulic heads and model-derived steady-state potentiometric surface, layer 1	5 1

FIGURES - - Concluded

			Page
Figure	23.	Map showing measured hydraulic heads and model-derived potentiometric surface for 1975-76, layer 1	52
	24.	Map showing measured hydraulic heads and model-derived potentiometric surface for 1984-85, layer 1	53
	25.	Hydrographs showing measured and model-derived hydraulic heads	55
	26.	Graphs showing effect of changes in diffusivity on the source of water for nonirrigation withdrawals	59
		TABLES	
Table	1.	Estimated values of hydraulic conductivity of the Santa Fe Group	69
	2.	Values used for estimating net irrigation flux	72
	3.	Thicknesses of hydrostratigraphic units assigned to each model layer, in feet	77
	4.	Specified inflow to river reaches by stress period	136
	5.	Schedule of nonagricultural ground-water withdrawal specified in the model	137
ı	6.	Differences between measured and model-derived hydraulic heads	145
	7.	Average differences, in feet, between model-derived and measured hydraulic heads for the standard and each sensitivity test	151

CONVERSION FACTORS AND VERTICAL DATUM

Multiply	<u>By</u>	<u>To obtain</u>
inch	25.40	millimeter
inch	0.02540	meter
foot	0.3048	meter
foot per day	0.3048	meter per day
foot per year	0.3048	meter per year
per foot	3.281	per meter
square foot	0.0929	square meter
foot squared per day	0.0929	meter squared per day
cubic foot	0.02832	cubic meter
cubic foot per second	0.02832	cubic meter per second
cubic foot	28.32	cubic decimeter
cubic foot per second	28.32	liter per second
cubic foot per second	28.32	cubic decimeter per second
foot per mile	0.1894	meter per kilometer
mile	1.609	kilometer
acre	4,047	square meter
acre	0.4047	square hectometer
acre	0.004047	square kilometer
square mile	2.590	square kilometer
gallon	3.785	liter
gallon	3.785	cubic decimeter
gallon	0.003785	cubic meter
gallon per minute	0.06309	liter per second
gallon per minute	0.06309	cubic decimeter per second
gallon per minute	0.00006309	cubic meter per second
gallon per minute per foot	0.2070	liter per second per meter
acre-foot	1,233	cubic meter
acre-foot	0.001233	cubic hectometer
acre-foot	0.000001233	cubic kilometer
acre-foot per year	0.001233	cubic hectometer per year
acre-foot per acre	0.003048	cubic hectometer per hectare

<u>Sea level</u>: In this report, "sea level" refers to the National Geodetic Vertical Datum of 1929--a geodetic datum derived from a general adjustment of the first-order level nets of the United States and Canada, formerly called Sea Level Datum of 1929.

SIMULATION OF GROUND-WATER FLOW IN THE MESILLA BASIN, DOÑA ANA COUNTY, NEW MEXICO, AND EL PASO COUNTY, TEXAS

SUPPLEMENT TO OPEN-FILE REPORT 88-305

By Peter F. Frenzel

ABSTRACT

The Santa Fe Group of Tertiary age and flood-plain alluvium constitute the major aquifer of the Mesilla Basin. The lateral extent and depth of the ground-water system are defined by bedrock that has much smaller values of hydraulic conductivity than those of the flood-plain alluvium and deposits of the Santa Fe Group. Most flow into and out of the ground-water system occurs at or near land surface in the Mesilla Valley and is the result of a complex interaction of the river, drains, canals, evapotranspiration, and water withdrawals from wells. However, a relatively small amount of recharge results from torrential surface runoff, mainly near the mountain fronts.

Anticipated ground-water withdrawals in the Mesilla Basin may reduce flow in the Rio Grande and thus reduce the quantity of water that is available to downstream users. A previous model (Frenzel, P.F., and Kaehler, C.A., 1990, Geohydrology and simulation of ground-water flow in the Mesilla Basin, Doña Ana County, New Mexico, and El Paso County, Texas, with a section on Water quality and geochemistry by S.K. Anderholm: U.S. Geological Survey Open-File Report 88-305; also to be available as U.S. Geological Survey Professional Paper 1407-C, in press) was revised to assess the effects of existing ground-water withdrawals on flow in the Rio Grande and to provide a tool for the assessment of the effects of future ground-water withdrawals. The revisions included a 10-year extension, from 1975 through 1985, of the original simulation period, a reduction in the simulated thickness of Santa Fe Group deposits based on recent geologic interpretations, and minor changes in simulated recharge and evapotranspiration based on additional data.

The revised ground-water model is basically the same as the previous model with some exceptions. The bottom layer (layer 5) was discarded because the basin is now considered to be shallower than was thought previously. Refined estimates of agricultural evapotranspiration averaged 2.4 acre-feet per acre, 0.2 acre-foot per acre per year more than those of Frenzel and Kaehler (1990) as a result of adding estimates of winter evapotranspiration for particular crops. The main difference between this model and the previous model is the way in which transmissivity is estimated. Whereas transmissivity of the previous model was estimated on the basis of depth below land surface and, in a general way, on lithology, transmissivity in this model primarily is estimated on the basis of recent definition of three hydrostratigraphic units of the Santa Fe Group. The current specified hydraulic conductivity of the upper hydrostratigraphic unit of the Santa Fe Group is 30 feet per day. middle hydrostratigraphic unit of the Santa Fe Group was subdivided on the basis of depth below the water table. The hydraulic conductivity of the middle unit is 20 feet per day (10 feet per day at Las Cruces) at depths shallower than 600 feet below the water table and 10 feet per day (5 feet per day at Las Cruces) at depths below 600 feet. The hydraulic conductivity of the lower unit is 13 feet per day and the lower unit was absent at Las Cruces.

About two-thirds of the simulated nonirrigation ground-water withdrawals were estimated to be derived from increased depletion of the Rio Grande. About one-sixth was estimated to have been derived from depletion of aquifer storage and about one-sixth derived from salvaged evapotranspiration. These proportions are valid for most of the time since the late 1950's, after which time simulated nonirrigation withdrawals ranged from about 35 cubic feet per second (1958-60) to about 60 cubic feet per second (1980-86). These results were not very sensitive to changing hydraulic diffusivity by a factor of 2.25, but could be very sensitive if a well field were located distant from the Mesilla Valley.

INTRODUCTION

The Mesilla Basin is located mostly in Doña Ana County, extending into El Paso County, Texas, and the State of Chihuahua, Mexico (fig. 1). The river alluvium and Santa Fe Group deposits of the Mesilla Basin are in direct hydraulic connection with the Rio Grande. Future ground-water withdrawals in the Mesilla Basin may reduce flow in the Rio Grande and thus reduce the quantity of water that is available to make deliveries to downstream users. The ground-water hydrology and geochemistry of the Mesilla Basin in central-southern New Mexico and western Texas had been studied as part of the Southwest Alluvial Basins Regional Aquifer-System Analysis program of the U.S. Geological Survey. That study produced a ground-water flow model of the Mesilla Basin (Frenzel and Kaehler, 1990; also to be available as U.S. Geological Survey Professional Paper 1407-C, in press) that simulated the effects of ground-water withdrawals on flow in the Rio Grande. The U.S. Geological Survey, in cooperation with the International Boundary and Water Commission, conducted the study described here to revise the Frenzel-Kaehler ground-water flow model to reassess the effects of ground-water withdrawals in the Mesilla Basin on flow in the Rio Grande.

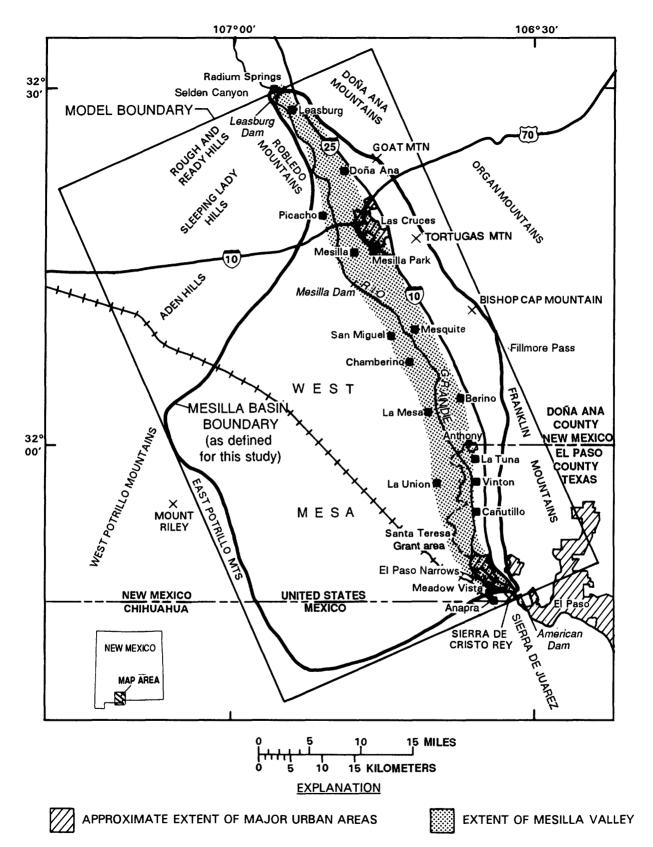


Figure 1.--Location of the study area.

Purpose and Scope

This report describes the revision of the Frenzel-Kaehler (1990) model and the assessment of the effects of existing ground-water withdrawals on flow in the Rio Grande. A companion report (Frenzel, 1992) lists model-input values on magnetic diskette for the revised model.

The revision includes a 10-year extension, from 1975 through 1985, of the original simulated historical period and the modification of the model to reflect recent geological interpretations concerning the depth and stratigraphy of the Santa Fe Group (J.W. Hawley and R.P. Lozinsky, New Mexico Bureau of Mines and Mineral Resources, written commun., 1988). Extension of the simulated period required assembling ground-water-withdrawal and streamflow data and preparing model input for additional stress periods. Modifying the model to reflect recent geological interpretations mainly involved changing the way in which model-layer transmissivity values and vertical-conductance values were derived, and adjusting the simulated thickness of basin fill. Other improvements were minor, such as estimating agricultural evapotranspiration on the basis of reported acreages of individual crops rather than on the basis of an assumed crop-type distribution for the 1950's and 1960's and estimating winter crop evapotranspiration.

As a supplement to the report by Frenzel and Kaehler (1990), this report describes the changes made to the model, and gives the reasons for the changes and the estimates of model input. The general description of the Mesilla Basin, its geohydrology, and water use are an abridged version of Frenzel and Kaehler's (1990) discussion with some additional information. The description of the model in "Simulation of ground-water flow" closely parallels that of Frenzel and Kaehler (1990) and describes in detail the differences between this model and the Frenzel-Kaehler model. In addition, this report describes the estimated effects of existing ground-water withdrawals on flow in the Rio Grande.

Previous Investigations

Ground-water investigations in the Mesilla Valley began in the early part of the 20th century. Most were reviewed by Frenzel and Kaehler (1990) who presented an overview of the geology and hydrology of the basin and a digital ground-water flow model. A section on geochemistry in the Mesilla Basin by S.K. Anderholm also was included in Frenzel and Kaehler (1990).

Recent works were not reviewed by Frenzel and Kaehler (1990). Wilson and White (1984) presented aquifer-test data for the central Mesilla Valley, and Myers and Orr (1985) presented aquifer-test data for the northern West Mesa area. A report by Hernandez and others (1987) included estimates of municipal water use for 1984. Nickerson (1989) reported aquifer-test data based on stage changes in the Rio Grande. Ground-water flow-model studies include Gates and others (1984) and Peterson and others (1984). J.W. Hawley and R.P. Lozinsky (written commun., 1988) reviewed electric logs, identifying upper, middle, and lower hydrostratigraphic units (informal usage) of the Santa Fe Group, and found that the Mesilla Basin is shallower than was thought previously. As aquifer characteristics are determined for these individual hydrostratigraphic units, Hawley and Lozinsky's conclusions may be of major significance in the estimation of the basinwide effects of ground-water withdrawals.

Well-Numbering Systems

Wells in New Mexico are identified by a location-number system based on the township-range system of subdividing public lands. The location number consists of a series of digits, separated by periods, corresponding to the township, range, section, and tract within a section (fig. 2). The townships and ranges are numbered according to their location relative to the New Mexico base line and the New Mexico principal meridian. The smallest division, represented by the third digit of the final number, is a 10-acre tract. If a well has not been located precisely enough to be placed within a particular section or tract, a zero is used for that part of the location number.

Wells in Texas are identified by a number consisting of five parts (fig. 3). The first part is a two-letter prefix used to identify the county. El Paso County is represented by JL, whereas NM signifies wells in New Mexico that have been given a number using the Texas system. The second part of the number is two digits indicating the l-degree quadrangle. Each l-degree quadrangle is divided into 64 7 1/2-minute quadrangles. This is the third part of the well number. The first digit of the fourth part indicates the 2 1/2-minute quadrangle and the last two digits comprise a sequence number that identifies the well from others in the same 2 1/2-minute quadrangle. The Texas wells used in the study also were given a latitude-longitude location number to aid data processing.

<u>Acknowledgments</u>

Many people contributed to this study. During the data collection phase, personnel of the U.S. Bureau of Reclamation office in El Paso, Texas, provided records of irrigated crop acreages and precipitation. Records of irrigated acreages were also provided by personnel of the Elephant Butte Irrigation District and El Paso County Water Improvement District. Personnel of the U.S. Section, International Boundary and Water Commission provided precipitation records and helped find additional records of irrigated crop acreages. Owen Lockwood of New Mexico State University supplied ground-water-withdrawal and well-construction data. City of Las Cruces personnel provided city ground-water withdrawals and return-flow information. Personnel of the Jornada Water Company and the Doña Ana Mutual Domestic Water Consumers Association contributed updated ground-water-withdrawal data. Discussions about the geology of the Mesilla Basin with J.W. Hawley of the New Mexico Institute of Mining and Technology were valuable.

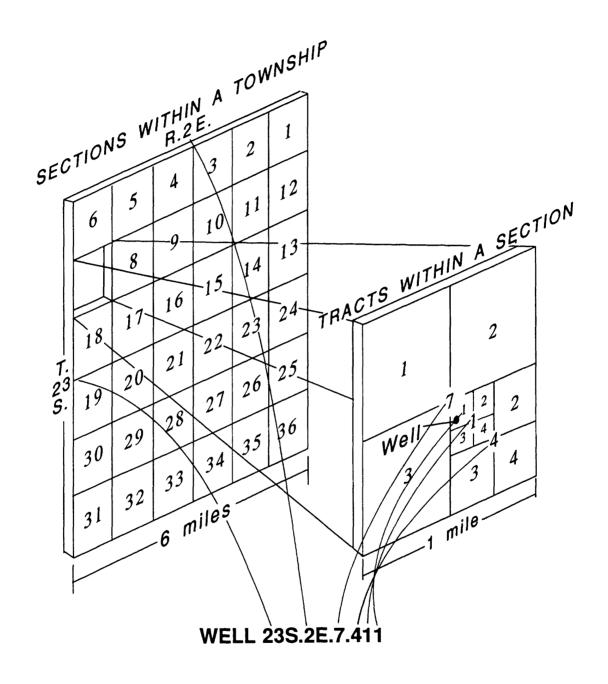


Figure 2.--System of numbering wells in New Mexico.

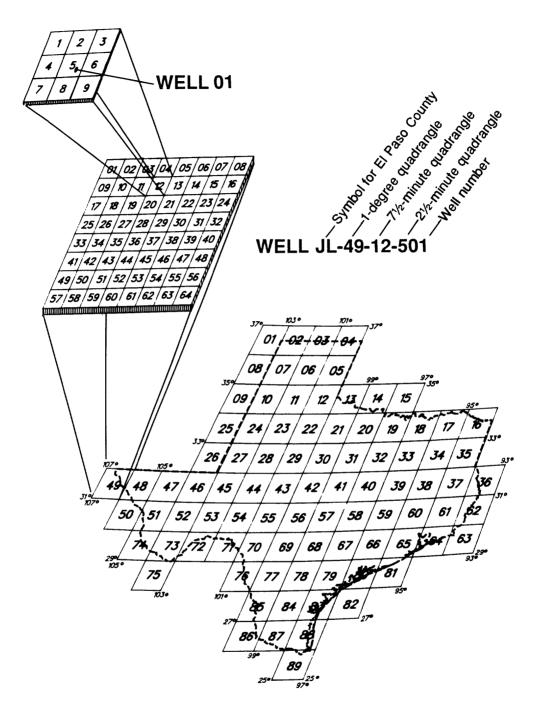


Figure 3.--System of numbering wells in Texas.

DESCRIPTION OF THE MESILLA BASIN

The following description is taken from Frenzel and Kaehler (1990). More recent information given in this section generally foreshadows a corresponding change in the updated model. Many of the physical features of the basin are shown in figure 1.

Physical Features

The Mesilla Basin is encircled by mountains. The East and West Potrillo Mountains, Aden Hills, and Sleeping Lady Hills are on the west. The Robledo Mountains and Doña Ana Mountains are on the north. Goat Mountain, Tortugas Mountain, the Organ Mountains, Bishop Cap Mountain, and the Franklin Mountains are on the east; on the southeast, the Sierra de Cristo Rey is on the international boundary and the Sierra de Juarez is just south of the International boundary. The highest point in Doña Ana County is in the Organ Mountains (9,012 feet above sea level), whereas mountains on the west side of the basin have relatively low relief (high point 5,979 feet).

The Rio Grande enters the basin through Selden Canyon between the Ro do and Doña Ana Mountains and exits through El Paso Narrows between the Franklin Mountains and the Sierra de Cristo Rev. broad area west of the Rio Grande from the international boundary to the Rough and Ready Hills is called the West Mesa (or La Mesa). The West Mesa is relatively level, has some closed drainage basins, and generally slopes slightly toward the southeast. The area contains scattered extinct volcanoes and an extensive area of lava flows. Incision of the Rio Grande into the mesa surface by as much as 400 feet has formed the Mesilla Valley, which is more than 50 miles long and as much as 5 miles wide. The total area of the Mesilla Valley is about 110,000 acres. The altitude of the Mesilla Valley ranges from 3,980 feet at Leasburg Dam to 3,729 feet at El Paso Narrows. East of the Mesilla Valley, the land slopes upward toward the Organ and Franklin Mountains (Frenzel and Kaehler, 1990, p. 9).

The basin, as defined for this report (fig. 1) is not a surface-water drainage area but is

* * * bounded by uplifted blocks of bedrock or by relatively impermeable volcanic rocks and is filled with alluvial sediment from the surrounding mountains and with fluvial sediment carried in by an ancestral Rio Grande. The area between the basin and the mountains generally is covered by similar but relatively thinner sediments. * * * Gravity anomalies and bedrock outcrops indicate that the basin extends about 7 miles into Mexico. However, no hydrologic data were available for that area (Frenzel and Kaehler, 1990, p. 9).

Climate

The climate is dry. Pan evaporation averages about 94 inches per year. The average annual precipitation is about 8 inches. Average annual and Aprilto-October (growing season) precipitation is shown in figure 4. The precipitation values shown in figure 4 were calculated as the mean of values reported for several sites. In addition to the stations at Las Cruces and La Tuna, which were used by Frenzel and Kaehler (1990), nine stations scattered throughout the Mesilla Valley operated since 1975 by the U.S. Bureau of Reclamation were used as well as a station at American Dam operated since 1938 by the U.S. Section of the International Boundary and Water Commission. The Bureau of Reclamation stations are located at Leasburg; Picacho, about 5 miles northwest of Las Cruces; Las Cruces; San Miguel; Mesquite; Chamberino; Anthony; La Union; and Cañutillo.

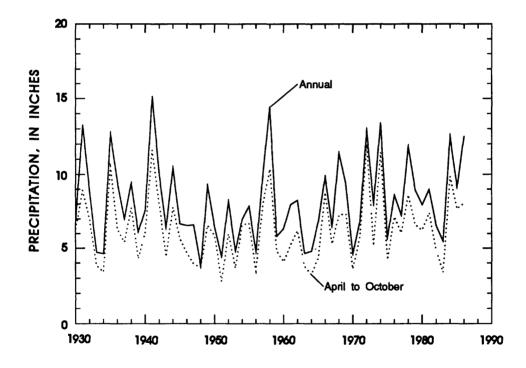


Figure 4.--Average annual and April-to-October precipitation at selected stations in and near the Mesilla Valley.

Population, Industry, and Agriculture

Frenzel and Kaehler (1990) discussed the population of the Mesilla Basin and its reliance on ground water.

The main population center within the basin is Las Cruces, which had a population of 8,325 in 1940, 29,367 in 1960, and 45,086 in 1980 (U.S. Department of Commerce, 1952, p. 31-8; 1982, p. 33). Other population centers in New Mexico include Doña Ana, Mesilla, Mesilla Park, San Miguel, La Mesa, Berino, Chamberino, Anthony, La Union, Meadow Vista, Anapra, and residential developments in the Santa Teresa Grant area * * *. New Mexico State University is adjacent to Las Cruces. In the Texas part of the valley are Anthony, Vinton, and Cañutillo. All these communities are dependent upon ground water for domestic and industrial water supply. El Paso, Texas, which had a 1980 population of almost 0.5 million, is mainly southeast of the basin and depends partly on ground water pumped from a well field at Cañutillo. Industry in the area includes produce canneries, an egg-production plant, a steel mill, and a knitting mill.

Agriculture is a major activity, and irrigation is the chief use of water in the Mesilla Basin. The Rio Grande is the primary source of irrigation water * * *. Surface water is supplemented by ground water. As of February 1948, there were about 70 irrigation wells in the Rincon and Mesilla Valleys (Conover, 1954, p. 107). During the drought of 1951-57, several hundred wells were drilled in the Mesilla Valley. Many wells were also drilled during a shortage of surface water from 1963 to 1966. As of 1975 there were about 920 usable irrigation wells in the Mesilla Valley (C.A. Wilson, U.S. Geological Survey, oral commun., 1979). The irrigation wells. mostly about 100 feet deep, primarily were completed in the floodplain alluvial aquifer. However, after 1975, a large number of irrigation wells were drilled deeper in order to obtain water of better quality than that available from shallow wells (Wilson and White, 1984). * * * Major crops * * * are cotton, pecans, alfalfa, cereal grains, and vegetables * * *. The number of irrigated acres in Mesilla Valley has increased from about 25,000 acres near the turn of the century to about 77,000 acres during 1940-75 * * * which is about two-thirds of the area of the valley. Much of the West Mesa area is used for grazing, where there are scattered stock and domestic wells.

Historical irrigation practices have effectively used the ground-water system as a reservoir in a combined stream-aquifer system. During years of plentiful surface water, most irrigation water is diverted from the Rio Grande. About one-third of applied irrigation water may replenish the ground-water system (Blaney and Hanson, 1965, p. 28). Some ground water seeps into drains that discharge to the Rio Grande. During years of inadequate surface-water supply, the shortfall is made up from ground water, causing lower than usual ground-water levels and diminished drain discharge. Ground-water levels generally return to normal after an irrigation season when surface water is plentiful (Frenzel and Kaehler, 1990, p. 11 and 14). [Withdrawal rates for irrigation wells are unknown.]

Surface-Water System

The surface-water system is comprised of the Rio Grande and its tributaries, a distribution system of canals and laterals, and drainage ditches that return water to the Rio Grande (Frenzel and Kaehler, 1990. pl. 1). Flow in the Rio Grande is almost completely controlled by upstream The channel of the Rio Grande is defined by levees in the Mesilla reservoirs. Basin. Several arroyos (ephemeral streams with straight-walled, flat-bottomed channels) flow into the Rio Grande mainly from the mountains on the east side of the basin. Flow in some of the large arroyos is blocked by retention dams near Las Cruces and El Paso. Flow in other arroyos reaches the valley, but probably does not contribute much flow directly to the discharge of the Rio The convergence of surface flow from time to time into arroyos where Grande. it can rapidly infiltrate deep into the sand may provide the principal mechanism for natural recharge of ground water. This is the essential mechanism that Frenzel and Kaehler (1990, p. 41-42) termed "mountain-front" or "slope-front" recharge, which are similar except for location. Mountain- and slope-front recharge together may be 15 cubic feet per second, but the error of this estimate may be very large (plus 100 percent or minus 50 percent (Frenzel and Kaehler, 1990, p. 42). Diversions from the Rio Grande in the Mesilla Basin are at Leasburg and Mesilla Dams. The canal system distributes diverted water to agricultural fields and the drain system returns unused water to the Rio Grande. Depletions of the Rio Grande are calculated as the discharge at Leasburg Dam near the north end of the Mesilla Basin minus the discharge at El Paso Narrows. For 71 years, 1915-85, annual depletions have averaged 218,000 acre-feet. Annual discharges and depletions for this period are shown in figure 5. Annual discharges and depletions for 1975-85, in acrefeet, are:

	Discharge		
Year	At Leasburg Dam	At El Paso Narrows	Depletion
1975	591,400	360,959	230,441
1976	663,710	402,835	260,875
1977	398,319	214,553	183,766
1978	338,819	156,024	182,795
1979	563,304	312,594	250,710
1980	631,380	353,983	277,397
1981	574,872	333,329	241,543
1982	599,051	326,642	272,409
1983	599,714	331,955	267,759
1984	630,894	359,361	271,533
1985	645,911	359,917	285,994

Ground-Water System

The ground-water system is controlled by the basin-fill geology, geologic boundaries, and the interrelation of ground and surface waters (Frenzel and Kaehler, 1990, p. 19). The interrelation of ground and surface waters occurs mainly in the Mesilla Valley and was termed "Mesilla Valley boundary" by Frenzel and Kaehler (1990). The geologic boundaries of the basin are formed

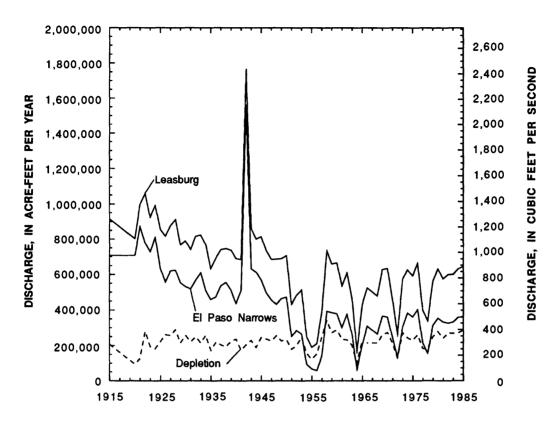


Figure 5.--Annual discharge of the Rio Grande at Leasburg and El Paso Narrows, and depletion of Rio Grande discharge.

by consolidated bedrock units that are generally much less permeable than the basin-fill sediments of the Santa Fe Group and flood-plain alluvium. The Santa Fe Group of Tertiary age and flood-plain alluvium constitute the major aquifer. Frenzel and Kaehler (1990, p. 19-29) described the bedrock, Santa Fe Group, and flood-plain alluvium based on the geologic literature of the late 1970's and early 1980's.

Hydrogeology of the Basin-fill Deposits

Since the report of Frenzel and Kaehler (1990), the description of the subsurface geology of the basin has been expanded by Hawley and Lozinsky (written commun., 1988) to include three "hydrostratigraphic" units of the Santa Fe Group. The lower unit is mainly composed of uniform fine sand, whereas the middle and upper units are lenticular and with grain sizes ranging from clay to gravel. The lower hydrostratigraphic unit is about 600 feet thick, ranging from 400 to about 1,000 feet in thickness where it is fully penetrated by wells (Hawley and Lozinsky, written commun., 1988). Because the lower unit is easily distinguishable from the middle unit on electric logs (Hawley and Lozinsky, written commun., 1988), the depth of the base of the Santa Fe Group can be estimated to within a few hundred feet using wells that do not fully penetrate the lower unit. The maximum thickness of the Santa Fe Group is now considered to be about 2,500 feet, much less than the 3,800 feet that it was previously thought to be (Frenzel and Kaehler, 1990, p. 32).

The division of the Santa Fe Group into three hydrostratigraphic units facilitates the assignment of aquifer properties on the basis of stratigraphy instead of strictly on the basis of depth. Although the lower unit consists of more uniform sand, cementation may cause its horizontal hydraulic conductivity to be less than that of the other two units. The middle hydrostratigraphic unit may be somewhat less permeable than the upper unit on the basis of grain size and degree of cementation. The upper hydrostratigraphic unit is saturated mainly in the northern part of the basin and may be the most permeable on the basis of grain size and relative lack of cementation. Horizontal-to-vertical anisotropy is probably greatest in the middle and upper units because of lenticularity (Hawley and Lozinsky, written commun., 1988).

Each estimated hydraulic conductivity in table 1 (tables are in back of the report) was assigned to one of the three hydrostratigraphic units on the basis of the depth tested. The depths of the hydrostratigraphic units at each well site were estimated from plates 16 and 17 of Hawley and Lozinsky (written commun., 1988). Most of the estimated values of hydraulic conductivity in table 1 are from table 1 of Frenzel and Kaehler (1990). Additional values were estimated from values of transmissivity divided by open intervals reported in references given in table 1.

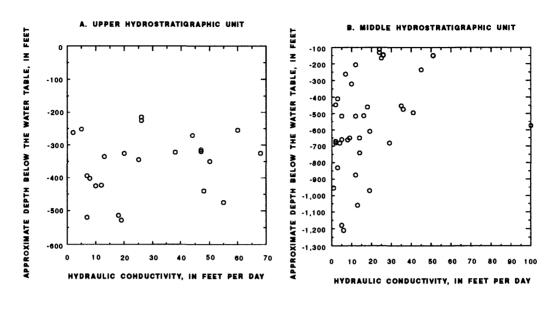
The relation of hydraulic conductivity with depth that had been apparent in figure 13 of Frenzel and Kaehler (1990) was no longer apparent (fig. 6A and 6C) for the upper and lower hydrostratigraphic units of the Santa Fe Group. However, an inverse relation of hydraulic conductivity with depth was still apparent in figure 6B for the middle hydrostratigraphic unit.

For the lower unit, estimated values of hydraulic conductivity given in table 1 range from 1 to 34 feet per day. Of 12 values, the lower quartile is between 2 and 3, the middle quartile (median) is between 11 and 14, and the upper quartile is 17 feet per day.

For the middle unit, estimated hydraulic conductivity ranges from 1 to 100 feet per day (table 1). Of 38 values, the lower quartile is 5, the middle quartile (median) is between 13 and 14, and the upper quartile is 25 feet per day. However, because of the apparent inverse correlation between hydraulic conductivity and depth (fig. 6B), it may be reasonable to divide these data on the basis of depth.

Estimated values of hydraulic conductivity for the middle hydrostratigraphic unit were divided into two groups on the basis of depth below the water table (table 1). Estimated hydraulic conductivity for depths between 600 and 1,210 feet ranges from 1 to 29 feet per day. Of 18 values, the lower quartile is 4, the middle quartile (median) is between 8 and 9, and the upper quartile is 14 feet per day. Estimated hydraulic conductivity for depths shallower than 600 feet ranges from 2 to 100 feet per day. Of 20 values, the lower quartile is between 10 and 12, the middle quartile is 24, and the upper quartile is between 35 and 36 feet per day.

For the upper unit, estimated hydraulic conductivity ranges from 2 to 68 feet per day (table 1). Of 23 values, the lower quartile is between 10 and 12, the middle quartile is 25, and the upper quartile is 47 feet per day.



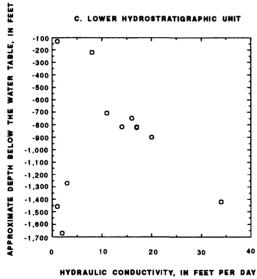


Figure 6.--Distribution of hydraulic conductivity with depth for the upper, middle, and lower hydrostratigraphic units of the Santa Fe Group.

In the Las Cruces area, faults may cause a reduction in the overall transmissivity of the upper and middle hydrostratigraphic units of the Santa Fe Group. Plate 1 of Hawley (1984) shows many more faults in the Las Cruces area than in other parts of the Mesilla Basin. Faults may cause sand or gravel lenses to be juxtaposed against clay lenses, causing a reduction in transmissivity. Also, fault planes have been observed in the Las Cruces area that were more cemented than the rest of the formation (J.W. Hawley, New Mexico Bureau of Mines and Mineral Resources, oral commun., 1989).

The flood-plain alluvium is about 4 miles wide along the Rio Grande, and is much more permeable than the Santa Fe Group deposits. Wilson and others (1981, pl. 11) reported values of transmissivity of 10,000 to 30,000 feet squared per day for the upper 150 feet of the aquifer, which is composed of about 80 feet of river alluvium overlying about 70 feet of the Santa Fe Group. If the 70-foot thickness of Santa Fe Group deposits is assumed to have a hydraulic conductivity of about 25 feet per day, it then has a transmissivity of 1,750 feet squared per day, and by subtraction, the transmissivity of the river alluvium is between 8,250 and 28,250 feet squared per day. Dividing transmissivity values by the 80-foot thickness yields hydraulic-conductivity values for the river alluvium ranging from about 100 to 350 feet per day.

Under water-table conditions, specific yield may be about 0.2 as indicated by previous studies (Richardson and others, 1972, p. 86; Lizarraga, 1978, p. 29; Wilson and others, 1981, fig. 10). Under fully saturated conditions, the specific storage may be about 1×10^{-6} per foot of thickness. This value is common in sedimentary rocks (Lohman, 1972, p. 8) and accounts for expansion of water and some elastic matrix compression.

Hydrogeologic Boundaries

The Mesilla Basin (fig. 1) is generally surrounded by a pediment-like structure where the bedrock is relatively shallow between the basin and the surrounding mountains (Frenzel and Kaehler, 1990, p. 29-33). In the following discussion the pediment-like structure will be called a "pediment" to signify the general concept even though it is not a true pediment in all locations around the basin. Overlying the pediment are Santa Fe Group deposits. The saturated interval of these deposits is thin relative to the saturated thickness of Santa Fe Group deposits within the basin.

A small quantity of water probably enters the Mesilla Basin around its boundaries. Underflow into the basin from the Santa Fe Group deposits overlying the pediments probably occurs although it is probably small compared to the quantity of water involved in the interaction of ground water with surface water near the axis of the basin (Frenzel and Kaehler, 1990, p. 29). Most mountain-front recharge probably enters the basin as underflow from Santa Fe Group deposits overlying the pediments. A much smaller amount of water may enter the basin from the bedrock surrounding the basin.

A small quantity of water also enters the basin as underflow in the river alluvium near Radium Springs and a similarly small quantity, estimated by Slichter (1905, p. 9-13) to be 0.1 cubic foot per second, leaves the basin in the river alluvium in El Paso Narrows. These quantities are small because of the small cross-sectional area of the river alluvium.

A small, unknown quantity of water probably leaves the basin through Fillmore Pass. The east boundary of the basin generally is a recharge boundary (Frenzel and Kaehler, 1990, fig. 18) and mountain-front recharge enters the basin largely as underflow along the pediments. Along most of the east boundary, a steep water-table gradient toward the basin exists. Almost no gradient exists between Fillmore Pass and the Mesilla Valley, but a steep gradient does exist toward the east from Fillmore Pass (Wilson and others, 1981, pl. 9). If no water flowed eastward through Fillmore Pass, some gradient toward the Mesilla Basin from Fillmore Pass might exist because of mountain-front recharge north and south of the pass. Possibly some of the mountain-front recharge in the vicinity of Fillmore Pass flows eastward through the pass and some flows westward toward the Mesilla Basin.

Mesilla Valley Boundary

Most ground-water recharge and discharge in the Mesilla Basin occurs in the Mesilla Valley. The interaction of ground and surface water is complex in the Mesilla Valley and has been referred to as the "Mesilla Valley boundary" by Frenzel and Kaehler (1990, p. 60). A flow chart for the Mesilla Valley boundary is shown in figure 7.

The following excerpt taken from Frenzel and Kaehler (1990) describes the relation of ground water and surface water:

The ground and surface waters in the Mesilla Valley are closely interrelated. In addition to mountain- and slope-front recharge, recharge takes place when excess irrigation water percolates through the flood-plain alluvium, as shown by the rise in water levels during the summer irrigation season (Spiegel, 1958). After long periods of heavy pumping, ground-water levels rapidly recover when surface water is applied on agricultural lands (Taylor, 1967).

Ground water is discharged to drains and eventually flows back to the Rio Grande when ground-water levels rise above the drain bottoms. However, during periods of drought, the limited surfacewater supply is supplemented by ground-water withdrawals from wells, and drain discharge is decreased (King and others, 1971, p. 57). Ground water also is discharged by evapotranspiration from crops and natural vegetation in the valley.

A schematic diagram shown in figure 15 [fig. 7 in this report] indicates the following ground-water/surface-water interactions:

(1) A major group made up of (a) net diversions, (b) effective rainfall on both irrigated and nonirrigated lands (effective rainfall is here defined as that part of rainfall that either recharges aquifers or reduces ground-water discharge), (c) evaporation from canal surfaces, and (d) evapotranspiration from irrigated lands, is summed to provide an overall flux for the valley area * * *. This summation is termed 'net irrigation flux' in this report, and it includes by implication the leakage from irrigation canals to ground water and that part of irrigation ground-water pumpage that is not recirculated back to the ground water.

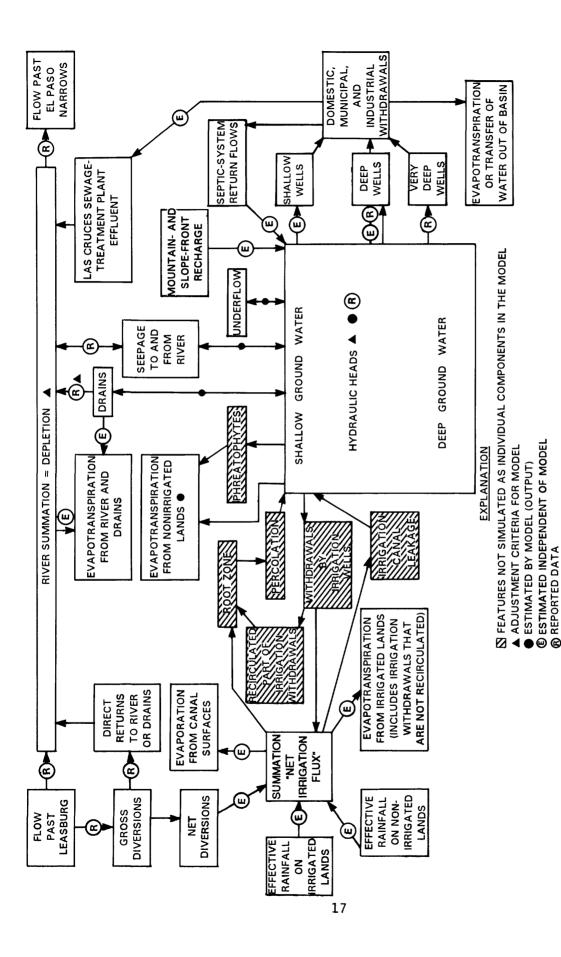


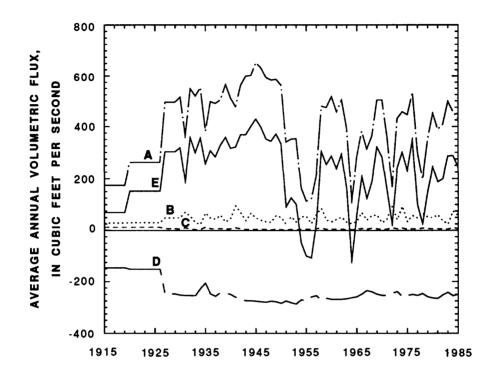
Figure 7.--Interactions between ground water and surface water in the Mesilla Valley (modified from Frenzel and Kaehler, 1990, fig. 15).

- (2) A second major group is comprised of seepage to and from the river and drains.
- (3) A third major group is evapotranspiration from nonirrigated lands.
- (4) A fourth group, of intermediate significance, consists of water pumped from wells for domestic, municipal, and industrial purposes, and septic-system return flows.
- (5) Mountain- and slope-front recharge constitutes a relatively minor amount of flow. Most flows have been reported or estimated. Most of the flows to and from ground water occur at or near the land surface in the Mesilla Valley, with the exceptions of discharges from deep wells and mountain-front recharge. These flows fluctuate seasonally, in the short term, but in the intermediate term (1-5 years), they fluctuate with the availability of surface water, and in the long term, they do not fluctuate much at all (Frenzel and Kaehler, 1990, p. 34).

Some of the flow estimates have been revised on the basis of more complete data. Estimates of net irrigation flux were revised on the basis of additional precipitation data and more complete crop acreages. Additional precipitation data were used to revise recharge estimates as well. Timevarying estimates were updated from 1975 through 1985 on the basis of data for those years.

Net irrigation flux

Net irrigation flux (fig. 8 and table 2) was estimated by essentially the same procedure as used by Frenzel and Kaehler (1990, p. 37) except that the estimates of evapotranspiration and effective precipitation were revised and extended from 1975 through 1985. The net irrigation flux (fig. 8, curve E) was a summation of estimates for annual net diversions (fig. 8, curve A), effective rainfall on irrigated lands (fig. 8, curve B), effective rainfall (ground-water recharge) on nonirrigated lands (fig. 8, curve C), and evapotranspiration from irrigated lands in the Mesilla Valley (fig. 8, curve D). As defined, net irrigation flux implicitly includes all groundwater withdrawals and percolation from irrigated lands. Estimates for these individual components were not made. Estimates of annual net diversions were made for 1976-85 in the same way as estimates for previous years had been made (Frenzel and Kaehler, 1990, p. 37). Estimates of effective rainfall were made in the same way, but were revised slightly on the basis of additional rainfall data and were extended through 1985. Evapotranspiration from irrigated lands for a given year was estimated as a per-acre evapotranspiration rate times the total irrigated acreage of the Mesilla Valley for that year. The per-acre evapotranspiration rate was estimated on the basis of reported annual acreages of individual crop types for each of the years since 1949, whereas Frenzel and Kaehler (1990) based the estimation on an assumed mixture of crops. Also, some irrigated acreage on the east side of the Mesilla Valley that was not accounted for by Frenzel and Kaehler (1990) was included in the estimates.



EXPLANATION

- A NET DIVERSION LESS ESTIMATED EVAPORATION FROM CANALS OF 5.5 CUBIC FEET PER SECOND
- B EFFECTIVE PRECIPITATION ON IRRIGATED LANDS
- C EFFECTIVE PRECIPITATION ON OTHER LANDS
- D ESTIMATED EVAPOTRANSPIRATION FROM IRRIGATED LANDS
- E SUMMATION OF A, B, C, AND D, TERMED "NET IRRIGATION FLUX"

Figure 8.--Estimate of net irrigation flux in the Mesilla Valley, 1915-85.

In addition to the main part of the Mesilla Valley that is served by a canal system to distribute surface water (Bureau of Reclamation project area) there is a narrow strip of land on the east side of the valley that is irrigated exclusively with ground water. Although records of irrigated acreage outside the Bureau of Reclamation project area generally are not kept, a few estimates have been made from aerial photography. In 1967, this irrigated acreage was about 2,800 acres (New Mexico State Engineer Office, written commun., 1979); in 1969 it was about 3,200 acres (estimated from data in Lansford and others, 1974, table 16); and in 1983, it was about 3,400 acres. This area may not be significant to the total irrigated acreage; however, the evapotranspiration from this acreage (possibly about 10 cubic feet per second) is significant compared to municipal withdrawals in the Mesilla Basin (about 60 cubic feet per second). Irrigation of this area may have begun in the 1950's, when many of the irrigation wells in the Mesilla Valley were drilled. An average of 3,000 acres, beginning in 1958, was assumed (table 2).

Similar to Frenzel and Kaehler (1990), the Blaney-Criddle method was used (Blaney and Hanson, 1965) to estimate annual evapotranspiration rates (acrefeet per acre) for each crop. However, winter transpiration by crops such as alfalfa, pasture, and pecans was accounted for, thereby increasing the estimated evapotranspiration rate. The average annual evapotranspiration rate was about 2.4 acre-feet per acre of irrigated land--an increase of almost 10 percent from that of Frenzel and Kaehler (1990, p. 37), which was 2.2 acrefeet per acre.

An evapotranspiration rate for the Mesilla Valley that reflected the mixture of crop types for the year was estimated for each year since 1949. Acreage information for individual crop types generally is available for most of the Mesilla Valley for years following 1949. For 1953-67 and 1975, acreages of individual crops are reported (Bureau of Reclamation Form 7-316 or equivalent) for the Mesilla Valley part of the Elephant Butte Irrigation These reports generally include most but not all of the irrigated acreage in the Mesilla Valley. For 1949-78 and 1980-85, similar reports exist; however, the acreages are for the entire Elephant Butte Irrigation District, combining areas upstream from the Mesilla Valley. The assumption was made that the annually reported individual crop acreages represent the mixture of crops for the entire Mesilla Valley for a given year even though the total of reported individual crop acreages often does not equal the entire irrigated acreage in the Mesilla Valley for the year. The average evapotranspiration rate (table 2) for the valley for a given year was estimated as the sum of evapotranspiration for all reported individual crops divided by the sum of reported individual crop acreages for the year.

The annual evapotranspiration from irrigated land for the entire Mesilla Valley (table 2 and fig. 8, curve D) was estimated as the average evapotranspiration rate times the annual irrigated acreage. Additional details of the estimation of net irrigation flux are given in table 2.

Evapotranspiration from nonirrigated lands

Evapotranspiration from nonirrigated lands was not estimated independently of the model. As explained by Frenzel and Kaehler (1990):

* * * These lands have highly variable hydrologic properties. They include such features as paved roads that may reduce evapotranspiration and collect rainfall, allowing it to infiltrate to the water table in places. Also included are vegetated areas that lie relatively high above the water table, minimizing the availability of ground water to plants, and areas where trees grow near surface water or shallow ground water where evapotranspiration may be near a maximum. Evapotranspiration of ground water from surfaces covered by natural vegetation generally is dependent on the depth to the water table * * *.

Seepage to and from streams

Seepage to and from streams, including drains * * * accounts for a major amount of water. The Rio Grande has both gaining and losing reaches that probably change seasonally. The drains gain when ground-water levels are high but may lose in their lower reaches where they join the river or in localities where large amounts of ground water are withdrawn.

The type and amount of interconnection between streams and the aquifer depend on the stream stages and heads in the aquifer and on streambed permeability. Where the water table is below the level of the stream, surface-water recharge to the aquifer can take place; however, the rate of recharge is dependent on the head difference and the permeability of the streambed. The beds of streams that recharge ground water tend to be somewhat more plugged than the beds of streams that receive ground-water discharge because suspended sediment tends to follow the water from the stream into pores of the aquifer, a condition that does not prevail where flow is from the aquifer to the stream.

* * * (Individual drain discharges are shown, for convenience of comparison with model-derived drain discharges, in the section on model adjustment.) The accuracy of drain-discharge measurements probably is poor [accuracy not within 15 percent of the true value] because the stream gradients are low, causing problems with backwater and moss or other vegetation. An additional uncertainty is that drains may receive water from redirected surface water such as excess water from irrigation canals. All drain discharges were assumed to have come from ground water (Frenzel and Kaehler, 1990, p. 38-39).

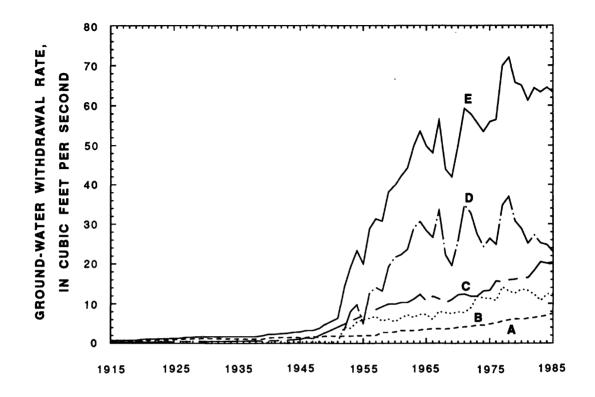
Nonirrigation withdrawals and returns

Reported and estimated ground-water withdrawals for municipal and industrial uses are shown in figure 9. Withdrawals are in four groups: curve A is for wells north of the \tilde{S} tate line at Anthony excluding the \bar{C} ity of Las Cruces, curve B is for wells south of the State line at Anthony excluding the City of El Paso, curve C is for City of Las Cruces wells, and curve D is for City of El Paso wells. Curve E is the summation of curves A-D. Groundwater withdrawals before 1976 are the same as those given by Frenzel and Kaehler (1990, p. 39-41, and fig. 17), and values for 1976-85 are reported and estimated as follows. Ground-water withdrawals at the City of Las Cruces are reported values, as are withdrawals at Doña Ana and a group of communities served by Jornada Water Company. Ground-water withdrawals at Mesquite and Anthony, New Mexico, were derived by interpolation between the 1975 estimate of Frenzel and Kaehler (1990, table 4) and values given by Hernandez and others (1987, table 9) for 1984. For these two towns, 1985 values were derived by extrapolation of the 1975-84 values. Ground-water withdrawal-data for wells south of the State line at Anthony are from the files of the U.S. Geological Survey in El Paso (Don White, U.S. Geological Survey, written commun., 1987), including withdrawals at Santa Teresa through 1984. For Santa Teresa, 1985 withdrawals were assumed to be equal to 1984 withdrawals.

Return flows of nonirrigation withdrawals are described by Frenzel and Kaehler (1990, p. 41):

Return flows of municipal, industrial, and domestic water take several forms. Water withdrawn by El Paso is not returned in the Mesilla Basin. Slightly more than one-half of the wintertime withdrawals by Las Cruces are returned as surface discharge to the Rio Grande (City personnel, oral commun., 1983). On the basis of data in Sorensen (1977, tables 1 and 2), possibly one-half of the water withdrawn by communities where no surface disposal systems are in use is returned by means of septic systems. Errors in estimates of these return flows probably are not critical because the amount is small.

Sewer-plant effluent from Anthony and La Tuna has been returned to the surface-water system via a nearby drain. Some of the ground-water withdrawals at Santa Teresa supply sprinkler-irrigation systems and some are for municipal use. Sprinkler-irrigation systems may not have a large percentage of return flow, although some return flow is required to avoid a salt buildup in the soil. Disposal of treated municipal wastewater at Santa Teresa is by way of land disposal, a method that probably causes evapotranspiration of a large percentage of the effluent.



EXPLANATION

- A RATE FOR WELLS NORTH OF THE STATE LINE AT ANTHONY
 EXCLUDING THE CITY OF LAS CRUCES WELLS
- B RATE FOR WELLS SOUTH OF THE STATE LINE AT ANTHONY
 EXCLUDING THE CITY OF EL PASO WELLS
- C RATE FOR CITY OF LAS CRUCES WELLS
- D RATE FOR CITY OF EL PASO WELLS
- E SUM OF A, B, C, AND D

NOTE: All of the rates shown were estimates of net ground-water withdrawal made exclusively for this study and may be neither complete enough nor appropriate for any other purpose

Figure 9.--Estimated net ground-water withdrawal rates for nonirrigation uses in the Mesilla Valley.

Mountain- and slope-front recharge

Recharge to the aquifer probably occurs along ephemeral streams in response to intense, local storms. Beds of ephemeral streams are composed of sand and gravel, and their relatively flat gradients allow for infiltration of water that comes from upstream, relatively steep reaches that are incised in bedrock around the perimeter of the ground-water basin. This recharge is termed mountain-front Hydrologically similar conditions occur on the sides of recharge. the Mesilla Valley where sand channels indicate that rainwater from steep slopes flows into less steep sand channels that do not extend to the river. Recharge under these conditions is termed slope-front recharge. * * * Mountain- and slope-front recharge was estimated * * * by J.D. Dewey (U.S. Geological Survey, written commun., 1983) using an empirical formula developed from a log-multiple regression analysis of measured streamflows. The method has been documented by Hearne and Dewey (1988). * * * The total mountain- and slope-front recharge was estimated to be 15.3 cubic feet per second, most recharge occurring on the eastern side of the basin. The potential error in these estimates is great--perhaps plus 100 percent or minus 50 percent. Any recharge on gently sloping areas, such as the West Mesa, probably occurs only occasionally and in small amounts. hypothesized that in such areas, where average annual rainfall (8 inches) is much less than potential evapotranspiration (50-80 inches), the natural grass and desert shrubs capture all rainfall before it percolates beyond the root zone. * * * Recharge on the West Mesa was assumed to be negligible (Frenzel and Kaehler, 1990, p. 41-42).

The effect of using the additional precipitation data in the calculation was to reduce estimated recharge by about 0.1 percent on a long-term average, an insignificant change. Although the inclusion of more data should make estimated annual recharge rates more accurate, the change in annual estimated recharge is probably not significant because short-term variations are damped out by storage in the aquifer.

Summary of Geohydrology

Frenzel and Kaehler (1990, p. 44) summarized the geohydrology of the Mesilla Basin as follows:

* * * (1) The ground-water system is three dimensional: (a) The flood-plain alluvium has a larger hydraulic conductivity than the underlying and laterally contiguous Santa Fe Group; (b) although several water-producing zones have been identified in the southern end of the valley, each with a different hydraulic head, no laterally extensive confining beds have been observed; and

(c) wells generally do not produce from the full thickness of the (2) The lateral extent and depth of the ground-water flow system are defined by bedrock with values of hydraulic conductivity that are much less than those of the Santa Fe Group. (3) Most water that flows into and out of the ground-water system is at or near the land surface in the Mesilla Valley. These flows are the result of complex interactions of the river, drains, canals, evapotranspiration, and withdrawals from wells. These flows fluctuate seasonally, in the short term, but in the intermediate term (1-5 years), they fluctuate with the availability of surface water, and in the long term (more than 5 years), they do not fluctuate much at all. (4) Basinwide withdrawals of water from deep wells (deeper than about 200 feet) before 1975 were small compared to evapotranspiration, but they were locally significant, especially in the Cañutillo and Las Cruces areas. After 1975, withdrawals from deep wells increased because more deep irrigation wells were installed. (5) A small amount of water recharges the ground-water system near mountain fronts. This water comes from surface runoff from the steep mountain drainages. By a similar mechanism, a small amount may recharge the system along the toe of the steeply sloping bluff along the west side of the Mesilla Valley. (6) Recharge of any significant amount over most of the West Mesa area is unlikely, but occasionally may occur in places.

SIMULATION OF GROUND-WATER FLOW

The ground-water flow model was essentially the same as that of Frenzel and Kaehler (1990) except for the simulated time, which was extended 10 years from 1975 to 1985, and the way in which values of transmissivity and leakance were derived. These items are discussed under the separate headings. Several minor refinements and corrections were made to the Frenzel-Kaehler model. Because it is generally similar and used the same computer program (the U.S. Geological Survey modular model by McDonald and Harbaugh, 1988, with addition of the river-routing package by Miller, 1988), this flow model is most easily described in terms of the similarities and differences between it and the Frenzel-Kaehler (1990) model.

Description of the Model

The model required specification of a three-dimensional grid, aquifer characteristics, time periods, and boundary conditions. Some values were changed during model adjustment. Aquifer characteristics reported in this section are the adjusted values.

Model Grid

Except for refinements in the vertical dimension, spacial discretization was the same as that of Frenzel and Kaehler (1990). In the horizontal direction, the grid had 36 rows and 64 columns (fig. 10), and was exactly the same as in Frenzel and Kaehler (1990, fig. 19). In the vertical dimension, layer 5 was discarded because the basin is not as deep as was formerly thought. The thicknesses of layers 1-3 were the same as those of Frenzel and Kaehler (1990) (300, 400, and 600 feet, respectively) except that the thickness of some layers was reduced near the edges of the model to better conform to the shape of the basin. Because the bottom altitude of layer 1 was defined to be 200 feet below the 1975 water table (Frenzel and Kaehler, 1990, p. 47), the saturated thickness of layer 1 was generally about 200 feet for the calculation of transmissivity. Figure 11 shows diagrammatic sections through the model along row 22 (northwest-southeast) and along column 17 (southwest-northeast). On the left side of figure 11A, the thickness of layers 1, 2, and 3 has been reduced; on the right side of figure 11A, thickness of layers 1 and 2 has likewise been reduced. Layers 1, 2, and 3 are not full thickness on the right side of figure 11B. The thickness of layer 4, the bottom layer, was variable throughout its extent. In figure 11, layer 4 varies from about 200 to 1,000 feet. The thickness of layer 4 was as much as 1,030 feet in the deepest part of the basin.

The layer thicknesses define a mixed grid (McDonald and Harbaugh, 1988, p. 2-29) that is orthogonal throughout most of the model, but deformed near the extremities and at the bottom. Model nodes were located at the center of each model block. Therefore, throughout most of layers 1-3, where model layer thickness was the nominal value, model nodes formed a uniform, orthogonal grid. However, throughout layer 4 and near the extremities of other layers, where they were less than full thickness, model nodes formed a deformed grid. As described below, vertical distances between model nodes were used in the calculation of leakances for each model block. (Leakance is vertical hydraulic conductivity divided by thickness, the term "Vcont" of McDonald and Harbaugh, 1988, p. 5-11.)

Aquifer Characteristics

Aquifer characteristics were assigned on the basis of estimated or theoretical values and then adjusted to make the model simulate measured values of hydraulic heads and stream depletions as described in the "Model adjustments" section. Aquifer characteristics are reported in this section as ranges rather than exact values because they tend to vary from block-to-block. Although aquifer characteristics could be approximated from the description given here, exact block-by-block values of aquifer characteristics are available on magnetic diskette in a companion report (Frenzel, 1992). The aquifer characteristics reported here are the most recently assigned (adjusted) values.

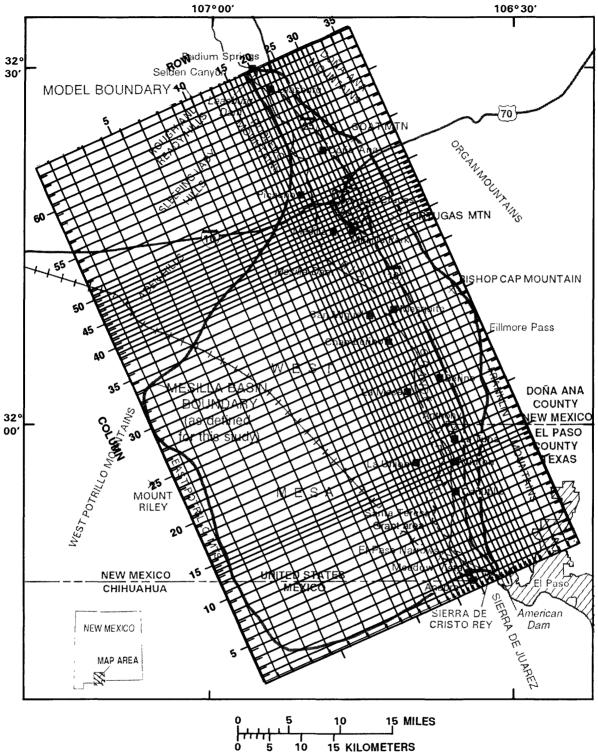


Figure 10.--Model grid.

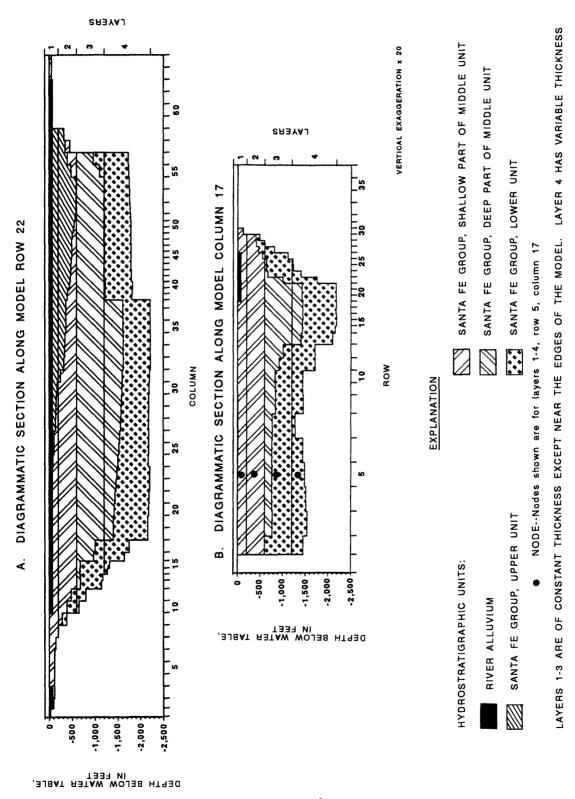


Figure 11.--Diagrammatic sections along row 22 and column 17 showing layer thicknesses and relation between hydrostratigraphic units and layers.

Hydraulic conductivity and transmissivity

Hydraulic conductivity is assigned to layer 1, from which the model estimates transmissivity on the basis of the saturated thickness (hydraulic head minus the bottom altitude). Layers 2-4 were assigned transmissivity values that were estimated from the hydraulic-conductivity values and thicknesses of hydrostratigraphic units represented in each model block.

The hydraulic conductivity assigned to each block in layer 1 (fig. 12) was a thickness-weighted average for the upper 200 feet of saturated sediments at that block. The hydraulic conductivity (K) assigned to each model block was calculated as:

$$K = \frac{TRA*KRA + TUSF*KUSF + TMSF*KMSFS}{200}$$
 (1)

where TRA = the thickness of river alluvium, in feet;

KRA = the hydraulic conductivity of river alluvium, in feet per second;

TUSF = the thickness of the upper hydrostratigraphic unit of the Santa Fe Group that is within the upper 200 feet of saturated sediments, in feet;

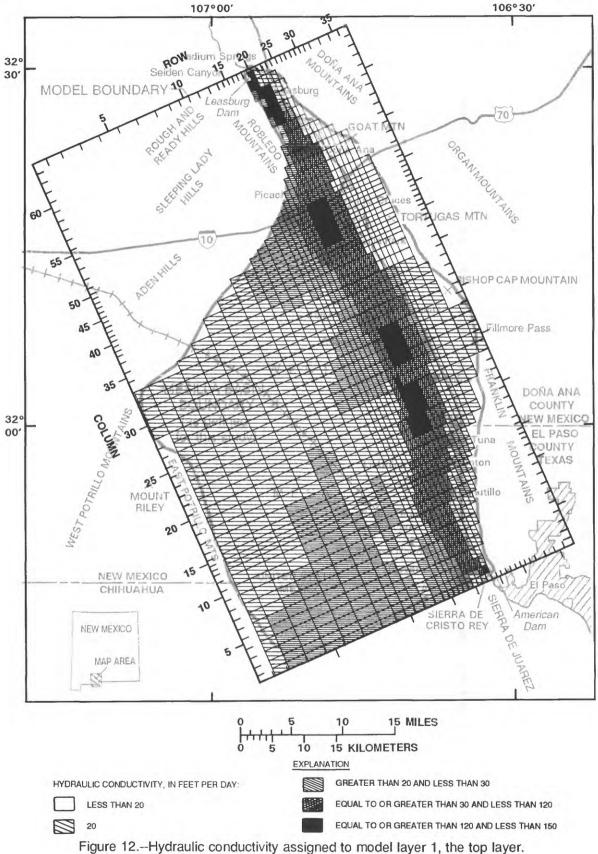
KUSF = the hydraulic conductivity of the upper hydrostratigraphic unit
 of the Santa Fe Group, in feet per second;

TMSF = the thickness of the middle hydrostratigraphic unit of the Santa Fe Group that is within the upper 200 feet of saturated sediments, in feet; and

KMSFS = the hydraulic conductivity of the shallow part of the middle hydrostratigraphic unit of the Santa Fe Group, in feet per second.

Individual thicknesses may equal zero at any given model block.

The transmissivity values of layers 2, 3, and 4 (figs. 13-15) were calculated in a similar fashion by setting the denominator of equation (1) to unity and substituting the appropriate thicknesses and hydraulic-conductivity values of the hydrostratigraphic units represented. The thicknesses of each hydrostratigraphic unit of the Santa Fe Group and the river alluvium represented by each model layer are given in table 3. The river alluvium is entirely in layer 1 (fig. 11). In row 22 (fig. 11A), the upper Santa Fe Group hydrostratigraphic unit traverses layer 1 and comprises part of layer 2, the middle Santa Fe Group hydrostratigraphic unit traverses layers 1-3 and comprises part of layer 4, and the lower Santa Fe Group hydrostratigraphic unit comprises parts of layers 2-4. Figure 11B shows a similar pattern along model column 17.



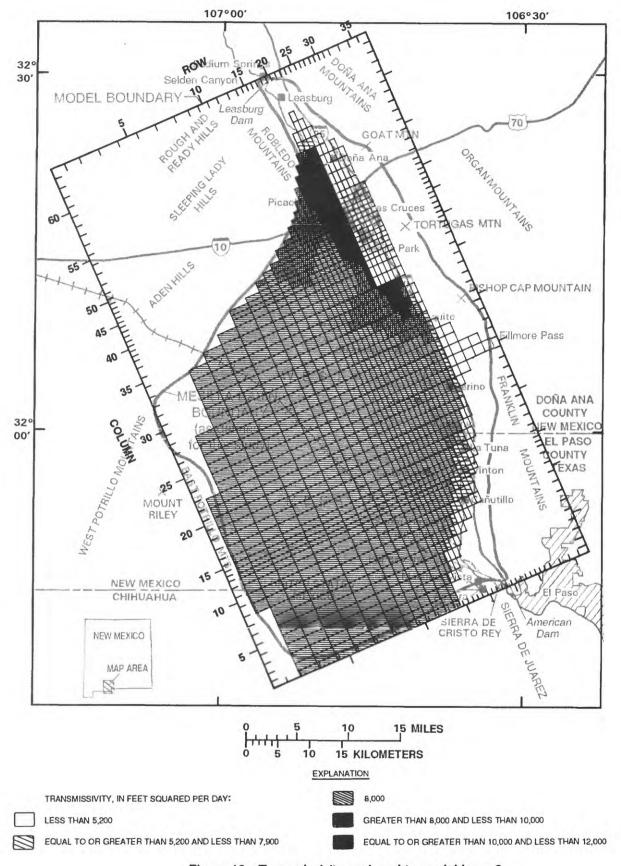


Figure 13.--Transmissivity assigned to model layer 2.

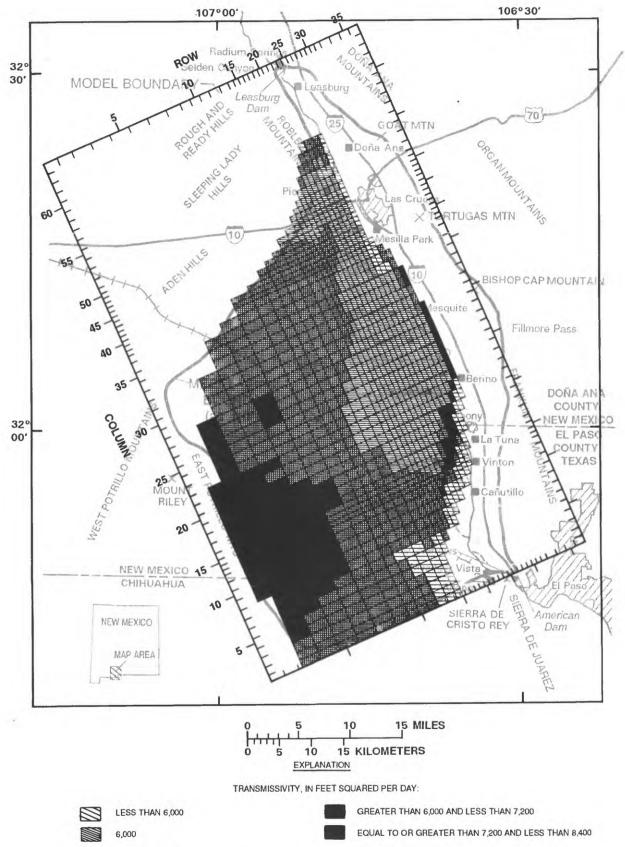


Figure 14.--Transmissivity assigned to model layer 3.

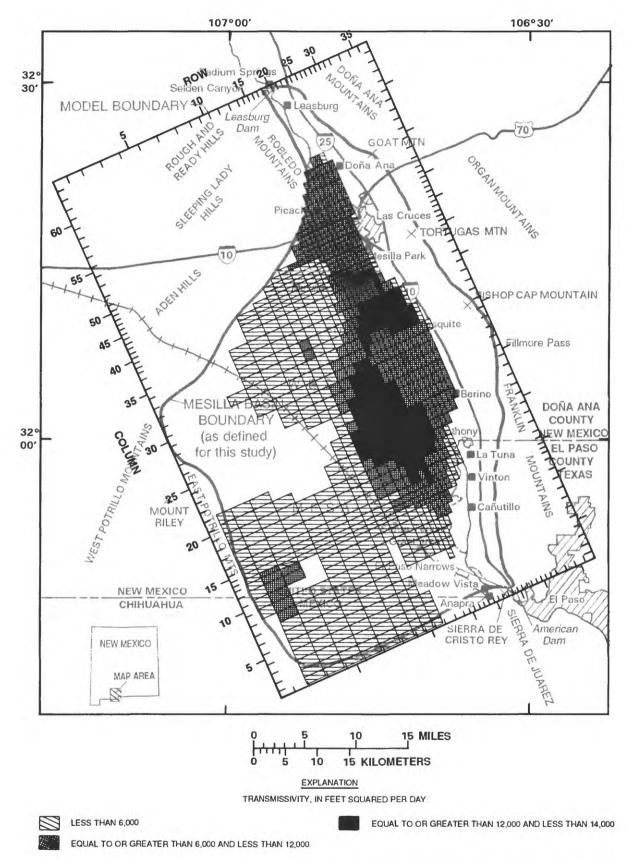


Figure 15.--Transmissivity assigned to model layer 4.

The hydraulic conductivity for the river alluvium was specified as 150 feet per day (1.71 x 10^{-4} foot per second) in most of the area of the Mesilla Valley and 350 feet per day (4.05 x 10^{-3} foot per second) in a few localities (areas marked Kl and K2 on fig. 20 of Frenzel and Kaehler, 1990). These hydraulic-conductivity values were approximately equal to those derived in the previous section on geology.

Values of hydraulic conductivity for the Santa Fe Group were assigned on the basis of location, depth, and hydrostratigraphic unit and changed during model adjustment. The Las Cruces area (model rows 25 and greater, columns 35 and greater) was given smaller values of hydraulic conductivity than the rest of the basin to simulate measured hydraulic heads in Las Cruces. A physical justification for the smaller values is the large number of faults, which are more numerous in the Las Cruces area than elsewhere. As noted previously, reduced transmissivity may be associated with faults because of cementation along fault planes or juxtaposition of clay lenses against sand and gravel lenses.

The horizontal-hydraulic-conductivity values, derived during model adjustment, fall within the upper and lower quartiles of the hydraulicconductivity values discussed in the previous section on geology. hydraulic conductivity for the upper Santa Fe Group hydrostratigraphic unit were 12 feet per day (1.39 x 10^{-4} foot per second) in the Las Cruces area and 30 feet per day (3.47×10^{-4}) foot per second) elsewhere. Values of hydraulic conductivity for the middle Santa Fe Group hydrostratigraphic unit where that unit is less than about 600 feet below the water table (layers 1 and 2) were 10 feet per day (1.16 x 10^{-4} foot per second) in the Las Cruces area and 20 feet per day (2.31 x 10^{-4} foot per second) elsewhere. Where the middle Santa Fe Group hydrostratigraphic unit is deeper than 600 feet below the water table (layers 3 and 4), hydraulic-conductivity values were 5 feet per day (5.79 x 10-5 foot per second) in the Las Cruces area and 10 feet per day $(1.16 \times 10^{-4} \text{ foot per second})$ elsewhere. The hydraulic conductivity of the lower hydrostratigraphic unit of the Santa Fe Group was set to 13 feet per day $(1.50 \times 10^{-4} \text{ foot per second})$ throughout the model. Because layers 3 and 4 are inactive in most of the Las Cruces area (figs. 14 and 15), the specification of hydraulic-conductivity values of these two layers for that area had little effect on the simulation.

The hydraulic-conductivity values shown in figure 12 are derived from the hydraulic conductivity assigned to various hydrostratigraphic units. On the left side of figure 12, where layer 1 represents only the shallower part of the middle hydrostratigraphic unit of the Santa Fe Group, the hydraulic conductivity is 20 feet per day. In other places outside the Mesilla Valley, the hydraulic conductivity is between 20 and 30 feet per day, representing partial thicknesses of the middle and upper hydrostratigraphic units of the Santa Fe Group. In the Mesilla Valley, the hydraulic conductivity is influenced by the river alluvium. The largest values of hydraulic conductivity in figure 12 occur where the river alluvium was assigned a hydraulic conductivity of 350 feet per day, except in the north where the large value represents almost entirely river alluvium, caused by the near absence of the Santa Fe Group. The smallest value in figure 12 is in the Las Cruces area outside the Mesilla Valley.

Most of layer 2 was given a transmissivity of 8,000 feet squared per day (fig. 13) where only the shallower part of the middle Santa Fe Group was represented. Where transmissivity values shown in figure 13 are greater than 8,000 feet squared per day, part or all of the 400-foot thickness of layer 2 represents the upper Santa Fe Group hydrostratigraphic unit. Figure 13 shows where the smaller transmissivity values were assigned in the Las Cruces area. On the eastern edge of layer 3 outside the Las Cruces area, the small transmissivity values are influenced by the thinning of the aquifer.

Where layer 3 represented only the deeper part of the middle hydrostratigraphic unit of the Santa Fe Group, the transmissivity is 6,000 feet squared per day (fig. 14). Smaller values represent thinning in the southeast and smaller hydraulic-conductivity values assigned to the Las Cruces area. Larger values are shown where the lower hydrostratigraphic unit is represented.

The differences in the transmissivity of layer 4 (fig. 15) mainly reflect the variable thickness of the Santa Fe Group represented. The largest transmissivity values are shown for the deepest part of the basin and the smallest values are shown for the shallowest parts.

<u>Leakance</u>

Leakances were calculated on the basis of the thicknesses of hydrostratigraphic units between nodes in the vertical direction. Equation 49 of McDonald and Harbaugh (1988, p. 5-12) describes the general method used:

$$Vcont_{i,j,k+1/2} = \frac{1}{\sum_{g=1}^{n} \frac{\Delta z_g}{K_g}}$$
 (2)

Where Vcont is the leakance;

i is the row:

j is the column;

k+1/2 indicates that the leakance is calculated between layers k and k+1.

n is the number of hydrostratigraphic units that fall between the two nodes:

 Δz_g is the thickness of the gth hydrostratigraphic unit that falls between the two nodes, in feet; and

K is the vertical hydraulic conductivity of the gth hydrostratigraphic unit that falls between the two nodes. K is equal to the horizontal hydraulic conductivity divided by the horizontal to vertical ratio that was specified for each hydrostratigraphic unit.

Consider nodes located at row 5, column 17, layers 2 and 3 (fig. 11). Between these two nodes, 200 feet of the upper part of the middle Santa Fe hydrostratigraphic unit is simulated with an assigned horizontal hydraulic conductivity of 2.31×10^{-4} foot per second (20 feet per day), and an assigned horizontal to vertical ratio of 500:1; 160 feet of the lower part of the middle Santa Fe Group hydrostratigraphic unit with an assigned hydraulic conductivity of 1.16×10^{-4} foot per second (10 feet per day), and an assigned horizontal to vertical ratio of 500:1; and 140 feet of the lower hydrostratigraphic unit with an assigned hydraulic conductivity of 1.50×10^{-4} foot per second (13 feet per day) and an assigned horizontal to vertical ratio of 100:1. Converting horizontal hydraulic conductivity to vertical hydraulic conductivity and solving the equation yield a value of Vcont equal to 8.22×10^{-10} per second. After model adjustment, the horizontal to vertical ratios were 100:1 for river alluvium, 500:1 for upper and middle hydrostratigraphic units of the Santa Fe Group, and 100:1 for the lower hydrostratigraphic unit of the Santa Fe Group.

Nodes in layer 1 were placed 150 feet from the bottom of the layer. This is consistent with the 300-foot layer-1 thickness of Frenzel and Kaehler (1990).

Specific yield and storage coefficients

The specific yield for layer 1 was 0.2, and the storage coefficients for layers 2-4 were estimated as the thickness times a specific storage of 1×10^{-6} per foot (the same as Frenzel and Kaehler, 1990, p. 54). Thus, in much of the area of layer 2, the thickness was 400 feet and the storage coefficient was 0.0004. Near the edges of the active part of layer 2, where thickness was less than 400 feet, the storage coefficient was less than 0.0004. Similarly, where the thickness of layer 3 was 600 feet, the storage coefficient was 0.0006, and less near the edges. Because the thickness of layer 4 varied to as great as 1.030 feet, the storage coefficient varied to as great as 0.00103. Although the exact specified storage coefficients are not tabulated, they can be calculated on the basis of thickness (table 3).

Initial Condition and Time Periods

The initial (pre-1915) condition was assumed to be steady-state. Because the steady-state model was kept consistent with the transient model by simultaneously adjusting aquifer characteristics in the steady-state transient simulations, the initial condition was not the same as that of Frenzel and Kaehler (1990). However, the method and rationale for generating the initial condition were exactly the same (Frenzel and Kaehler, 1990, p. 45); that is, the model properties used in the steady-state simulation were identical with those in the transient simulation except for time-dependent characteristics.

The first 16 stress periods were exactly the same as those of Frenzel and Kaehler (1990, p. 56). The additional six stress periods were defined to simulate a relative shortage of water in 1977-78 and to stop at the end of 1984 (end of stress period 21) to enable a comparison to the large number of water-level measurements that were made in the winter of 1984-85. Stress periods 1-16 simulated 1915-19, 1920-26, 1927-40, 1941-47, 1948-50, 1951-53, 1954-57, 1958-60, 1961, 1962-63, 1964, 1965-66, 1967-68, 1969-71, 1972, and 1973-75, respectively. Stress periods 17-22 simulated 1976, 1977-78, 1979-80, 1981-82, 1983-84, and 1985, respectively.

Boundary Conditions

Boundary conditions were defined, similar to those of Frenzel and Kaehler (1990, p. 56). Refinements and corrections led to differences that may not be individually significant, except possibly those in the Mesilla Valley boundary.

No-flow boundary

The exterior boundary of the basin, which is the extent of active model blocks. was different from that of the Frenzel-Kaehler model, but was defined on the same basis. The extent of the active part of layer 1 (fig. 12) was the same as that shown in figure 20 of Frenzel and Kaehler (1990) except southwest of Picacho where the boundary was straightened by activating two blocks (row 6, column 37 and row 7, column 39). Differences in the extent of the active parts of layers 2-4 reflected the more recent interpretation of the geology and the boundary between the Santa Fe Group and bedrock (Hawley, 1984; Seager and others, 1987; and J.W. Hawley and R.P. Lozinsky, written commun., 1988). The main change in the lateral extent of the simulated basin occurs in A comparison of figure 15 with figure 23 of Frenzel and Kaehler (1990) shows that layer 4 has been made less extensive in the middle of the west side and more extensive in the southwest corner of the model. Other differences around the eastern and northern edges of the active part of layer 4 generally result in a shift of the simulated boundary of about 1 mile and probably are not significant with respect to the basinwide flow system. Similarly, shifts of the boundaries of layers 2 and 3 with respect to the boundaries simulated by Frenzel and Kaehler probably are not significant. Layer 3 was deactivated in the vicinity of the Las Cruces well field to improve the simulation of hydraulic heads there, as discussed in the section on model adjustment. The basin is not as deep in the Las Cruces well-field area as it is farther west (Hawley, 1984), although the depth is not well known.

Recharge

Mountain- and slope-front recharge was located in the same places as shown in figure 25 of Frenzel and Kaehler (1990), and the values were modulated in the same way using precipitation data (Frenzel and Kaehler, 1990, p. 58). However, during model adjustment, recharge on the west side of the basin was doubled to improve the simulation of hydraulic heads in the West Mesa area. This additional recharge on the west side caused an overall increase in mountain- and slope-front recharge of 20 percent. Also, because more precipitation data were available as previously mentioned, the estimated recharge on the east side of the basin was slightly different from that of the Frenzel-Kaehler model. Figure 16 shows location and rates of steady-state recharge for layer 1 and can be compared to figure 25 of Frenzel and Kaehler (1990). The overall steady-state recharge rate was 18 cubic feet per second. Transient recharge rates are shown in figure 17A.

EXPLANATION

0.353	BLOCK WHERE RECHARGE WAS SPECIFIEDNumber is rate, in cubic feet per second				
-0.136	SPECIFIED-HEAD BLOCKNumber is model-derived underflow, in cubic feet per second. Negative sign indicates outflow				
MODEL-DERIVED EVAPOTRANSPIRATION, IN ACRE-FEET PER ACRE PER YEAR:					
	0				
	GREATER THAN 0 AND LESS THAN OR EQUAL TO 1.5				
	GREATER THAN 1.5 AND LESS THAN OR EQUAL TO 3.5				
	GREATER THAN 3.5 AND LESS THAN OR EQUAL TO 5.5				

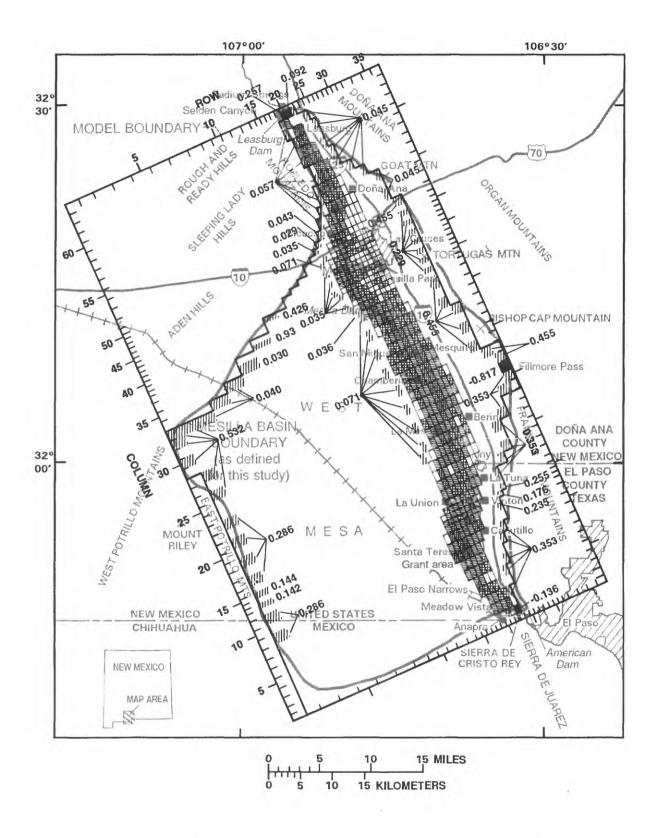


Figure 16.--Recharge, underflow, and evapotranspiration boundaries for layer 1, steady-state simulation.

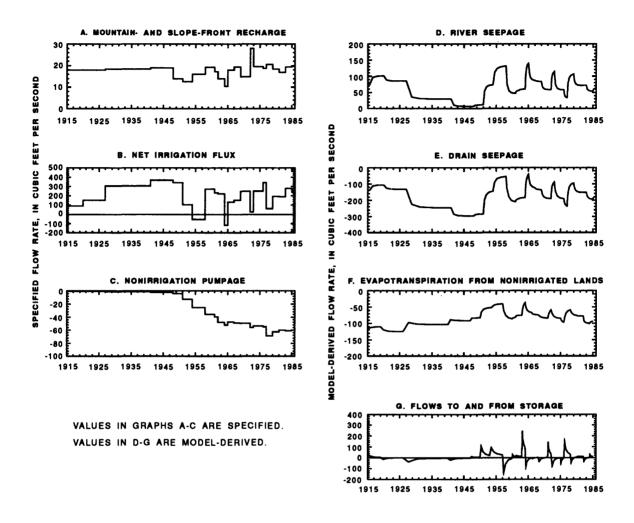


Figure 17.--Major basinwide flow rates.

Underflow

Underflow through the flood-plain alluvium in Selden Canyon and El Paso Narrows was simulated with specified-head boundaries exactly as was done by Frenzel and Kaehler (1990, p. 60), and an additional specified-head boundary was used to simulate flow out of the basin through Fillmore Pass. A specified head of 3,814 feet in row 36, column 26, layers $\bar{1}$ and 2 simulated the measured heads near Fillmore Pass. The model-derived steady-state inflow rate of about 0.35 cubic foot per second at Selden Canyon and outflow of about 0.14 cubic foot per second at El Paso Narrows compare with 0.15 and 0.09 cubic foot per second in the Frenzel-Kaehler model. The steady-state outflow of 0.99 cubic foot per second at Fillmore Pass (0.82 cubic foot per second from layer 1 and 0.17 cubic foot per second from layer 2) was not simulated in the Frenzel-These values are not significant with respect to the overall mass balance of the model, which was 300 cubic feet per second under steadystate conditions. Steady-state underflows at specified-head blocks are shown in figure 16 for layer 1. Under transient conditions, the model-derived inflow and outflow rates at specified-head boundaries were about 2 cubic feet per second.

Mesilla Valley boundary

As described in Frenzel and Kaehler (1990, p. 60):

A complex boundary at or near the land surface in the Mesilla Valley was approximated by three overlapping boundary conditions. They were: (1) A head-dependent flux to approximate flow to and from the river and drains; (2) another head-dependent flux to approximate evapotranspiration from nonirrigated lands; and (3) a specified net irrigation flux to approximate the summation of effective rainfall and net diversions (positive) and evapotranspiration from irrigated lands (negative) [fig. 17, graph B]. By implication, pumpage for irrigation from layer 1 is included in this summation because evapotranspiration is accounted for directly.

Flow to and from the river and drains. -- The locations of flow to and from the river and drains are shown in figures 27 and 28 of Frenzel and Kaehler (1990). The conductances assigned to the river and drains were the same as those of Frenzel and Kaehler (1990, p. 60-65, table 3). The altitudes assigned to the river and drains were set 1 foot lower than those of Frenzel and Kaehler, consistent with their suggestion (p. 101).

Surface-water inflow was assigned to the farthest upstream river node for Selden Canyon exactly the same as in Frenzel and Kaehler (1990, p. 65), and inflow from 1975 to 1985 was calculated in the same way. Inflows to river reaches including municipal return flows are shown in table 4. Other river specifications are shown in table 3 of Frenzel and Kaehler.

Evapotranspiration from nonirrigated lands. -- Evapotranspiration from nonirrigated lands in the Mesilla Valley was simulated in the same way as described by Frenzel and Kaehler (1990, p. 65) except that the evapotranspiration surface (McDonald and Harbaugh, 1988, p. 10-1) was corrected to more closely correspond to the land surface, and the extinction depth (the depth below the evapotranspiration surface at which model-derived evapotranspiration is zero) was reduced from 15 feet to 12 feet. The effect of corrections to the evapotranspiration surface was to make evapotranspiration more uniform (compare fig. 16 of this report with fig. 25 of Frenzel and Kaehler, 1990). The extinction depth was reduced in order to reduce the model-derived depletion. The maximum evapotranspiration from nonirrigated lands was specified to be 5.5 acre-feet per nonirrigated acre, exactly the same as that of Frenzel and Kaehler (1990, p. 65-66). Evapotranspiration from nonirrigated land was simulated for the entire valley, and to avoid overestimating it, the maximum rate of 5.5 feet was multiplied by a factor equal to the nonirrigated acreage divided by the total valley acreage.

The overall rate of evapotranspiration in the transient simulation is shown in figure 17F. The model-derived evapotranspiration for 1975 was 85 cubic feet per second. If this evapotranspiration is assumed to occur on 32,000 acres (Frenzel and Kaehler, 1990, p. 66), this is the equivalent of 1.9 acre-feet per acre. This compares with the Frenzel-Kaehler value of about 2.5 acre-feet per acre. Model-derived evapotranspiration for 1975 shown in figure 18 is similar to figure 29 of Frenzel and Kaehler (1990) except that evapotranspiration shown in figure 18 is more evenly distributed as a result of corrections to the simulated altitude of the evapotranspiration surface.

<u>Evaporation from streams</u>.-- Evaporation from streams was treated outside of the model the same way as by Frenzel and Kaehler (1990, p. 66). That is, 16,000 acre-feet was subtracted from annual measured flows passing Leasburg Dam before routing through the stream system, and this amount (equivalent to 22 cubic feet per second) was added in the calculation of model-derived depletion.

Net irrigation flux.-- Net irrigation flux was treated the same way as by Frenzel and Kaehler (1990, p. 66). The flux of the first stress period was the same as that of Frenzel and Kaehler and later stress periods were derived from figure 8, curve E in the same way. The net irrigation fluxes for each stress period are shown in figure 17B. Although values in figure 17B are not the same as those in Frenzel and Kaehler (1990, fig. 26B), the curves have a similar shape.

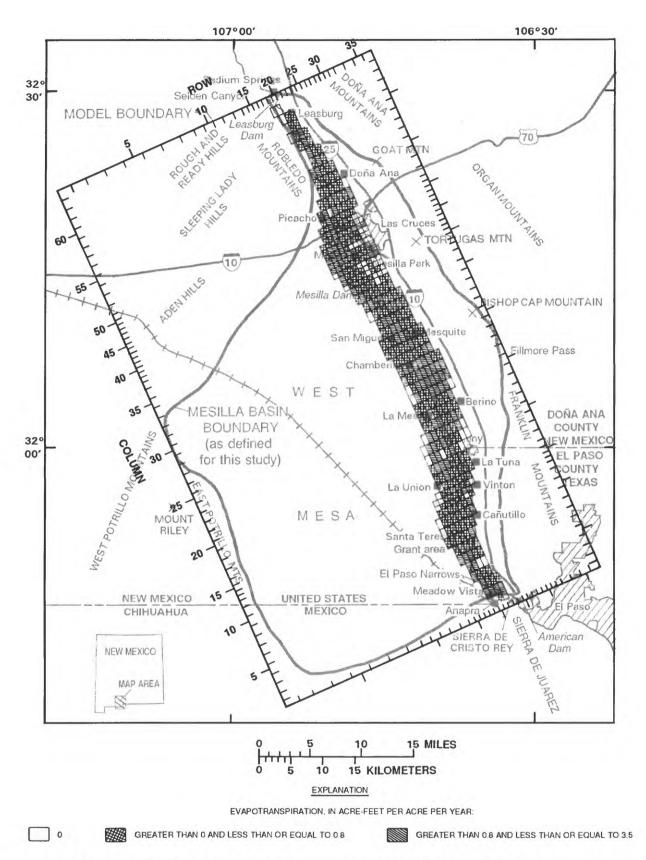


Figure 18.--Model-derived evapotranspiration from nonirrigated lands for 1975 and area of net irrigation flux.

Nonirrigation withdrawals.-- Nonirrigation ground-water withdrawals were treated the same way as by Frenzel and Kaehler (1990, p. 69). Return flows generally were treated in the same way also except that return flows at Anthony were assumed to have been via the East Drain after 1975 and return flows at Santa Teresa were assumed to have taken the form of ground-water recharge after 1975. Because part of the withdrawals at Santa Teresa were for irrigation of sod and the sewage effluent was disposed of by land application, return flows were assumed to be 25 percent of withdrawals, a value less than the assumed 50-percent returns from other municipal usage. Nonirrigation ground-water withdrawals for the first 16 stress periods were the same as those of Frenzel and Kaehler (1990, table 4), except for a few corrections. Simulated nonirrigation ground-water withdrawals are shown in table 5. The totals for each stress period are shown in figure 17C, which is comparable to figure 9, line E.

Model Adjustments

The general goal of model adjustment was to make model-derived values, such as hydraulic head, match measured values reasonably well while keeping simulated properties, such as hydraulic conductivity, within generally recognized ranges for the type of deposits simulated. The model, with the refinements described above, needed little adjustment. Boundary conditions were similar to those in the Frenzel-Kaehler model (1990), and with few exceptions were assumed not to need more adjustment.

The emphasis in adjustment of the revision of the model was on the simulated values of hydraulic conductivity of the Santa Fe Group. The main difference between this model and the Frenzel-Kaehler model was the way in which transmissivity values and leakances were derived. However, the correct simulation of the effects of ground-water withdrawals on streamflow in the Rio Grande depends on the specification of correct aquifer characteristics for the Santa Fe Group. As Frenzel and Kaehler (1990, p. 98) showed, this becomes more critical as the point of withdrawal is moved more distant from the Mesilla Valley boundary. Because transmissivity and leakance largely determine the time required for the effects of ground-water withdrawals to reach the river, the adjustment of Santa Fe Group hydraulic-conductivity values was given special emphasis.

System Properties Used for Comparison

As in the Frenzel-Kaehler model, measured values of hydraulic head, river depletion, and drain discharge were compared with model-derived values. Because depletion and drain discharge are closely related to the Mesilla Valley boundary, which was not emphasized in the adjustment of this model, hydraulic-head comparisons received more emphasis.

Surface-water depletions

Annual surface-water depletions were calculated as the measured discharge of the Rio Grande at Leasburg Dam minus the measured discharge of the Rio Grande at El Paso Narrows. This depletion is referred to as "measured" depletion (fig. 19). Similarly, depletion calculated as stress period-averaged discharge at Leasburg Dam minus model-derived discharge at El Paso Narrows is referred to as "model-derived" depletion. Measured and model-derived depletions were calculated the same way as by Frenzel and Kaehler (1990, p. 78). Figure 19 compares favorably with figure 39 of Frenzel and Kaehler (1990). Each point of the curve representing model-derived values in figure 19 is for the end of a time step.

Differences between measured and model-derived depletion were large before the 1950's. This could be partly because the model simulated the progress of drainage ditching and the concurrent changes on ground-water storage with only a few stress periods. Also, irrigated agriculture may have changed over the years in ways that the model did not simulate. For example, the consumption of water by any given crop might have changed over the long term because of changing varieties and irrigation practices.

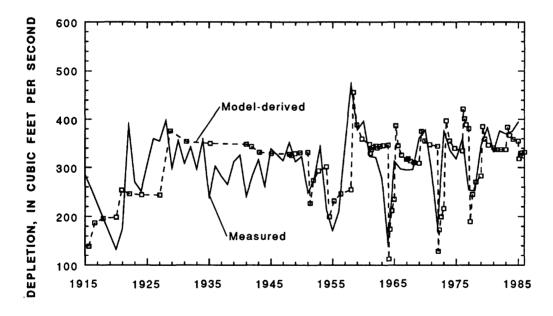


Figure 19.--Relation of annual surface-water depletions calculated from measured discharges to depletions calculated from model-derived discharges.

The summation of measured depletions versus the summation of annual model-derived depletions is shown in figure 20. Each point plotted in figure 20 represents the summation 1 year later than the previous point. straight, diagonal line in figure 20 is the line of equality between measured and model-derived values. During periods when the slope of the line defined by the points is approximately parallel to the line of equality (such as 1945-55), the model-derived depletions are approximately equal to measured values. When the points define a line that is steeper than the line of equality (such as 1930-45), model-derived depletions are larger than measured During much of the 16-year period 1965-80, model-derived depletions. depletions are slightly larger than measured values; after that the trend changes slightly so that, taken as a whole during the last 20 years of simulation, model-derived depletions are approximately equal to measured To construct figure 20, annual model-derived depletions were estimated by linear interpolation between time steps.

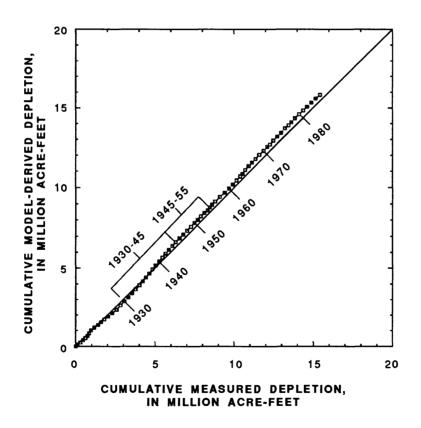


Figure 20.--Relation of cumulative surface-water depletions calculated from measured discharges to cumulative depletions calculated from model-derived discharges.

Drain discharges

Measured and model-derived drain discharges were compared to assure that model adjustments did not adversely affect model-derived drain discharge. Model-derived drain discharge is largely dependent on specified drain-bottom hydraulic conductivity, which was not adjusted; therefore, model-derived drain discharges were not expected to be much different from those of Frenzel and Kaehler (1990). The comparison between measured and model-derived discharges before 1975 (fig. 21) is similar to that of Frenzel and Kaehler (1990, fig. 40). The comparison between measured and model-derived discharges after 1975 is similar to before 1975 except at East Drain. There, measured values trend upward whereas model-derived values continue a 30-year downward trend. The reason for this divergence is not known. The accuracy of drain-discharge measurements is poor (Frenzel and Kaehler, 1990, p. 39 and 83).

Hydraulic heads

The same measured hydraulic heads were used for 1947 and 1975 comparisons as used by Frenzel and Kaehler (1990). In addition, hydraulic heads for 1985 were used. The comparison of individual measured and model-derived hydraulic heads is shown in table 6, which is similar to table 5 of Frenzel and Kaehler (1990). The average of the differences shown in table 6 was calculated in four ways: the arithmetic mean, the mean of absolute values of the differences ("mean absolute" difference), the median, and the root-mean-square as explained in Frenzel and Kaehler (1990, p. 72). Differences for 1947, 1975, and 1985, calculated from 181 locations, were averaged together (bottom of table 6). The arithmetic mean difference of 1.4 feet and median of 0.90 foot indicate that the model-derived hydraulic heads are, on average, slightly higher than measured values. The mean absolute difference of 7.34 feet compares favorably with values in table 8 of Frenzel and Kaehler (1990). mean absolute difference was used as the primary statistic in deciding if one simulation was better than another. The simulated hydraulic-conductivity values of the Santa Fe Group hydrostratigraphic units were adjusted to minimize the mean absolute difference. The root-mean-square difference of 12.43 feet emphasizes the large differences for 1947 on the west side of the basin, but is similar to root-mean-square differences of Frenzel and Kaehler (1990, tables 5 and 8).

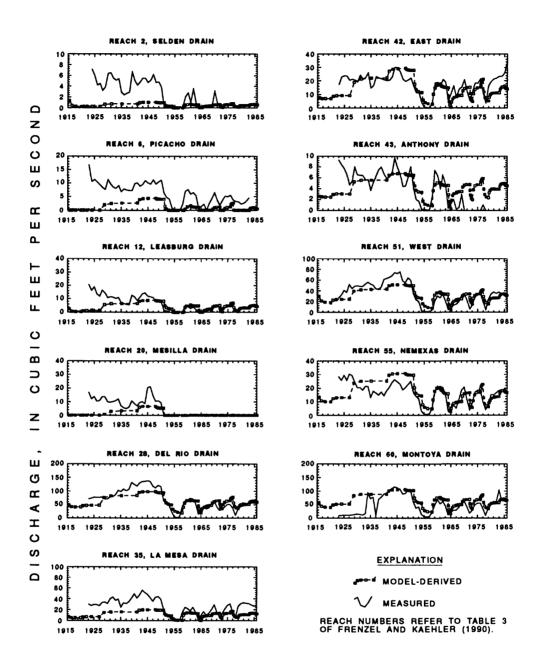


Figure 21.--Measured and model-derived drain discharges.

To increase the sensitivity to changes in simulated Santa Fe Group aquifer characteristics, only hydraulic heads outside of the Mesilla Valley or in the deeper parts of the aquifer (represented by layers 2 and 3) were compared to model-derived values during the adjustment of Santa Fe Group aguifer characteristics. Specified Santa Fe Group aguifer characteristics can be adjusted only on the basis of a statistic that is sensitive to the adjustments. Frenzel and Kaehler (1990, p. 95-101) found that the differences between measured and model-derived values generally were not very sensitive to Santa Fe Group aquifer characteristics. This is especially true of hydraulic heads in the shallow part of the aquifer in the Mesilla Valley, which are largely determined by the Mesilla Valley boundary. About one-half of the sites shown in table 6 are in this category. The remaining 90 sites, marked by an asterisk in table 6, were used in calculating the averages shown at the bottom of table 6. Adjustments of Santa Fe Group aquifer characteristics were made mainly on the basis of minimizing the mean absolute difference between measured and model-derived hydraulic heads at the sites outside of the Mesilla Valley.

Measured hydraulic heads are compared to model-derived potentiometric surfaces in figures 22-24. The comparison of measured hydraulic heads and the model-derived steady-state potentiometric surface for layer 1 (fig. 22) is slightly better than that of Frenzel and Kaehler (1990, fig. 31) in the West Mesa area. Similarly, the 1975-76 comparison for the West Mesa area (fig. 23) is slightly better than that of Frenzel and Kaehler (1990, fig. 33). The 1975-76 comparison for the Mesilla Valley area generally is good and similar to that of Frenzel and Kaehler (1990), and the Mesilla Valley boundary was the same in both models. The model-derived potentiometric surface may be about 5-20 feet too high in the Cañutillo area because several measured values there are all lower than the model-derived surface. The 1984-85 comparison (fig. 24) also shows a reasonably good match between measured hydraulic heads and the model-derived potentiometric surface for the West Mesa area, but no measured values were found for the well-field areas. Hydrographs of measured and model-derived hydraulic heads for well-field locations are compared in figure 25. Generally the trends are the same as those of Frenzel and Kaehler (1990, fig. 36) with some improvement in the Las Cruces area.

System Properties Adjusted

Simulated extinction depth of evapotranspiration (part of the Mesilla Valley boundary), recharge rates on the west side of the basin, basin depth near Las Cruces, specific yield, and specific storage were adjusted during transient simulations. Values of hydraulic conductivity of the Santa Fe Group were adjusted during and after the other adjustments.

Extinction depth

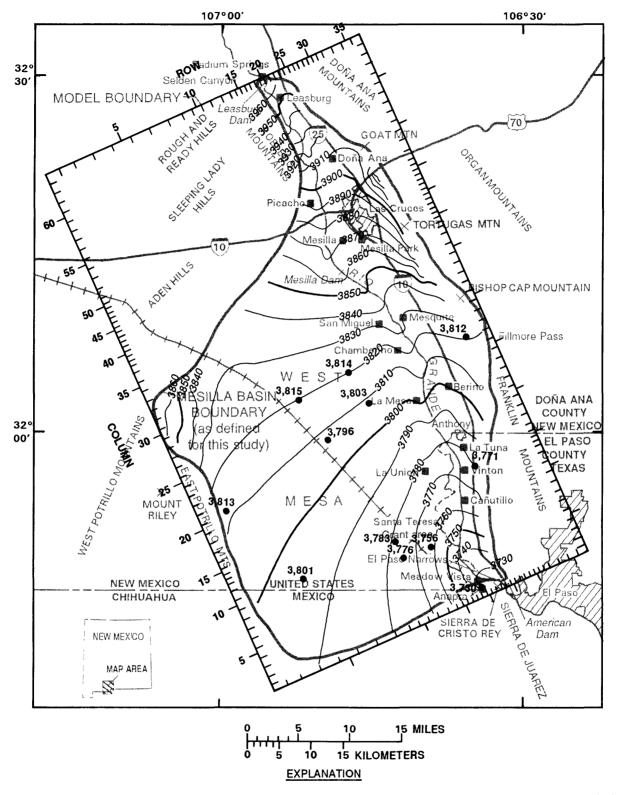
The increase in simulated evapotranspiration from irrigated land (over that of the Frenzel-Kaehler model) caused model-derived depletion of the Rio Grande to be too large. To correct this, the extinction depth (McDonald and Harbaugh, 1988, p. 10-1) of evapotranspiration from nonirrigated land was reduced from 15 to 12 feet. At this point, model-derived drain discharges (fig. 21), depletion of the Rio Grande (figs. 19 and 20), and hydraulic heads in layer 1 (table 6) for the Mesilla Valley matched measured values reasonably well. Additional model adjustments were made on the basis of minimizing the mean absolute difference between measured and model-derived hydraulic heads at depth and outside of the Mesilla Valley (marked with an asterisk in table 6).

Recharge

Mountain-front recharge for the west side of the basin was adjusted to improve the simulation of hydraulic heads for that area. As can be seen in table 5 of Frenzel and Kaehler (1990), the differences between measured and model-derived hydraulic heads for 1947 and 1975 are generally negative for rows 1-19 and positive for rows 20-36. Because trial-and-error adjustments of aquifer characteristics did not seem to correct this "skewness," simulated recharge values on the west side of the basin were doubled to achieve the desired effect. This increased recharge resulted in an overall increase of about 20 percent in simulated mountain-front recharge. From this point on, model adjustment was focused on Santa Fe Group aquifer characteristics, although storage was adjusted for the entire model.

Specific yield and specific storage

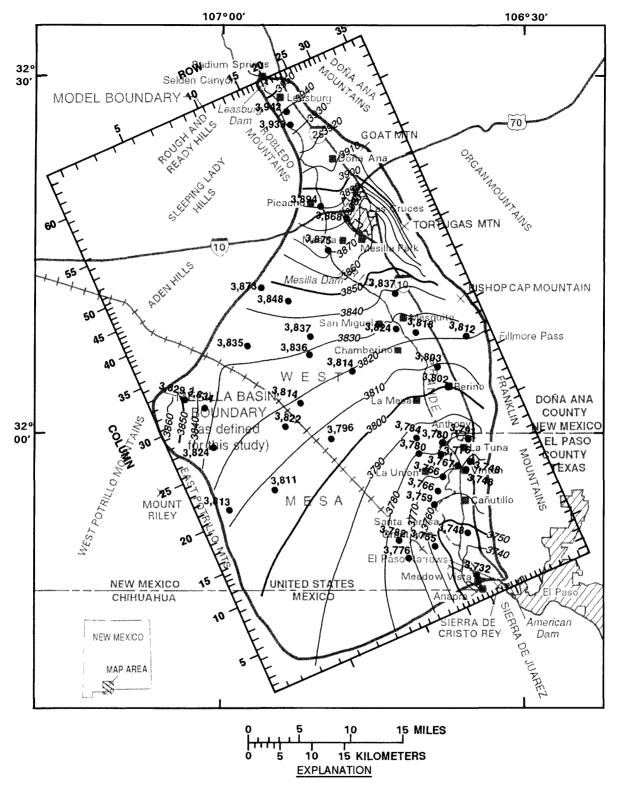
Specific yield for layer 1 was originally specified as 0.2 and specific storage for layers 2-4 as 1 x 10^{-6} per foot. Although the specific yield and specific storage were adjusted, the final values were the same as the values originally specified.



-3730- MODEL-DERIVED POTENTIOMETRIC CONTOUR--Shows simulated altitude at which water would have stood in tightly cased wells. Contour interval 10 feet. Datum is sea level

• 3,730 WELL--Number is measured hydraulic head, in feet above sea level

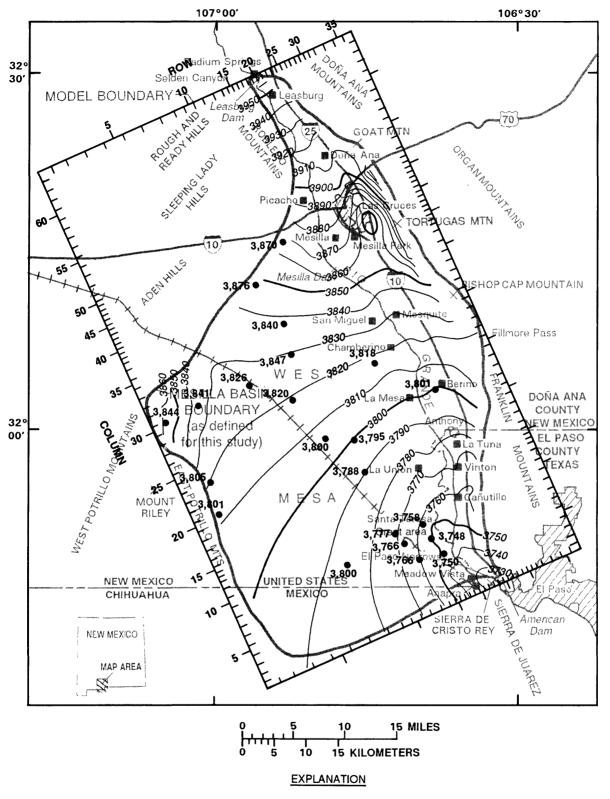
Figure 22.--Measured hydraulic heads and model-derived steady-state potentiometric surface, layer 1.



__3750— MODEL-DERIVED POTENTIOMETRIC CONTOUR.-Shows simulated altitude at which water would have stood in tightly cased wells. Contour interval 10 feet. Datum is sea level

● 3,732 WELL--Number is measured hydraulic head in 1975-76, in feet above sea level

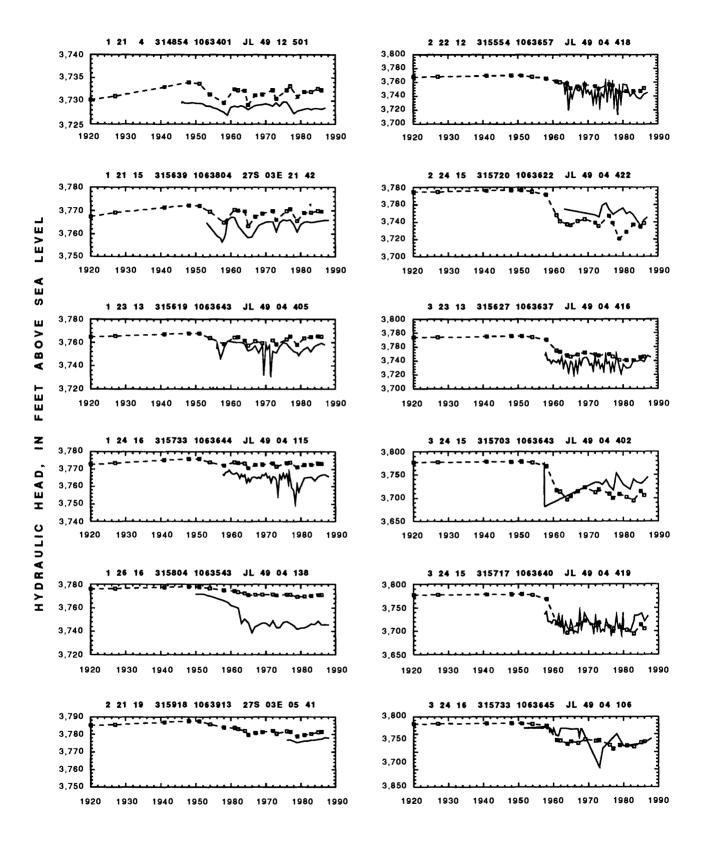
Figure 23.--Measured hydraulic heads and model-derived potentiometric surface for 1975-76, layer 1.



-3800- MODEL-DERIVED POTENTIOMETRIC CONTOUR--Shows simulated altitude at which water would have stood in tightly cased wells. Contour interval 10 feet. Datum is sea level

●3,801 WELL--Number is measured hydraulic head in 1984-85,in feet above sea level

Figure 24.--Measured hydraulic heads and model-derived potentiometric surface for 1984-85, layer 1.



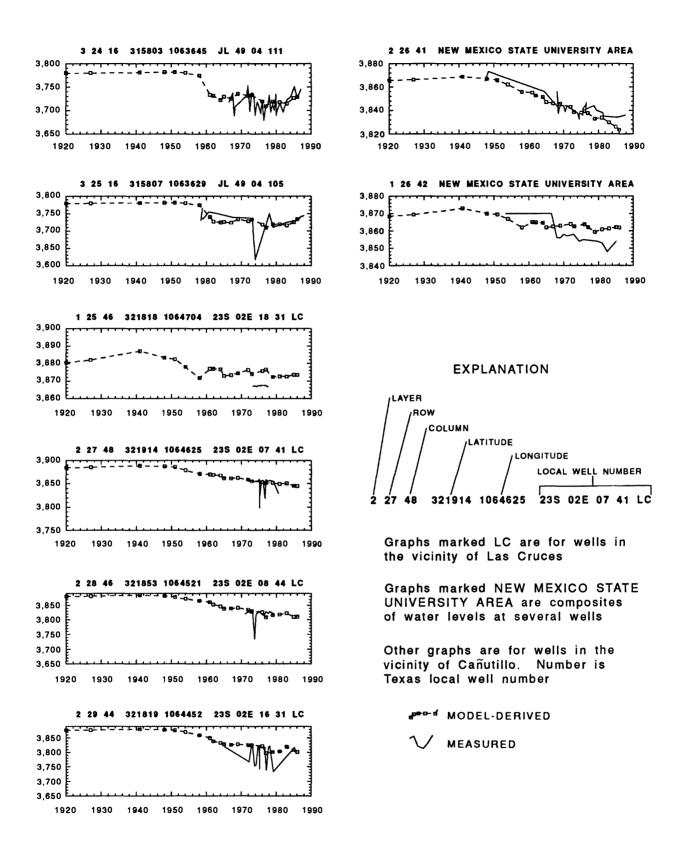


Figure 25.--Measured and model-derived hydraulic heads.

Model boundary

The no-flow boundary of model layer 3 was moved westward about 1 mile in the Las Cruces area (rows 26-27, columns 40-56) to increase simulated drawdowns and thereby improve the simulation of hydraulic heads in the Las Cruces well field. The basin is not as deep in the Las Cruces well field as elsewhere, and is simulated by two layers. The locality where the basin becomes deeper is not precisely known. The model-derived heads of Frenzel and Kaehler (1990, fig. 36) were between 10 and 40 feet higher than measured values, differences that are difficult to attribute entirely to errors or anomalies in the measured values. To improve the fit between measured and model-derived hydraulic heads, measures were taken to increase simulated This was partially accomplished by moving the no-flow boundary of drawdowns. layer 3, thus slightly reducing the simulated thickness of Santa Fe Group This adjustment resulted in a minor (about 2 feet) increase in drawdown. Adjustments of simulated hydraulic-conductivity values also served to increase drawdowns.

Hydraulic conductivity

Hydraulic-conductivity values specified for each hydrostratigraphic unit of the Santa Fe Group were changed individually in a trial-and-error manner to minimize the mean absolute difference between model-derived hydraulic heads and 90 measured values (bottom of table 6). Santa Fe Group hydraulicconductivity values for the Las Cruces area (rows 25 and greater, columns 35 and greater, figs. 12-14) were adjusted separately from hydraulic-conductivity values in the rest of the model. Final values of hydraulic conductivity for the area outside of the Las Cruces area were approximately equal to the median values of estimated hydraulic conductivity, and within the Las Cruces area values of hydraulic conductivity were about one-half as large as outside. Although changes in hydraulic conductivity were usually about 10 percent, The final value for the lower round numbers were preferred. hydrostratigraphic unit of the Santa Fe Group (13 feet per day) was the same as that used by Frenzel and Kaehler (1990, p. 86) for layer 3 in the Canutillo well field.

Vertical hydraulic conductivity was specified for each hydrostratigraphic unit as a ratio of horizontal hydraulic conductivity. The ratio was adjusted by units of 100 starting with values in the range of 100-500, considering that Frenzel and Kaehler (1990, p. 86) derived a ratio of 200:1. On the basis of the greater lenticularity of the upper and middle hydrostratigraphic units of the Santa Fe Group relative to the lower unit, the lower unit was assumed to be less anisotropic. Upon the last adjustment, the ratios were set to 100:1 for the river alluvium and lower hydrostratigraphic unit of the Santa Fe Group and 500:1 for the upper and middle hydrostratigraphic units of the Santa Fe Group.

Effects of Existing Withdrawals on Flow in the Rio Grande

Ground-water withdrawals were derived from three simulated sources: increased depletion of the Rio Grande, salvaged evapotranspiration, and temporary depletion of storage in the aquifer. Part of the depletion of the Rio Grande is due to evapotranspiration from nonirrigated lands. Ground-water withdrawals cause a lowering of the water table and may cause a reduction in evapotranspiration from nonirrigated land. In the following discussion, this reduction is termed "salvaged evapotranspiration" because water derived from this source does not increase the depletion of the Rio Grande. The lowering of the water table also causes decreased flow toward and increased flow from streams, which result in increased depletion of the Rio Grande. The simulated depletion of aquifer storage is eventually replenished by salvaged evapotranspiration or increased depletion of the Rio Grande, or both. However, the time of replenishment may vary greatly depending on the location and continuation of ground-water withdrawal and on aquifer characteristics. Although it is temporary, depletion of aquifer storage is simply referred to as "depletion of storage" in the following discussion. The term "standard" model in the following discussions refers to the final adjusted version described in the section "Description of the model."

The effects of existing nonirrigation withdrawals on flow in the Rio Grande were estimated by comparing the mass balance of the standard model with the mass balance of a simulation that did not include nonirrigation withdrawals. The following is a mass balance for the end of 1985 for the standard model, which includes nonirrigation withdrawals (all units are in cubic feet per second):

	Ground- water with- drawals and recharge	Evapotrans- piration	River and drains	Constant- head boundaries	Aquifer storage	Total
Inflows Outflows	259.67 -61.02	0 -94.18	76.51 -191.73	2.04 -2.01	12.38 -2.15	350.60 -351.09
Net	198.65	-94.18	-115.22	.03	10.23	49

"Ground-water withdrawals and recharge" are specified values including mountain- and slope-front recharge, net irrigation flux, and nonirrigation withdrawals. "Evapotranspiration" is from nonirrigated lands, a model-derived rate. "River and drains" are seepage to the river and drains combined. "Constant-head boundaries" are model-derived underflows in Selden Canyon, El Paso Narrows, and Fillmore Pass. "Aquifer storage" is a model-derived rate of depletion (inflow) or replacement (outflow) of water in aquifer storage. "Total" is the summation of the previous five items and "net" is the summation of inflows and outflows. By definition, inflow must equal outflow in a mass balance or, because they are given opposite sign, they must sum to zero. Due to mathematical truncation and rounding errors, there is always a discrepancy between simulated inflows and outflows. The discrepancy (total net) is considered acceptable.

The following is a mass balance for the end of 1985 for a model that is the same as the standard except that it does not include nonirrigation withdrawals:

	Ground- water with- drawals and recharge	Evapotrans- piration	River and drains	Constant- head boundaries	Aquifer storage	Total
Inflows Outflows	258.52 0	0 -105.20	63.35 -216.88	2.04 -2.07	3.38 -2.92	327.29 -327.07
Net	258.52	-105.20	-153.53	03	.46	. 22
Differ- ence from standard model	-59.87	11.02	38.31	0.06	9.77	71

The difference is calculated between the net of this model and the net of the standard. The "difference" row in the table shows that the change in "groundwater withdrawals and recharge" (nonirrigation withdrawals) is made up by changes in the other four categories with an accumulated discrepancy of 1.2 percent of the nonirrigation withdrawals (-0.71 cubic foot per second). Differences were calculated for all time steps in a similar manner and are shown in figure 26A. Figure 26A was constructed so that the differences are represented by the distances between the curves, and the sum of all the positive differences are represented by the uppermost curve, which is the rate of nonirrigation withdrawals. The areas between the curves represent the quantity of water taken from each source. Figure 26A shows that the standard model simulated about two-thirds of nonirrigation withdrawals taken from increased depletion of the Rio Grande, about one-sixth from depletion of aquifer storage, and about one-sixth from salvaged evapotranspiration. In a similar figure, Frenzel and Kaehler (1990, fig. 42A) showed about four-fifths of nonirrigation withdrawals taken from increased depletion of the river. In either simulation most nonirrigation withdrawals are simulated to have been taken from an increase in depletion of the Rio Grande. However, the exact percentage is uncertain and the percentage would be different if the locations of withdrawals were different.

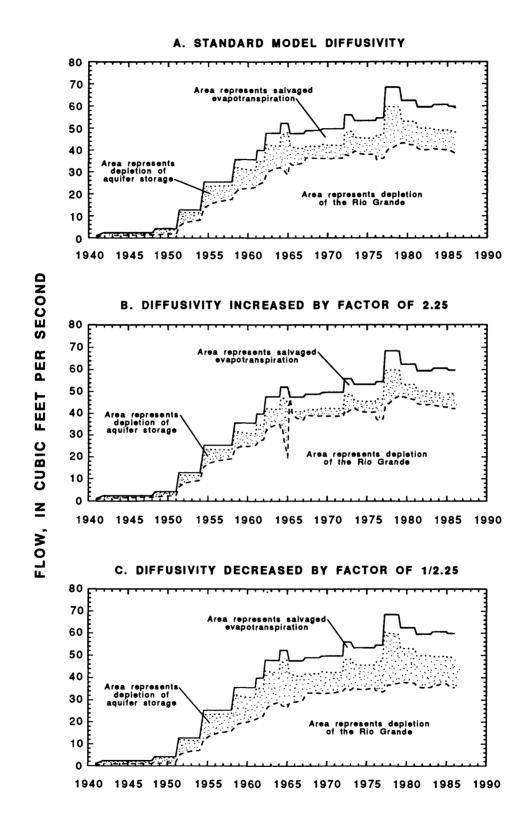


Figure 26.--Effect of changes in diffusivity on the source of water for nonirrigation withdrawals.

Items in the mass balance of the model can be divided differently to extract additional information. Similar to the mass balance of Frenzel and Kaehler (1990, p. 88), the following table shows the mass balance for the standard model where ground-water withdrawals and recharge, river and drains, and constant-head boundaries have been divided into additional groups. The main components of this mass balance are shown in figure 17. The balance is shown for the initial condition (pre-1915, which is assumed to be steady state), for 1975, and for the end of the simulated time (1985) (all units are in cubic feet per second):

	Initial condition (steady state)	1975	End of simulated time (1985)
Inflow from:			
Net river seepage Net irrigation flux Mountain- and slope-front recharge Underflow in flood-plain alluvium	280.53 .00 17.92 35	57.83 251.13 19.53 	62.13 238.81 19.71
Total inflows	298.80	330.53	322.69
Outflow to:			
Net drain seepage Net nonirrigation ground- water withdrawal	.00 .00	189.89 53.48	177.35 59.86
Evapotranspiration from nonirrigated lands	297.49	84.56	94.18
Underflow in flood-plain	.14	.96	.91
Underflow through Fillmore Pass	.99	1.08	1.10
Net flow to storage	00	39	<u>-10.24</u>
Total outflows	298.62	330.36	323.16
Difference	.18	.17	47
Percentage difference	.06	.05	.15

The main differences between this mass balance and that of Frenzel and Kaehler (1990, p. 87) are in the net river seepage and in evapotranspiration from nonirrigated lands. Both items were reduced by about 12 percent for steady-state conditions and by about 25 percent for 1975. Although underflows were increased by as much as 100 percent, they constitute less than 1 percent of the total mass balance. Underflow in Fillmore Pass was not simulated by Frenzel and Kaehler (1990).

Because river and drain seepages are net values, including outflows in the river seepage and inflows in the drain seepages, the inflows and outflows of this table do not compare with those of the above mass-balance table for the standard (the standard includes nonirrigation withdrawals). However, the discrepancy (shown above as "difference") for 1985 is nearly equal to the discrepancy for the standard, as it should be.

With some assumptions, an estimate can be made of the simulated sources of water consumed by evapotranspiration from irrigated lands. Frenzel and Kaehler (1990, p. 96) found that the ultimate sources of nearly 100 percent of water withdrawals that are accounted for by this model were the boundaries-specifically, salvaged evapotranspiration and depletion of the river. Although storage is an important temporary source, the time lag between the depletion and the replenishment of storage is short (perhaps annual) because simulated irrigation activities are all within or very close to the Mesilla Valley boundary along with the river and nonirrigated lands. If it is assumed that evapotranspiration from nonirrigated lands would be approximately equal to the simulated value for steady-state conditions except for human activities, then the difference between steady-state and 1985 evapotranspiration from nonirrigated lands, about 200 (297.49 minus 94.18) cubic feet per second, is equal to the simulated evapotranspiration salvaged by human activities. The human activities that might salvage evapotranspiration mainly have been nonirrigation ground-water withdrawals, and activities associated with irrigation such as drainage, ground-water withdrawals, and replacement of areas of natural evapotranspiration with irrigated cropland. The simulated quantity of evapotranspiration salvaged by nonirrigation withdrawals for 1985 has been shown to be about 10 (11.02) cubic feet per second. The difference, 190 cubic feet per second, constitutes an estimate of salvage by irrigation. The net flow from aquifer storage was approximately 10 (-10.24 flow to storage) cubic feet per second, most of which (9.77 cubic feet per second) was taken by nonirrigation withdrawals, leaving almost no net change in storage caused by irrigation. In the simulation, therefore, all the estimated evapotranspiration from irrigated lands, about 250 cubic feet per second (fig. 8, curve D), came from salvaged evapotranspiration (190 cubic feet per second) and from increased depletion of the Rio Grande (60 cubic feet per second, which is 250 minus 190).

Steady-state (pre-1915) evapotranspiration from nonirrigated lands may be much in error, resulting in large errors in the estimate of evapotranspiration salvaged by irrigation, and hence the estimate of increased depletion of the Rio Grande. For example, whereas a 3-foot error in simulated ground-water altitudes would be acceptable in view of the accuracy of water-level and land-surface altitudes, the same 3-foot error could cause more than a 1-acre-foot-per-acre error in simulated evapotranspiration.

Model Evaluation and Sensitivity Tests

The model's usefulness for projections may be evaluated on the basis of how well it accounts for geologic properties and how well it simulates the historical trends in the hydrologic data. This model accounts for the most recent geologic interpretations, and simulates a decade more of hydrologic data than the Frenzel-Kaehler (1990) model. The model described in this report simulates the hydrologic data reasonably well, as can be seen from comparisons of depletions (figs. 19 and 20), drain flows (fig. 21), and hydraulic heads (figs. 22-25). This does not prove that the model is correct, only that it is reasonably consistent with the data. As is the case for many ground-water models, this model is not unique. Other models also may be reasonably consistent with these data and as more data become available, the models found to be the most consistent with all the data might be retained instead of the others. Hence, as more data become available, this model may be altered or discarded. In the meantime, this model is a reasonable tool for projection.

The adjustment of simulated aquifer characteristics for the Las Cruces area separately from elsewhere improved the overall fit of model-derived to measured hydraulic heads but has some drawbacks. The basinwide adjustment of aquifer characteristics to simulate drawdowns in well fields essentially uses ground-water withdrawals (stress) and drawdowns (response) as a large-scale, long-term aquifer test. The main well fields in the Mesilla Basin are at Cañutillo and Las Cruces. The Cañutillo well field probably could be considered to be part of a larger lower Mesilla Valley field. The adjustment of the Las Cruces area separately implies that the stress-response relation for the Las Cruces area is not considered useful for adjusting the aquifer characteristics of the basin as a whole, leaving the adjustment of Santa Fe Group aquifer characteristics dependent only on the stress-response relation in the lower Mesilla Valley. As stress-response relations become better known for the West Mesa, the model can be adjusted more accurately for that area.

Sensitivity tests were used to evaluate the model. The sensitivity of the model was tested by setting selected aquifer properties at either 1.5 or 0.67 times the values used in the standard model. Because the main emphasis of model adjustment was directed toward the Santa Fe Group aquifer properties, the criterion for assessing sensitivity was the same as the criterion for adjusting Santa Fe Group properties. The criterion was the mean absolute difference between the model-derived hydraulic heads and the 90 measured hydraulic heads outside of the shallow part of the aquifer in the Mesilla Valley (marked with an asterisk in table 6). The properties tested were specific yield, specific storage, and hydraulic-conductivity values of each hydrostratigraphic unit of the Santa Fe Group. Average differences between model-derived and measured hydraulic heads are listed in table 7. The results of sensitivity tests are discussed in order of ascending mean absolute Most tests yielded mean absolute differences greater than that of the standard except when storage was set to 1.5 times the standard value, in which case the mean absolute difference was 0.01 foot less than that of the standard. The insensitivity of hydraulic head to changes in storage makes adjustment of storage on the basis of this criterion futile.

The main conclusion to be drawn from these tests is that the model is not very sensitive to changes in Santa Fe Group aquifer properties even though the sites used in the averages were selected purposefully to increase sensitivity. None of the mean absolute differences are greater than 0.8 foot from the mean absolute of the standard. This insensitivity was found by Frenzel and Kaehler (1990, p. 90 and 93). The implication of this is that model adjustment does not provide a good method of determining Santa Fe Group aquifer properties in the absence of good stress-response data for the West Mesa.

Additional sensitivity tests were conducted to assess the importance of simulated aquifer properties in determining the effects of existing nonirrigation withdrawals on flow in the Rio Grande. Hydraulic diffusivity (defined as transmissivity divided by storage coefficient) was increased by multiplying values of hydraulic conductivity and transmissivity by 1.5 and dividing specific yield and storage coefficients by 1.5. This increased hydraulic diffusivity by a factor of 2.25. The results are shown in figure 26B, in which the depletion of aquifer storage is reduced, depletion of the Rio Grande is increased, and salvaged evapotranspiration is nearly unchanged. Reducing hydraulic diffusivity by a factor of 1/2.25 had an opposite effect of similar magnitude (fig. 26C). Figure 26 shows that about 70 percent of 1985 nonirrigation withdrawals were derived from simulated depletion of the Rio Grande with the large diffusivity and about 60 percent were derived with the small diffusivity. As Frenzel and Kaehler (1990, fig. 42) noted, these percentages would be much different if the withdrawals were at some distance from the Mesilla Valley boundary.

Estimates of salvaged evapotranspiration may be much in error. The quantity of water available for salvage is determined by the quantity of evapotranspiration from nonirrigated lands. The complexity of evapotranspiration from nonirrigated lands is not well simulated by the simplistic approach that consists of a constant maximum evapotranspiration rate at land surface and a constant extinction depth with a linear variation between maximum rate and zero at extinction depth. Therefore, the quantity of water that is simulated as being available for salvage may not be reliable. Moreover, no method is known to estimate the possible error.

The estimated depletion of the Rio Grande by nonirrigation withdrawals is not greatly affected by errors in evapotranspiration. Salvaged evapotranspiration by nonirrigation withdrawals is small compared to the estimated increase in depletion of the Rio Grande. For example, if the correct salvage of evapotranspiration by nonirrigation withdrawals were one-half the estimated rate, increased depletion of the Rio Grande would be about 10 percent greater than the estimated rate.

SUMMARY AND CONCLUSIONS

Anticipated ground-water withdrawals in the Mesilla Basin may reduce flow in the Rio Grande and thus reduce the quantity of water available to downstream users. An existing model was revised and extended to assess the effects of existing ground-water withdrawals and to provide a tool for the assessment of the effects of future ground-water withdrawals.

The revised model extended the original 1915-75 simulated historical period an additional 10 years through 1985 and had modifications that were based on recent geological interpretations concerning the depth and stratigraphy of Santa Fe Group deposits. Extension of the simulated period required assembling appropriate data, such as ground-water withdrawals and streamflows, and preparing model input for the additional stress periods. Modifying the model to reflect recent geological interpretations mainly required changing the way that model-layer transmissivity values and vertical-conductance values were derived. Whereas the earlier modeling efforts specified hydraulic-conductivity values mainly on the basis of depth, hydraulic-conductivity values in the model described here are specified mainly on the basis of three hydrostratigraphic units of the Santa Fe Group.

The division of the Santa Fe Group into three units facilitates the assignment of aquifer properties on the basis of stratigraphy instead of strictly on the basis of depth. Estimated hydraulic-conductivity values were correlated with hydrostratigraphic units on the basis of the tested intervals. These estimates provided a basis for the initial assignment of aquifer properties. The model was then adjusted with particular attention focused on adjustment of Santa Fe Group aquifer properties. Simulated hydraulic-conductivity values after adjustment fell within the middle two quartiles of the estimated values except in the Las Cruces area.

The upper hydrostratigraphic unit is below the water table mainly in the northern part of the basin and is considered the most permeable of the three units because of larger grain size and less cementation. The simulated values of horizontal hydraulic conductivity of the upper unit were 30 feet per day throughout most of the basin and 12 feet per day in the Las Cruces area. The middle hydrostratigraphic unit is considered less permeable than the upper unit, judging by grain size and degree of cementation. The middle unit was divided into two parts, less than 600 feet deep and greater than 600 feet deep, because estimated values of hydraulic conductivity exhibited an inverse correlation with depth. The simulated horizontal hydraulic conductivity of the shallower part of the middle unit was 20 feet per day throughout most of the basin, but was 10 feet per day in the Las Cruces area. The simulated horizontal hydraulic conductivity of the deeper part of the middle unit was 10 feet per day and in the Las Cruces area was 5 feet per day. Although the lower unit consists of more uniform sand, cementation may make its horizontal hydraulic conductivity less than that of the other two units. The simulated horizontal hydraulic conductivity of the lower unit was 13 feet per day. Horizontal-vertical anisotropy is probably greatest in the upper and middle units because of lenticularity. The simulated horizontal-to-vertical anisotropy ratio for the upper and middle units was 500:1 and for the lower unit was 100:1.

The simulated depth of the ground-water system is significantly less than that simulated in the earlier model. The maximum saturated thickness of Santa Fe Group deposits is now considered to be about 2,500 feet, much less than the 3,800 feet that it was previously thought to be.

Except for refinements in the vertical dimension, spacial discretization was the same as that in earlier simulations. In the horizontal direction, the grid had 36 rows and 64 columns and was exactly the same as that of the previous model. In the vertical dimension, layer 5 was discarded because the basin is not as deep as was formerly thought. The thicknesses of layers 1-3 were 300, 400, and 600 feet, respectively, except where layers thinned near the edges of the model. Layer 1 simulated the first 200 feet of basin fill below the water table. Layer boundaries crossed hydrostratigraphic unit boundaries.

Assigned values of hydraulic conductivity and transmissivity were estimated on a different basis than in the earlier model. The hydraulic conductivity (layer 1) or transmissivity (layers 2-4) for a given model block was calculated on the basis of the thickness of each hydrostratigraphic unit represented in the model block.

Several refinements were made on the basis of additional precipitation and crop-acreage data. An evapotranspiration rate was estimated for each year since 1949 on the basis of individual crop acreages. The average estimated evapotranspiration from irrigated lands was revised upward to account for winter transpiration by crops such as alfalfa, pasture, and pecans. This resulted in an average annual estimated evapotranspiration rate of 2.4 acrefeet per acre, almost 10 percent greater than estimated by the previous model. Because the model was adjusted to match measured river depletions, the increase in estimated evapotranspiration from irrigated lands resulted in a decrease in simulated evapotranspiration from nonirrigated lands. The model-derived value for evapotranspiration from nonirrigated lands was about 1.9 acre-feet per acre of nonirrigated land for 1975 as compared to the value of 2.5 acre-feet per acre of the previous model.

The purpose of this model was to estimate the effects of ground-water withdrawals on flow in the Rio Grande. By comparing the mass balances of simulations with and without nonirrigation ground-water withdrawals, the sources of those withdrawals were estimated. Simulated nonirrigation ground-water withdrawals were estimated to be about two-thirds derived from increased depletion of the Rio Grande. About one-sixth was estimated to have been derived from depletion of aquifer storage and about one-sixth from salvaged evapotranspiration. These proportions are valid for most of the time since the late 1950's, after which time simulated nonirrigation withdrawals ranged from about 35 cubic feet per second (1958-60) to about 60 cubic feet per second (1980-85). These results were not very sensitive to changing hydraulic diffusivity by a factor of 2.25, but would be more sensitive for well fields located more distant from the Mesilla Valley than are existing (1980's) well fields.

REFERENCES

- Blaney, H.F., and Hanson, E.G., 1965, Consumptive use and water requirements in New Mexico: New Mexico State Engineer Technical Report 32, 82 p.
- Conover, C.S., 1954, Ground-water conditions in the Rincon and Mesilla Valleys and adjacent areas in New Mexico: U.S. Geological Survey Water-Supply Paper 1230, 200 p.
- Frenzel, P.F., 1992, Listings of model-input values for the simulation of ground-water flow in the Mesilla Basin, Doña Ana County, New Mexico, and El Paso County, Texas: U.S. Geological Survey Open-File Report 91-455, p.. 2 diskettes.
- Frenzel, P.F., and Kaehler, C.A., 1990, Geohydrology and simulation of ground-water flow in the Mesilla Basin, Doña Ana County, New Mexico, and El Paso County, Texas, with a section on Water quality and geochemistry, by S.K. Anderholm: U.S. Geological Survey Open-File Report 88-305, 179 p. (also to be available as U.S. Geological Survey Professional Paper 1407-C [in press]).
- Gates, J.S., White, D.E., and Leggat, E.R., 1984, Preliminary study of the aquifers of the lower Mesilla Valley in Texas and New Mexico by model simulation: U.S. Geological Survey Water-Resources Investigations Report 84-4317, 21 p.
- Hawley, J.W., 1984, Hydrogeologic cross sections of the Mesilla Bolson area, Doña Ana County, New Mexico, and El Paso County, Texas: Socorro, New Mexico Bureau of Mines and Mineral Resources Open-File Report 190, 10 p.
- Hearne, G.A., and Dewey, J.D., 1988, Hydrologic analysis of the Rio Grande basin north of Embudo, New Mexico, Colorado and New Mexico: U.S. Geological Survey Water-Resources Investigations Report 86-4113, 244 p.
- Hernandez, J.W., Mapel, C.L., and Enis, P.E., 1987, Community 40-year water plan for the County of Doña Ana, New Mexico, 1980-2030: Las Cruces, New Mexico State University, 73 p.
- King, W.E., Hawley, J.W., Taylor, A.M., and Wilson, R.P., 1971, Geology and ground-water resources of central and western Doña Ana County, New Mexico: Socorro, New Mexico Bureau of Mines and Mineral Resources Hydrologic Report 1, 64 p.
- Lansford, R.R., Ben-David, Shaul, Gebhard, T.G., Jr., Brutsaert, W.F., and Creel, B.J., 1974, An analytical interdisciplinary evaluation of the utilization of the water resources of the Rio Grande in New Mexico, lower Rio Grande region: Las Cruces, New Mexico Water Resources Research Institute Report 024, 109 p.
- Lansford, R.R., Sorensen, E.F., Gollehon, N.R., Fisburn, M., Loslebon, L., Creel, B.J., and West, F.G., 1980, Sources of irrigation water and irrigated and dry cropland acreages in New Mexico, by county, 1974-1979:

 Las Cruces, New Mexico State University Agricultural Experiment Station Research Report 422, 39 p.

REFERENCES -- Continued

- Leggat, E.R., Lowry, M.E., and Hood, J.W., 1962, Ground-water resources of the lower Mesilla Valley, Texas and New Mexico: Texas Water Commission Bulletin 6203, 191 p.
- Lizarraga, S.P., 1978, A non-linear lumped-parameter model for the Mesilla Valley, New Mexico: Socorro, New Mexico Institute of Mining and Technology, unpublished M.S. thesis, 43 p.
- Lohman, S.W., 1972, Ground-water hydraulics: U.S. Geological Survey Professional Paper 708, 80 p.
- McDonald, M.G., and Harbaugh, A.W., 1988, A modular three-dimensional finite-difference ground-water flow model, <u>in</u> Modeling techniques: Techniques of Water-Resources Investigations of the U.S. Geological Survey, book 6, chap. Al, various pagination.
- Miller, R.S., 1988, User's guide for RIV2--A package for routing and accounting of river discharge for a modular, three-dimensional, finite-difference, ground-water flow model: U.S. Geological Survey Open-File Report 88-345, 33 p.
- Myers, R.G., and Orr, B.R., 1985, Geohydrology of the aquifer in the Santa Fe Group, northern west mesa of the Mesilla Basin near Las Cruces, New Mexico: U.S. Geological Survey Water-Resources Investigations Report 84-4190, 37 p.
- Nickerson, E.L., 1989, Aquifer tests in the flood-plain alluvium and Santa Fe Group at the Rio Grande near Cañutillo, El Paso County, Texas: U.S. Geological Survey Water-Resources Investigations Report 89-4011, 30 p.
- Peterson, D.M., Khaleel, Raz, and Hawley, J.W., 1984, Quasi three-dimensional modeling of groundwater flow in the Mesilla Bolson, New Mexico and Texas: Las Cruces, New Mexico Water Resources Research Institute Technical Completion Report, Project No. 1-3-45645, October 1984, 185 p.
- Richardson, G.L., Gebhard, T.G., and Brutsaert, W.F., 1972, Water-table investigation in the Mesilla Valley: Las Cruces, New Mexico State University Engineering Experiment Station Technical Report 76, 206 p.
- Seager, W.R., Hawley, J.W., Kottlowski, F.E., and Kelley, S.A., 1987, Geology of east half of Las Cruces and northeast El Paso 1-degree by 2-degree sheets, New Mexico: Socorro, New Mexico Bureau of Mines and Mineral Resources Geologic Map 57, scale 1:124,000, 2 sheets.
- Slichter, C.S., 1905, Observations on the ground waters of the Rio Grande valley: U.S. Geological Survey Water-Supply Paper 141, 83 p.
- Sorensen, E.F., 1977, Water use by categories in New Mexico counties and river basins, and irrigated and dry cropland acreage in 1975: New Mexico State Engineer Technical Report 41, 34 p.

REFERENCES -- Concluded

- Spiegel, Zane, 1958, Ground-water trends in New Mexico: New Mexico Professional Engineer, March, pt. 1, p. 8-12, and April, pt. 2, p. 8-11.
- Taylor, A.M., 1967, Geohydrologic investigations in the Mesilla Valley, New Mexico: Las Cruces, New Mexico State University, unpublished M.S. thesis, 130 p.
- U.S. Department of Commerce, Bureau of the Census, 1952, Census of population, 1950, v. II, Characteristics of the population: pt. 31, New Mexico, 138 p.
- _____1982, Census of population, 1980, v. 1, Characteristics of the population, Chapter A, Number of inhabitants: pt. 33, New Mexico, 37 p.
- Wilson, C.A., and White, R.R., 1984, Geohydrology of the central Mesilla Valley, Doña Ana County, New Mexico: U.S. Geological Survey Water-Resources Investigations 82-555, 144 p.
- Wilson, C.A., White, R.R., Orr, B.R., and Roybal, R.G., 1981, Water resources of the Rincon and Mesilla Valleys and adjacent areas, New Mexico: New Mexico State Engineer Technical Report 43, 514 p.

Table 1.--Estimated values of hydraulic conductivity of the Santa Fe Group

[Open interval is the depth below land surface of the interval of hydraulic connection to the aquifer. Both intervals of well screen and open hole are included. K is the estimated hydraulic conductivity. Adjusted interval is the open interval below the water table. No adjustment was made for wells in the valley where the water table is relatively shallow (about 15 feet). Depth is the approximate depth from the water table to the middle of the open interval. Depth, in reference to plates 16 and 17 of J.W. Hawley and R.P. Lozinsky (New Mexico Bureau of Mines and Mineral Resources, written commun., 1988), was used to determine the hydrostratigraphic unit. Reference A is Nickerson (1989, p. 19 and 24), B is Myers and Orr (1985, p. 16), and C is Leggat and others (1962, p. 32; table 1). If no reference is shown, values are from Frenzel and Kaehler (1990, table 1)]

Latitude	Longitude	Well number	Open interval (feet)	K (feet per day)	Adjusted interval (feet)	Depth (feet)	Refer- ence
	Upper h	ydrostratigraphic	unit of	the San	ıta Fe Group)	
32°22′23"	106°49′16"	22S.1E.22.444	252-273	2		262	
32°09′42"	106°44′17"	25S.2E.4.141	242-262	5		252	
32°19′17"	106°48′20"	23S.1E.11.214	384-404	7		394	
32°19′17"	106°48′20"	23S.1E.11.214	510-530	7		520	
32°10′47"	106°43′26"	24S.2E.36.131	392-412	8		402	
32°18′32"	106°45′13"	23S.2E.17.243	410-700	10	280-570	425	
32°18′56"	106°45′28"	23S.2E.8.433	430-716	12	280-566	423	
32°18′19"	106°44′52"	23S.2E.16.314	381-591	13	231-441	336	
32°22′23"	106°49′16"	22S.1E.22.444	504-525	18		514	
32°18′30"	106°47′30"	23S.1E.13.411B	429-629	19		529	
32°20′09"	106°45′23"	23S.2E.5.321	392-620	20	212-440	326	
32°19′46"	106°50′28"	23S.1E.4.434	335-355	25		345	
32°16′42"	106°46′04"	23S.2E.30.243A	205-225	26		215	
32°19′14"	106°46′33"	23S.2E.7.122	213-360	26	153-300	226	
32°13′08"	106°45′38"	24S.2E.17.322	180-464	38		322	
32°19′14"	106°46′25"	23S.2E.7.411	281-381	44	221-321	271	
32°14′10"	106°46′27"	24S.2E.7.231	170-460	47		315	
32°16′42"	106°46′04"	23S.2E.30.243A	310-330	47		320	
32°16′42"	106°46′04"	23S.2E.30.243A	430-450	48		440	
32°16′28"	106°45′58"	23S.2E.29.331	243-458	50		350	
32°19′17"	106°48′20"	23S.1E.11.214A	465-485	55		475	
32°13′35"	106°47′21"	24S.1E.13.221A	140-370	60		255	
32°12′36"	106°44'45"	24S.2E.21.123	170-480	68		325	

Table 1.--<u>Estimated values of hydraulic conductivity of the</u>
<u>Santa Fe Group</u>--Continued

Latitude	Longitude	Well number	Open interval (feet)	K (feet per day)	Adjusted interval (feet)	Depth (feet)	Reference
		ydrostratigraphic ss than 600 feet)	
32°07′37"	106°39′57"	25S.3E.17.111A	437-457	2		447	
32°03′37"	106°39′13"	26S.3E.8.143	400-420	3		410	
32°09'42"	106°44'17"	25S.2E.4.141	505-525	5		515	
32°06′29"	106°42′51"	25S.2E.26.114	251-272	7		262	
32°02′45"	106°37′47"	26S.3E.15.322	310-330	10		320	
31°56′25"	106°39′17"	27S.3E.20.432	195-215	12		205	
32°10′47"	106°43′26"	24S.2E.36.131	507-527	12		517	
32°06′29"	106°42′51"	25S.2E.26.114	503-524	16		514	
31°56′25"	106°39′17"	27S.3E.20.432	450-470	18		460	
32°00′54"	106°53′39"•	26S.1W.25.414	443-563	24	53-173	113	
31°51′11"	106°39′18"	28S.3E.20.432	163-320	24	53-210	132	
31°50′50"	106°38′49"	28S.3E.28.114	240-350	25	110-220	165	
31°50′46"	106°39′29"	28S.3E.29.231	201-350	26	71-220	146	
31°57′10"	106°36′23"	JL-49-04-423	97-200	26		148	Α
32°04′14"	106°39′58"	26S.3E.6.442	307-597	35		452	
32°13'24"	106°43′26"	24S.2E.15.231A	463-484	36		474	
32°13′14"	106°45′10"	24S.2E.17.423A	310-680	41		495	
32°05′40"	106°36′40"	25S.3E.28.434	225-245	45		235	
31°50′44"	106°38′08"	28S.3E.28.241	135-285	51	75-225	150	
32°02′45"	106°37′47"	26S.3E.15.322	565-585	100		575	
	Middle h	ydrostratigraphic re than 600 feet	unit of the	the San	ta Fe Group)	
					cas ₁ o ₁		
32°03′37"	106°39′13"	26S.3E.8.143	945-965	1		955	
32°09'42"	106°44′17"	25S.2E.4.141	660-680	2		670	
32°07′37"	106°39′57"	25S.3E.17.111A	675-685	2		680	
32°02′45"	106°37′47"	26S.3E.15.322	820-840	3		830	
32°22′23"	106°49′16"	22S.1E.22.444	672-693	4		682	
32°16′42"	106°46′04"	23S.2E.30.243A	650-670	5		660	
32°02′45"	106°37'47"		1,170-1,190			1,180	
32°05′40"	106°36′40"		L,200-1,220			1,210	
32°06′29"	106°42′51"	25S.2E.26.114	651-672	8		662	
32°19′17"	106°48′20"	23S.1E.11.214	640-660	9		650	

Table 1.--Estimated values of hydraulic conductivity of the Santa Fe Group--Concluded

Latitude	Longitude	Well number	Open (interval (feet)	K (feet per day)	Adjusted interval (feet)	Depth (feet)	Refer- ence
32°08′26"	106°51′12"	25S.1E.16.114	600-1,650	12	250-1,300		
32°02′45" 31°56′25"	106°37′47" 106°39′17"	26S.3E.15.322 27S.3E.20.432	1,050-1,070 640-660	13 14		1,060	
32°05′40"	106°36'40"	25S.3E.28.434	730-750	14 14		650 740	
32°16′40"	106°52′20″	23S.1E.30.422	720-1,220	14	400-900	650	В
J2 10 40	100 32 20	233.16.30.422	720-1,220	14	400-700	030	ь
32°18′30"	106°47′30"	23S.1E.13.411	600-620	19		610	
32°18′30"	106°47′30"	23S.1E.13.411	961-981	19		971	
32°02'45"	106°37′47"	26S.3E.15.322	670-690	29		680	
31°47′24"	106°35′07"	29S.3E.13.223	300-320	1	120-140	130	
32°18′30" 32°03′37"	106°47′30" 106°39′13"	23S.1E.13.411 26S.3E.8.143	1,448-1,468	1		1,458	
32°18′30"	106°39°13"	23S.1E.13.411	1,660-1,680 1,260-1,280	2 3		1,670 1,270	
31°47′24"	106°35′07"	29S.3E.13.223	390-410	8	210-230	220	
31°57′17"	106°36′22"	JL-49-04-401	510-900	11		705	Α
31°57′33"	106°36'45"	JL-49-04-106	544-1,090	14		817	С
31°58′07"	106°36′29"	JL-49-04-105	544-950	16		747	С
31°57′17"	106°36′40"	JL-49-04-419	585-1,050	17		817	С
31°57′03"	106°36′43"	JL-49-04-402	586-1,060	17		823	С
31°57′58"	106°36′57"	JL-49-04-104	650-1,150	20		900	C
32°03′37"	106°39′13"	26S.3E.8.143	1410-1,430	34		1,420	

Table 2.--Values used for estimating net irrigation flux

EXPLANATION

- Adjusted net diversion is the total diverted water at Leasburg and Mesilla minus water returned to the river or drains (net diversion) minus estimated evaporation from canals (4,000 acre-feet).
- <u>Growing-season precipitation</u> is mean precipitation during April through October at all reporting sites.
- Annual precipitation is the mean of the year's precipitation at all reporting sites.
- Agricultural effective precipitation is 0.9 times growing-season precipitation, times the total irrigated acreage, divided by 12. These multipliers are the same as those of Frenzel and Kaehler (1990, p. 37).
- Nonagricultural effective precipitation is 0.2 times the annual precipitation, times the nonirrigated part of the valley, divided by 12.
- Project-irrigated acreage is the acreage irrigated in the Bureau of Reclamation project (sources of irrigated acreage: 1915-29, Frenzel and Kaehler, 1990, fig. 6; 1930-46, Conover, 1954, table 4; 1947-75, Wilson and others, 1981, table 10; 1976-78, R.K. Patterson, U.S. Bureau of Reclamation, written commun., 1983; 1979, Angel Colon, El Paso County Water Improvement District, written commun., 1988; 1980-85, Edd Fifer, El Paso County Water Improvement District, written commun., 1988, for the Texas part of the Mesilla Valley, and Elephant Butte Irrigation District records, Bureau of Reclamation Form 7-2045 for the New Mexico part).
- <u>Total irrigated acreage</u> is project-irrigated acreage plus 3,000 acres for 1958 and later.
- Evapotranspiration is evapotranspiration rate times total irrigated acreage.

 Net irrigation flux (curve E of fig. 8) is adjusted net diversion (curve A of fig. 8) plus agricultural effective precipitation (curve B of fig. 8), plus nonagricultural effective precipitation (curve C of fig. 8), minus evapotranspiration (curve D of fig. 8).
- Model flux is net irrigation flux divided by the area of the model blocks, in acres, that simulate net irrigation flux (fig. 17). The area of net irrigation flux was 114,000 acres for stress periods 2-22. For stress period 1, the area of net irrigation flux was divided into two parts. The part north of the State line at Anthony was assigned a net flux of 0.68 foot to account for irrigation (Frenzel and Kaehler, 1990, p. 66), and the part south of the State line was assigned a net flux of 0.30 foot to account for precipitation on nonirrigated acreage.

Values in this table for stress periods 1 and 2 are approximate average values. Evapotranspiration rates for years before 1949 are average values, as is growing-season precipitation for years before 1931.

Table 2.--Values used for estimating net irrigation flux--Continued

Model r flux) (feet)	0.68,0,30	0.97 0.97	1.93 1.93 1.93 2.02 1.20 2.27 1.90 1.95 1.95 2.09 2.09
Net Irri- gation flux (cubic feet per second)	.69	152.	303. 303. 303. 318. 189. 357. 256. 307. 283. 328. 316.
Net Irri- gation flux (acre-	50, 150	110,015	219,717 219,607 219,552 230,418 137,107 258,315 256,824 185,513 222,109 204,835 237,740 259,599
Evapo- transpi- ration (acre- feet)	105,600	1.10,400	180,000 180,000 180,000 183,295 184,102 184,946 164,652 149,220 179,551 186,264 176,398 177,682
Total irri- gated acreage (acres)	44,000	46,000	75,000 75,000 76,373 76,722 76,709 77,061 68,605 62,175 74,813 77,610 73,499
Project- irri- gated acreage (acres)	44,000	46,000 period 2	75,000 75,000 76,373 76,722 76,709 77,061 68,605 62,175 74,813 77,610 73,499 74,034
Evapo- transpi- ration rate (feet)	790 8,960 2,40 44,000 Average for stress period 1	2,40 stress	2.40 2.40 2.40 2.40 2.40 2.40 2.40 2.40
ctive Itation Non- agri- cul- tural (acre- feet)	8,960 rage for	735 8,680 Average for	138 5, 280 138 5, 170 138 5, 115 186 3, 627 227 6, 912 112 4, 605 143 3, 033 143 3, 033 148 3, 551 121 5, 417 120 3, 453 132 4, 132
Effective precipitation Non- Agri- agri- cul- cul- tural tural (acre- (acre- feet) feet)	20,790 Ave	21,735 8,680 Average fo	35, 438 35, 438 35, 438 36, 086 51, 327 39, 812 22, 194 17, 443 17, 443 31, 548 42, 721 23, 820 34, 372
tation Annual (inch-es)	8.40	8.40	9.60 9.40 9.30 6.88 13.26 8.83 4.71 4.71 7.01 9.50 7.01
Precipitation Growing season Annual (inch- (inch-	6,30	6,30	6.30 6.30 6.30 6.30 7.32 10.52 10.52 7.75 6.08
Adjusted net diver- sion (acre- feet)	915-19 126,000	1920-26 190,000	359, 000 359, 000 359, 000 263, 000 398, 000 377, 000 401, 000 276, 000 356, 000 410, 000 371, 000
Year	1915-19	1920-26	1927 1928 1929 1930 1933 1935 1936 1936 1938

Table 2,--Values used for estimating net irrigation flux--Continued

: Model ir flux I) (feet)	2.04 2.34 2.34 2.54 2.73 2.25 2.24	2.18 2.36 2.13 2.22	0.58 0.80 0.64	0.31 -0.64 -0.68 0.18
Net irri- gation flux (cubic feet per second)	321. 368. 368. 400. 429. 398.	343. 371. 334.	91. 126. 101.	-48. -100. -107. 29.
Net irri- gation flux (acre- feet)	232, 769 266, 713 266, 741 289, 951 311, 118 288, 566 254, 847	248, 156 268, 703 242, 260	65,662 90,938 73,184	-34,914 -72,731 -77,410 20,887
Evapo- transpi- ration (acre- feet)	189, 266 194, 921 198, 348 198, 828 199, 850 201, 386 202, 562	199, 673 201, 553 205, 498	199, 578 204, 332 208, 183	196,058 198,155 188,433 184,654
Total irri- gated acreage (acres)	78,861 81,217 82,645 82,845 83,271 83,911	83, 197 84, 686 84, 567	84,210 84,785 85,672	78,738 78,633 74,775 75,369
Project- irri- gated acreage	78,861 81,217 82,645 82,845 83,271 83,911 84,401	83,197 84,686 84,567 period 5	84,210 84,785 85,672 period 6	78,738 78,633 74,775 75,369
Evapo- transpi- ration rate (feet)	2.40 2.40 2.40 2.40 2.40 2.40 2.40 2.40	2,40 2,38 2,43 stress	2,37 2,41 2,43 stress	2.49 2.52 2.52 2.45
itation Non- agri- cul- tural (acre- feet)	568 7, 367 157 4,477 397 2,692 102 4,377 224 2,745 327 2,626 314 2,596 Average for	274 1,554 566 3,590 231 2,527 Average for	495 1,745 390 3,180 581 1,786 Average for	39 3,404 72 3,852 104 2,619 08 5,433
Effective precipitation Non- Agri- agri- cul- cul- tural tural (acre- (acre- feet) feet)	68, 668 49, 157 27, 397 48, 402 35, 224 29, 327 24, 814	23,274 41,666 37,231 Ave	17,495 38,090 23,581	38, 739 39, 572 19, 404 45, 108
Annual (Inches)	15.17 10.03 6.37 10.44 6.66 6.54 6.60	3.76 9.24 6.47	4.40 8.22 4.80	6.98 7.87 4.73 9.99
Precipitation Growing season Annua (inch- (inches)	11.61 8.07 4.42 7.79 5.64 4.66 3.92	3,73 6,56 5,87	2.77 5.99 3.67	6.56 6.71 3.46 7.98
Adjusted net diver- slon (acre-	346,000 408,000 435,000 473,000 458,000	423,000 425,000 408,000	246,000 254,000 256,000	119,000 82,000 89,000 155,000
Үөаг	1942 1942 1943 1944 1945 1946	1948	1951 1952 1953	1954 1955 1956 1957

Table 2.--Values used for estimating net irrigation flux--Continued

Mode! flux (feet)	1.93 1.61 1.82 1.79	1.51	1.87 1.01 1.44	-0.76 -0.76	0.41 1.32 0.86	0.78 1.20 0.99	2.04 1.80 0.98 1.61	0.16 0.16
Net irri- gation flux (cubic feet per	304. 253. 286.	237.	294 . 159 .	-119.	64 . 208 .	123.	321. 283. 153.	25.
Net irri- gation flux (acre- feet)	219, 926 183, 139 207, 307	171,917	212,801	-86,091	46,266	89,432 136,812	232,891 204,874 111,182	18,377
Evapo- transpi- ration (acre- feet)	196,728 192,258 195,054	195, 152	194,604	190,669	187,676	170,453	180,072 183,783 183,957	177,891
Total irri- gated acreage (acres)	79,647 79,775 79,614	79,654	81,085	80,113	77,874	73,471	78,634 78,877 77,293	74,744
Project- irri- gated acreage (acres)	76,647 76,775 76,614 period 8	76,654 period 9	78,085 78,742 period 10 -	77,113 period 11 -	74,874 72,719 period 12 -	70,471 74,370 period 13 -	75,634 75,877 74,293 period 14 -	71,744 beriod 15 -
Evapo- transpi- ration rate (feet)	2.47 2.41 2.45 stress	2.45 stress	2.40 2.37 stress	2,38 stress	2.41 2.38 stress	2,32 2,25 stress	2.29 2.33 2.38 stress	01 7,167 2,38 71,744 Average for stress period 15
Effective precipitation Non- Agri- agri- cul- tural tural acre- (acre- feet) feet)	326 6,828 559 2,738 362 2,999 Average for	304 3,765 Average for	705 3,701 235 2,035 Average for	547 2,231 Average for	523 3,419 304 5,294 Average for	150 3,735 170 5,825 Average for	.06 4,557 115 2,242 195 3,444 Average for	67,101 7,167 Average for
Effec precipi Agri- cul- tural (acre- (61,826 28,659 24,362	31, 304 Ave	37,705 23,235 Ave	19,347 Avei	25,523 49,804 Ave	29, 150 42, 070	43,406 21,415 32,695	67, 101 Ave
tation Annual (inch-	14.45 5.82 6.34	7.97	8.25 4.65	4.80	6.81 9.84	6.49	9.31 4.62 6.73	12,93
Precipitation Growing season Annua (inch- (inch	10,35 4,79 4,08	5,24	6.20	3,22	4.37 8.77	5.29	7.36 3.62 5.64	11,97
Adjusted net diver- sion (acre- feet)	348,000 344,000 375,000	332,000	366,000 284,000	83,000	205,000 276,000	227,000 263,000	365,000 365,000 259,000	122,000
Year	1958 1959 1960	1961	1962	1964	1965	1967	1969	1972

Table 2.--Values used for estimating net irrigation flux--Concluded

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet

									Mode	l column	E									
	-	2	3	4	5	9	7	İ	6	2	_					1			1	20
	21	22	23	24	25	56	27	28	59	ଛ	31	32	33	¥	35	Ж	37	82	39	40
Model	41	42	43	44	45	46	47		49	20										09
NO.	61	62	63	64																
						hickness	5	river	alluvium assigned	um ass		to model	el lay	- Je						
Row 16	0000			0000	• • •	· · ·	• • •	• • •	70. 0.	70. 0.	70. 0.	• • •	• • •	• • •				• • •		• • •
Row 17	0. 70.	0 0 0	70. 0.	0. 70.	0. 70.	0. 70.	70. 0. 70.	70. 0.	70. 0.	70. 0.	70. 0.	0.00	.00.00	70. 0.	• • •		• • •	• • •	• • •	
Row 18	70.	. 0. . 0.	. 0. 0.	0°. 0°.	70. 0. 70.	70. 0.	70. 0.	70. 0. 70.	70. 0.	70. 0.	70. 0.	70. 0.	70. 0.	70. 0.	0°. 0°.	70. 0.	• • •	70.	0°.	0. 70. 0.
Row 19	0. 70.	. 0 0 0	70. 0. 70.	70. 70.	70. 0.	70. 0.	70. 0.	70. 0. 70.	70. 0. 70.	70. 0. 70.	70. 0.	70. 0. 70.	70. 0.	0.00	70. 0.	70. 0.	70. 70. 0.	70. 70. 0.	70. 0.	70. 70. 0.
Row 20	0. 70.	0. 70.	70. 0. 70.	70. 0. 70.	70. 0. 70.	70. 70.	70. 0.	70. 0. 70.	70. 0. 70.	70. 70. 70.	70. 70. 70.	70. 0.	70. 0.	70. 70.	70. 70. 0.	70. 70. 0.	70. 70. 0.	70. 70.	70. 70. 0.	70 . 70 . 70 .

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	5	CILLET									
	- ;	2 5	2.5	4 2	2 2	9 %	1	8 %	6 6	2 8	= =	22 2	13	4 2	15	92 %	12 2	18	61 6	20
Model	1 4 1	42 62	63 43	4 4 4 4 4 4	45	46	47	48	49		5 15	52	53	¥ ¼	52	8 %	57	ያ ያ	59	9
Row 21	0. 70. 70.	0. 70. 70.	70. 70. 0.	70. 70. 0.	70. 70.	70. 70.	70. 70.	70. 70.	70. 70.	70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 0.	70. 70.	70. 70.	70. 70. 70.
Row 22	0. 70. 70.	70. 70. 70.	70. 70. 0.	0. 70. 0.	0. 70.	0. 70.	0. 70.	0. 70.	0. 70.	70. 70.	70 . 70 .	70 . 70 . 70.	70. 70.	70. 70. 70.	70. 70.	70. 70. 70.	70. 70. 70.	70. 70.	70. 70.	70. 70. 70.
Row 23	0. 70. 0.	70. 70. 70.	0. 70. 0.	70.	0. 70.	0. 70. 70.	0. 70. 70.	0. 70. 70.	0. 70. 70.	0. 70. 70.	70. 70.	70. 70. 70.	70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70.	70. 70.	70. 70. 0.
Row 24	0. 70. 0.	0. 70. 0.	0. 70. 0.	70. 70.	0. 70.	0. 70. 70.	0. 70. 70.	0. 70.	0. 70. 70.	0. 70.	0. 70.	0. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70.	70. 70.	70. 70. 0.
Row 25	0. 70. 0.	0. 70. 0.	.07 .00.	70. 70.	0. 70.	0. 70.	0. 70.	0. 70.	0. 70.	0. 70.	0. 70. 70.	0. 70. 70.	0. 70. 70.	0. 70. 70.	70. 70. 70.	70. 70. 70.	70. 70. 0.	70. 0.	70. 70.	70. 70. 0.
Row 26	0. 70. 0.	.0. 70.	.0 70.	0. 70. 0.	.07 .07	.07 .07	.07 .07	.07 .07	0. 70.	.070.	0°. 0°.	0. 70. 70.	0. 70.	0. 70.	0. 70.	0 0 0	70. 70. 0.	70. 70.	70.	70. 70. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l column	CW									
	1 21	2 22	3 23	4 24	5 25	6 26	7	8	9 29		11	12	13	4 4	15 35	46 36	17	18 38	19 39	20 40
Model	41	42	63	44	45	46	47	48	49		51	52	53	54	55	26	57	58	59	09
Row 27	70.	.070.0	0. 70. 0.	.0. 70. 0.	.0. 0.	70.	0. 70. 0.	0. 70. 0.	0. 70. 0.	.0 70.	0. 70. 0.	0. 70. 0.	70.	.070.	.0 .0 .0	.0. 0.	• • •	• • •	70.	70.
Row 28	.0. 0. 0.	.070.0	0. 70. 0.	70. 0.	70.	.0. 0.	0. 70. 0.	0. 70. 0.	0. 0.	0. 70. 0.	0. 70. 0.	.0. 0.	.0. 0.	0. 70.	· · ·	• • •	• • •		• • •	70. 0. 0.
Row 29	0000	****	0. 0. 0. 0. 0. 0. Thickness of	0. 0. 0.	0. 70. 0.	0. 70. 0. hydrost		0. 0. 70. 70. 0. 0.		0. 70. 0.	0. 70. 0.	0 0 0	0. 0. 0. Group as	0. 0. 0. assigned	0. 0. 0. 0. 0. 0.	0. 0. 0.	0. 0. 0.	• • •	• • •	
Row		0000	0000	· • • •	• • •	• • •			• • •	• • •	• • •	• • •	• • •	000		• • •	• • •	• • •	• • •	• • •
Row 2	0000	0000	0000	0000	• • •	• • •	• • •	000	• • •	000	000	• • •	• • •	000	• • •	• • •	• • •	· · ·	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

	.								Mode		٤									
	- 12	7 23	23	7 7 7	25	و 20	72	78 78	29	2 R	= 5	2 22	១ ន	<u> </u>	ر الآ	2 %	3. 1.	<u>≈</u> &	5 S	6 2 8 2 9 2
Model	61	42	43 63	44 49	45	46	47	48	49	20	51	52	53	54	55	26	57	28	59	09
Row 3	, 0 0 0	50.000	50.00	0.00	800	20.					• • •		000				• • •	•••		• • •
Row 4	50.00.00	50.00.00	50.000	50.	50.	. ° ° ° ° ° ° ° ° ° ° ° ° ° ° ° ° ° ° °	20°.	• • •	• • •		• • •	• • •					• • •	• • •		• • •
Row 5	50.	50.	50.	50.	50.	50.	0 0 0	o o o	20. 0.	• • •	• • •	• • •	· · ·			• • •	• • •	• • •	• • •	• • •
Row 6	0000	50.00.00	50.	50.	50.	50.	50.	50.	0.00	0.00	0 0 0	0.00	0.00	o o o	0.00	20.	0.00	• • •	• • •	• • •
Row 7	50.	50.00.00	50.	0000	50.	50.00	50.	50.	0.00	0.00	40. 0.	800	. ° °	° ° °	20. 0.	20.	0.00	• • •	• • •	· · ·
Row 8	60 .01 .0	0000	0000	0000	00 0	0000	0000	50.	50.	0,00	0°0°	800	20.00.00	20. 0.	0.00	0 0 0	000	000	0 0 0	0.00

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

								1 1	Mode	100		1 1	ll	1 1		1 1				
	- 5	2 6	ر د د	4 6	יט גי	9 %	7		o 6	5 t								2 8	<u>6</u> 8	7 50
Model	41	42	43	44	45	46	47	48	49	R &	2 12	22	53	7. 7.	22 72	ጸ አ	57	የ ଝ	20 7	9 9
№	61	62	63	64																
Row 9	70.	70.	70.	70.	.0°	70.	70.	•09	• 09	50.	40.	30.	· 0	.	• •			.	.01	.00
	. 6 0	200.			· •	· •	• • •	• •	• •	· •	• •	· •	· •	• •	· •	• •	• • •	å °	° °	• • •
Row 10	80.	80°	80.	80.	80.0	0.0	· ·	• •	• •	.00	• •	• •	· •	.00	•••	0°.	. 04	.08 80.	0.	0.120.
	200.	200.	200.	200.	•	•	•	•	•	•	•	•	•	•0				•	•	•
Row 11	80. 0. 200.	80. 0. 200.	80. 0. 200. 0.	80. 0. 200.	0. 0. 200.	•••	• • •	ō	ō ° °		0.00	· · ·	· · ·	• • •	0°. 0°.	50.	0°. 0°.	0.0150	0. 200. 0.	0°. 0°.
Row 12	80. 200. 0.	0. 200. 0.	0. 200. 0.	0. 10. 200.	0. 30. 200.	0. 40. 200.	• • •	• • •	• • •	<u>°</u> ° °	0000	0.00	0.00	0.00	.05 .0	10. 60. 1	0. 140. 0.	.071	0. 200. 0.	0°. 200°. 0°.
Row 13	50. 200. 0.	50. 200. 0.	40. 0. 200.	50. 10. 200. 0.	40. 20. 200.	0. 40. 200.	0. 90. 200.	0. 120. 200.	• • •	• • •	• • •	• • •	· · ·	• • •	•••	0. 20. 1	.001	0.0180.	0. 200. 0.	0°. 200°. 0°.
Row 14	130. 0. 200.	130. 0. 200.	120. 0. 200.	140. 10. 200.	150. 10. 200.	140. 30. 200.	50. 70. 200.	0. 120. 200.	0. 150. 200.	0. 150. 0.	• • •	• • •	• • •		. 0 0	20.	80.0	0.	0. 160. 0.	0°. 200°. 0°.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

								Model	100	UET			1	H					
7	<u> </u>	n	4	Ŋ	9	7	æ ¦	6	2	= :	12	5			9	11	8	19	50
7	2 (23	24	25	26	27	78	59	R 8	<u>ب</u> ۲	22	33	X :	35	አ	31	8 2	<u>۾</u>	4 6
4.0	7 7	63	44 44	7	04	4	4. X	4	2	<u>.</u>	75	55			ደ	/c	%	<u></u>	9
	09	60.	70.	70.	110.	100	°	°	o	1	Į.	i	ł	o	Ì		°	°	0
	°	°	10.	10.	20.	70.	130	150.	160.					0.			140.	200	200
.,	200 . 0.	200.	200.	200.	200•	200.	200.	200•	200•	°	•	° 0	°	°	°	•	°	°	•
	50.	50.	70•	70.	50.	40.	40.	°	•	• •	•	°		°			•	°	•
	•	°	10.	0.	•09	80.	120.	150.	160.	180.	•	•		ં			150.	200	200
• •	200.	200.	200.	200.	200•	200.	200.	200.	200•	200.	•	°	°	•	.	°	°	•	°
	50.	50.	. 09	40.	40.	0	°	•	°	°	°						°	o	°
	°	°	0.	20.	50.	70.	120.	140.	150.	180.	190.						150.	200	200
	130.	0.0	130. 0.	130.	130.	130.	200.	200.	200.	200.	200.	°	•	•	° 0	•	•	•	•
	40.	40.	40.	_•	°	°	°	°											0
	°	°	10.	20.	50.	110.	120.	130.											130.
	0.	0.	130. 0.	.•	130.	130	130.	130.	130.	200.	200.	200.	•	.	•0	•			•
	°	°	°	•	°	•	°	•	°										.
	°	°	01	20.	40.	110.	120.	130.	140.				200.	20	80.				130.
0.0	50.	0	130. 0.	130.	130.	130.	130.	130.	130.	130.	130.	200.				o	•	•	.
	°	°	°	•	•	°	•	°	°	•	•	•	•	•			•	•	•
	0.05.	0°.	130.	20 .	70 . 130 .	80 . 130 .	130.	160 . 130 .	120 . 130 .	130. 130.	200 . 130.	200 . 130 .	130 . 200 .	80 . 200 .	40.	°0°	.0.	130 . 200 .	0.
	°	o	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

								1	Mode	ol column	C III									
	-	2	3	4	5	9	7		6	10	=	12	13	14		19	11	22	19	20
	21	22	23	24	25	56	27	28	53	ନ୍ଧ	31	32	33	ጃ	35	ጽ	31	82	æ	40
Model	1 ;	42	43	44	45	46	47		49	23	21	52	53	24		26	21	28	29	09
M OL	<u> </u>	79	65	40																
20	c	c	c	c	c	c	c	c	c	c	c						c	c		c
	•	• 6	. 0	20.	40,	50.	80.	. 6	. 06	100	130.						8	8		. 8. . 8.
	130	1.30 •	130	원	왕	<u>동</u>	130	ا ا	<u>용</u>	8	130	130	ا ا	<u>동</u>	<u>용</u>	8	200.	71.	•	•
	•	•	•	•																
Row 22	•	•	•	•	•	°	•	•	°	°	•	•					•	•		•
	•	•	•	20.	40.	40.	70.	80.	•06	110.	130.	130.	130.	130.	130.	130.	130	ا ا	٦ ا	130.
	0. 0.	130°.	₩ •	0.	130 .	130.	ي ق	130.	130.	8	<u>ਲ</u>	130.					130.	8		•0
					,		,	,		,	,	,	,				,			
Row 23	•	•	•	•	•	ó°	•	°	•	•	•	•	•				•	°		•
	•	°	•	20.	20.	40.	70.	80•	° 06	ਲ ਲ	130.	130	130.				130.	130.		130.
	130°	130.	130	130	130.	130.	130.	130.	130.	8	130	130.	130.	130.	ගේ	130.	130.	<u>8</u>	<u>8</u>	200.
	•	•	•	•																
Row 24	0	•	•	•	•	•	•	•	•	•	•	•					•			•
	•	•	•	•	•	8	70•	80	06	130.	130.	130.			130.		130.			130°
	<u>중</u>	<u>8</u>	<u>8</u>	<u>8</u>	130.	130.	130.	130.	130.	130	130°	130.	130.	130.		130.	<u>.</u> ਲ	130.	200.	•
	•	•	•	•																
Row 25	•	•	•	•	°	0	•	•	•	0.	20.		•					•	•	•
	•	•	•	•	•	•	. 07	• 09	130.	130.	30.	130.	130.		130.			130.	130.	130.
	- 0 0 0	130	0.00	02.	٠ 8	8	8	30.	130.	130.	ا ا		ا ا	8.		130.	200.	200.	•	•
	•	•	•	;																
Row 26	•	•	•	•	•	•	•	•	•	20.	20.	20.	40.				•	•		•
	°	°	•	•	•	40•	40.	. 02	<u>8</u>	130	130.	130.	130	130.	교	130.	130.	130.	٦ ا	130.
	호 양 호	0. 0.	<u>ရှိ</u> ဝ	<u>8</u> °	<u>용</u>	130.	<u>ද</u>	원	원	200•	200.	٠ 8	ਲੂ ਲ				200.	°		.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mod	10	٤									
	-	2	2	4	5	9	7	8	6	10 11	=	12	13	14	1	16	11	18	19	20
	21	22	23	24	25	56	27	28	53		31	22	33	× 2	35	× 2	37	8 2	8 8	40
MOL	61	62	63	64		,	}	4 0	,		5	70	C.	<u></u>		8	ñ	8	N	8
Row 27	0	0	0	0	0	0	0	0	0	0	0	01	\ \\ \\ \\ \\ \\ \\ \\ \\ \\ \\ \\ \\ \	ا 8	20.		0	0	0	0
	0.	0.02	0°.	0.	0.200.	30.	60.	80 . 200 .	130.	130.	130.	130.	130.	130.	130 . 200 .	0.0	200.	200.	200.	200.
	0	•	0	•																
Row 28	0 0	° c	0 0	° 0	0 0	0.40	000	0.10	0.	0.130	0.130	0.0	0.	20.	200	10.	10.	000	000	0.000
	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	• 0	• 0	•	• 0	• •	• •
Row 29	o s	· 6	•	0 0	0	0	0	0 8	0	0	0	0 6	0 (.	° (
	0° 0	200.	200.	200.	200.	200.	,0°.	200.	200.	200.	0.0	•002	• 0 • 0	• 0	0,0	0,0	°°°	0.0	000	• • •
Row 30	0	o			0	0	o	0	0	0	0	0	0	0	0	0	0	0	0	.
	0 0	0 0	0.	0 0	20.	8 8	60.	130.	200.	200.	00	0.0	0.0	20.	0	ō. º	• •	20.0	20.0	0
	0	o	•	0					•		•	•	•	•		•	;	•	•	•
Row 31	0 0	0 0	0 0	• 0	• •	0 0	0 0	00	0 0	• •	0 0	00	• •	• o	• o	° 0	° 0	• o	° 0	• •
		• •	• •	• •	00.	00.	•	•	•	0	•	•	· •	•	· •	0	0	•	0	•
Row 32	• •	° 0	• •	• •	• •	° °	· ·	• •	• •	• •	· ·	• •	• •	• •	° °	· ·	• o	· ·	· ·	• •
	• •		• •	• •	•	• 0	•	0	0	•	0	•	•	•	0	•	•	•	•	•

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	00	uwn									
	-	2	~	4	5	9	7		6		=	12	13	4	15	16	11	<u>8</u>	16	20
	21	22	23	24	25	26	27	28	29	ጽ	31	32	33	¥	35	8	37	82	33	40
Model	41	42	43	44	45	46	47	48	49		51	52	53	54	25	26	21	28	23	90
MOL	61	62	63	64																
Row 33	••••				• • •	•••	• • •	• • •		· · ·	• • •		• • •	• • •	• • •	• • •	• • •	· · ·		• • •
Row 34	0000	• • • •	0000	••••	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	•••	• • •	• • •	• • •	• • •	• • •	000
Row 35	0000		••••	• • • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 36	0000	• • • •		••••	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	•••	• • •	• • •	• • •	• • •
		£	Thickness	oŧ	middle	hydros	hydrostratigraphic	raphic	t un :	of the	Santa	_	Fe Group assigned	signed	1 0 mo	to model layer	yer 1			
Row 1						•••	•••		0. 200. 0.	0. 200. 0.	0. 200. 0.	0. 200. 0.	0. 200. 0.	• • •		•••	•••	•••		• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mod	el column	UWIT									
	-	2	3	4	5	9	7	8	6	2	=	12	13	14	15	16	17	18	6	20
	21	22	23	24	25	56	27	28	53	ጽ	31	32	33	¥	35	88	37	88	85	40
Model	41	42	43 63	44 64	45	46	47	48	49	20	51	52	53	54	55	26	27	28	29	09
Row 2	200.	200.	200. 200. 0.	200.	200.	200.	200. 0. 0.	200.	200.	200.										
Row 3	200.	150. 200. 0.	150. 200. 0.	160. 200. 0.	170. 200. 0.	180. 200. 0.	200. 200. 0.	200. 0. 0.	200°. 0°.	200°. 0°.	200. 0. 0.	200. 0. 0.	200. 0.	200.						
Row 4	150. 200. 0.	200.	150. 200. 0.	150. 200. 0.	150. 200. 0.	160. 200. 0.	180. 200. 0.	200. 200. 0.	200. 200. 0.	200. 200. 0.	200. 200. 0.	200. 200. 0.	200. 200. 0.	200 . 200. 0.	200 . 0. 0.	200 . 0. 0.	200.	200. 0. 0.	200.	200. 0. 0.
Row %	200.	150. 200. 0.	150. 200. 0.	150. 200. 0.	150. 200. 0.	150. 200. 0.	160. 200. 0.	170. 200. 0.	180. 200. 0.	200. 200. 0.	200. 200. 0.	200. 200. 0.	200 . 200. 0.	200 . 200. 0.	200. 200. 0.	200°. 0°.	200.	200. 0. 0.	200. 0.	200. 0. 0.
Row M	200.	200.	150. 200. 0.	150. 200. 0.	150. 200. 0.	150. 200. 0.	150. 200. 0.	150. 200. 0.	160. 200. 0.	160. 200. 0.	160. 200. 0.	160. 200. 0.	170. 200. 0.	170. 200. 0.	170. 200. 0.	180. 200. 0.	190. 200. 0.	200. 0.	200.	200.
Row 7	200°.	150. 200. 0.	160. 200. 0.	160. 200. 0.	160. 200. 0.	170. 200. 0.	170. 200. 0.	170. 200. 0.	180. 200. 0.	180. 200. 0.	190. 200. 0.	200. 200. 0.	200. 200. 0.	200. 0. 0.						

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	l column	nmr									
	-	2	3	4	5	9	7	80	6		11	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	8	31	32	33	×	35	Ж	37	82	39	40
Modei	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	56	57	58	59	09
TOW	19	62	63	64																
Row 8	140.	140.	140.	140.	140.	140.	140.	150	150.	160	170.	170.	180	180.	190	190	190	190	190	190.
	190.	200	200	200	200	200	200	200	200	200	200	200	200	200	200.	200.	200	200	200	200
	200.	200	° c	o c	°	°	•	°	•	•	•	•	•	•	• 0	•	•	•	•	•
	•	•	•	• •																
Row 9	130	130	130	130.	130	130	130	140.	140.	150.	160.	170.	200	200	200	200	200	200	190	190.
	200	200	200	200	200	200	200	200	200.	200•	200	200	200	200	200	200•	180	170	150.	140.
	0° 0°	· ·	· ·	• •	°	.	.	•	•	•	.	•	°	•	•	•	•	•	•	•0
10 10	120	120	120	120	120	5	000	000	200	190	200	200	000	190	200	200	200	200	200	200
	200	200	200	200	200	200	200	200	200.	200	200	200	200	200	200.	180.	160	120	100	80.
	0	0	0	0	0	0	0	0	•	•	0	0	0	0	0	0	0	0	0	•
	•	0	0	•																
Row 11	120.	120.	120.	120.	200.	200.			190.	190	190.	200.	200.	200.	200.	200.	200.	200.	200.	200.
	200	200	200	200	200.	200	200	200	200.	200	200	200	200	200	180.	150.	110.	50	0	•0
	o c	o c	ં c	o c	° 0	0			•	•	•0	•	•	•	•	•	•	•	• 0	•
	•	•	•	•																
Row 12	120.	200.	200	200	200.	200	200	200	200.	190•	190•	190.	190.	190.	190.	190.	200	200	200.	200.
	200	200	200	190	170.	160.	200	200.	200.	200.	200	200	200.	200.	170.	140.	• 09	8	o ·	• •
	• • •	• •	• •	• •	• 0	°	•	•	•	•	•	°	°	•	•	•	•	• •	°	•
Row 13		150.	160.	150.	160	200.	200	200.	200.	200.	200.	200.	200.	200.	200.	200.	200.	200	200.	200.
	200.	200	200	190	180.	160.	110.	80	200.	200	200	200	200	200	200.	180.	100	20.	0	•0
	• •	. .	°	•	°	°	°	°	• 0	• •	°	•	•	•	•	°	.	•

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	8	nmn									
	-	2	2	4	5	9	7	8	6		=	12	13	14	15	16	17	_∞	6	20
	21	22	23	24	25	56	27	28	53	ନ୍ଧ	31	32	33	¥	35	8	37	æ	33	40
Model	4	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	57	28	29	09
¥0	19	62	63	64																
Row 14	70.	70.	80.	•09	50.	•09														200.
	200	200•	200	190.	190.	170.	130.	80.	50.	50	200	200	200	200	190	180.	120.	•09	40.	•
	o 0	o o	o o	.	•	•													•	•
	•	•	•	•																
Row 15	140.	140	140.	130°	30.	90	100	200											200	200.
	200	200•	200•	190.	190.	180.	130.	70.	50.	40	200°	200.	200.	200	.061	170.	130.	•09	•	•
	°	o	•	•	°	•	•	•											•	•
	•	.	.	•																
Row 16	150.	150.	150.	130.	130	150.	160.	160.	130.	130.	130	200.							200.	200.
	200.	200	200	.061	.061	140.	120.	80.	50.	40.	20.	200.	200	200.	200	160.	140.	50.	•	•
	•	°	•	•	•	•	•	•	•	•	•	•						•	•	•
	o o	o	o°	°																
Row 17	150.	150.	150.	140.	160	160.	130	130.	130.	130.	130		130				200		200.	200.
	200	200	200.	190•	180.	150.	130.	80.	•09	50.	20.	0	200.	190	170.	150.	130.	50.	•	•
	°	•	•	.	.	.	•	•	°	°	°		•				•		•	•
	°	o	o'	°																
Row 18	160.	160.	160.	160.	130.	130.	130.	130.	130.	130.	130.	130.	130.	130.	130.	130.	200.	200	200.	200.
	200	200	200•	190•	180.	150.	° 06	80.	70 °	•09	20.	•	•	170.			10.	40.	•	•
	• •	• •	•••	•••	.	•	°	•	•	°	.	.	°	•			•	°	•	•
Row 19	150.	150.	80.	100	130	130.	130.	130	130.	130	130.	130.	130.	130.	130.	130.	130.	130.	130.	130.
	200	200	200	190•	180.	160.	•06	80.	70.	•09	ಜ್ಞ	°	•	•	150.	120.	• 06	40	•	°
	o	o o	o o	o o	.	°	°	•	•	•	°	°	•	°	•	•	°	•	•	•
	•	•	•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	el column	ımı									
	-	2	3	4	5	9	7	8	6	10	11	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	ጽ	31	32	33	×	35	Я	37	88	39	40
Model	41	42	43	44	45	46	47	48	49	22	51	52	53	54	55	26	21	82	29	09
NO.	19	62	63	64																
Row 20	130	140.	70.	80.	%	130	130.	30.	<u>8</u>	130.	130.	양	130.				130.	130.	130.	130.
	130.	200	200	200•	180.	130.	120.	110.	40.	10.	•	•	•	•	50.	90•	70.	20.	•	•
	.	°	·°	.	.	•	.	.	.	°	<i>.</i>	.	.				.	•	.	•
	°	°	°	°																
Row 21	50.	%	70.	70.	70.	70.	130	130°	138.	130.	130	동	130	130.		130.	130.	130.	130	130.
	130	130	130.	110.	•06	80.	50.	50.	40.	20.	•	•	•	•	•	•	•	•	•	•
	•	.	•	•	•	•	•	.	.	•	.	•	•	•		.	•	.	•	•
	.	.	50•	50.																
Row 22	8	ଞ୍ଚ	50.	120.	120.	120.	150.	200	200	130	130°	<mark>ද</mark>	130	130.		130.	130	130.	130	130.
	130	130.	130	110.	•06	• 06	•09	50.	40.	20.	•	•	•	•	•	•	.	.	•	°
	•	.	•	•	•	•	.	.	°	°	°	•	.	•		•	•	ċ	.	•
	°	°	20.	50.																
Row 23	•	•	70.	70.	80.	80.	180.	200	200	200.	130.	130.	130.	130.		130.	130.	130.	130.	130.
	130	-8 -8	130.	110.	110.	%	•09	50.	40.	•	•	°	•	•	•	•	•	•	•	•
	ં છ	ં છ	o 8	.	°	.	<i>•</i>	°	°	°	°	•	°	<i>•</i>		.	•	•	•	•
	• 0	• 08	•00	•																
Row 24	•		•	•	40.	40.	• 06	130.	200	200.	200	200	130.	30.	130.	130.	130.	130.	130.	130.
	130		ا گ	130°	130	100	•09	20.	40.	.	ċ	ċ	ċ	ċ		•	ċ	•	ċ	.
	001	.00	001	<i>.</i> .	o	°	°	o	°	°	°	°	°	o		•	o	ċ	°	120.
Row 25	ď	ď	d	d	d	Ċ	20	6	160.	190	180	200	200	200-		55	130	130	30	130.
	130	ا ا	ا ا	130	130	13 8	09	70.	0	0	0	0	0	0		0	o	o	o	0
	0	· 0	· 0	.	.	.	•	.	0	°	.	.	•	.	•	•	.	•	170.	170.
	150	140	o	°																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	1 column	mC Cm									
	-	2	~	4	5	٥	_		6	2	_	1				Ì		l	19	20
Model	21 41 61	22 42 62	23 43 63	24 44 64	25	26 46	27	28 48	49	20 20	51	32 52	53	¥ ¥	35	36 56	57	38 58	39	40
Row 26	0. 130. 0.	130.	0. 130. 0.	0. 130. 0.	0. 130. 0.	90.0	.06 .09	0,00	06 0	.0 .0 .0	180.	180.	160.	00.0	200.	200.	130. 0.	130 . 0. 200.	130. 0. 170.	130. 0.
Row 27	0. 130. 0.	0. 130. 0.	0. 130. 0.	0. 130. 0.	0. 130. 0.	.000	0.	50.	• • •	50.	130. 0.	.0 0.	170. 0.	170. 0.	180. 0.	180. 0. 200.	200. 0. 200.	200.	130. 0.	130. 0. 0.
Row 28	0. 130. 0.	0. 130. 0.	0. 130. 0.	0. 130. 0.	0. 130. 0.	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	30°.	0. 20. 0.	• • •		40° 0°	70. 0.	120. 0.	0.00	180. 0. 200.	190. 0. 200.	190. 0. 170.	200.	200 . 0.	130. 0. 0.
Row 29	0. 200. 150. 0.	200.	200. 0.	200. 0.	0. 130. 0.	.000	0 000	0°.	• • •		0. 0. 200.	0. 0. 200.	0.000.	40. 0. 200.	90 . 160 . 200 .	190. 160. 150.	190. 150. 0.	190. 140. 0.	200. 140. 0.	200. 140. 0.
Row 30	0. 200. 190. 0.	200. 190. 0.	0. 200. 180. 0.	200. 170. 0.	0. 180. 170.	0. 170. 170.	0. 140. 180.	0. 70. 180.	0. 0. 190.	0. 0. 190.	0. 200. 200.	0. 190. 180.	0. 190. 190.	0. 180. 180.	60. 190. 170.	100. 190. 160.	120. 190. 0.	200. 180. 0.	200 . 180 . 0.	200 . 190 . 0.
Row 31	0. 200. 40.	200. 40.	0. 200. 40.	200. 170. 0.	0. 200. 180.	0. 200. 190.	0. 200. 160.	0. 200. 140.	0. 200. 140.	0. 200. 140.	0. 200. 140.	0. 200. 130.	0. 200. 130.	0. 130. 140.	0. 190. 130.	.0 .0 .0	0. 190. 0.	40. 140. 0.	120. 40. 0.	200. 40. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l column	٤									
	- 12	22	3 23	4 42	5	92	7 72	88	6 8	은 R		12	2 8	-	35	2 %	71	<u> </u>	5 8 8	20 40
Model	41	42 62	43 63	44 64	45	46	47		46									28		09
Row 32	0. 140. 50.	0. 200. 50.	200.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 150. 50.	0° 90°	0. 60. 50.	0° 80°	0. 50. 1	0. 130. 40.	40°.	40.	200° 50°	20. 50. 0.	60. 50. 0.
Row 33	0. 200. 50.	50.	0. 100. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 200. 50.	0. 130. 50.	0. 90. 50.	0. 40. 50.	0. 40. 50.	° ° °	° ° °	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	0. 50.	.050.	. 000	200° 50° 0°	200 . 50 . 0.
Row 34	0000	200.	0°. 0°.		0. 200. 0.	0. 200. 0.	200°.	0.00	70.	50.	50.	50. 0.	0°. 0°.	• • •	• • •	• • •	• • •	· . · ·	• • •	• • •
Row 35	0000	0000	• • • •	0. 0. 0.	0. 200. 0.	0. 200. 0.	0° 200° 0°	0. 0.	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 36	0000	0000	• • • •	100.	0. 200. 0.	0. 200. 0.	0. 200. 0.	0. 100. 0.	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

0	38 39 40	29			0 0 0	• 0	.00.00	0.	•
	36 57			model layer 2	• • • • •	• 0		• • • •	•
1	× × × × × × × × × × × × × × × × × × ×	54		assigned to	0.00.00.00.00.00.00.00.00.00.00.00.00.0	• 0	.0 .0 .0	0 0 0	•
1 12	31 32 33	52		Santa Fe Group	0.000	•0	.00	0 0 0	•
odel colum	29 38 5	20		unit of the S	0 0			0 0 0	
α	27 28	48		hydrostratigraphic	.00			0 0 0	
1	25 26			upper	0 0				
7	7		63 64	Thickness of		.0		.000	
1	21 22		61 62			.00	2 0. 0. 0. 0. 0. 0. 0.	3 0. 0.	
		Model	MOL		Row		Row	Row	

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

	-				ر د			a	Model		E =	1 2	=	5	Į į	4		٥		º
Model	21 41 61	22 42 62	63	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	25 45	26 46	, 27 47	28 48	29 49	2 R G	- 15 15	32 22 22	53 33	<u> </u>	35	2 % 5	37		28 88 28	18 38 39 58 59
Row 6	0000	0000	0000	0000	000	000	000	000		• • •	• • •	• • • •	• • • •	000	000	000	•••			
Row 7		• • • •	0000	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •		• • •	
Row 8		0000	0000	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •		• • •	
Row 9		50.000		0000	• • •	• • •	• • •	• • •		• • •	•••	• • •	• • •	• • •	• • •	• • •	• • •		• • •	
Row 10			0000	0000	•••		•••	• • •	•••	•••			• • •	•••	• • •	• • •	•••		•••	
Row 11	. 0 0 0		0000	.00.00.00	•••	• • •	• • •	• • •	000	• • •	• • •	• • •		• • •	• • •	• • •	• • •		• • •	

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	n co la	E									
	-	2	٣	4	5	9	7	8	6	10	=	12	13						19	20
	21	22	23	24	25	26	27	28	53	ጽ	31	32	33	¥	35	×	37	88	33	40
Mode!	41	42	43	44	45	46	47	48	49	20	51	52	53						29	09
NO.	19	62	63	64																
Row 12	•	•	•	•	•	•	•	°	•	•	•	•	•0	•	•	•	•	•	•	•
	0	•	•	•	•	•	•	•	•0	•	•	.°	•0	•	•	•	•	•	50.	80.
	•06	100	110.	130.	130	130	•	•	•	•	•	•	•	•	•	•0	•	•	•	•0
	•	•0	•	•																
Row 13	•	•	0	•	0	•	•	0	0	0	•	•	•	•	•	•0	•0	•	•	•0
	•	•	•	•	•	•	•	•	•	•	•	•	•0	•	•	•	•	30.	70.	•06
	100	110.	120.	130.	130.	130	130.	•	•	•0	•	•	•0	•	•	•	•	•	•	•
	•	•	•	•																
Row 14	•	•	0	•	0	•	•	•	•	0	•	•	•	•	•	•	•	•	•	•0
	• 0	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•00
	100	100	110.	120.	140.	140.	130	<u>ਲ</u>	140.	•	•0	•	•	•	•	•	•	•	•	•0
	. '	°	°	°																
Row 15	•	•	0	•	•	•0	•	•	•	0	•	•	•0	•	•0	•0	•	•	•	•0
	•	•	•	•	•	•	•	•	•	•0	•	•	•	•	•	•	•	•	•	•0
	100.	100 .	110.	110.	120.	120.	110.	120.	130.	140.	• 0	•0	•	•	•	.	.	•0	• 0	•0
Row 16	.	• 0	.	° 0	.	· •	.	• o	° 0	• o	• •	.	• •	• 0	• •	• •	• o	• •	• •	• •
		• •	10.	10.	120.	120.	100	10.	120.	140.	150.	• •	• •	• •	• •	• •	• •	• •	• •	• •
	•0	•	•	•0																
Row 17	0	•	0	•	0	0	•	•0	•	•	•	•	•	•	•0	•0	•	•0	•	•
	.	.	o 0	o :	• <u> </u>	• c	• <u>•</u>	• <u>•</u>	• <u> </u>	0 0	• 0	8 E	.							
	• •	• •	• •	0.0	•071	•071	•	2	•021	•	•001	•	•	•	•	•	•	•	•	•

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	100	E									
	-	2	3	4	5	و	_	l	6	2	=	12	13	4	15	19	17	82	62	20
	21	22	23	24	25	56	27	28	53	କ୍ଷ	31	32	33	¥	35	8	37	æ	33	40
Model	41	42	63	4 4 64	45	46	47		40	S S	51	52	53	54	55	26	57	28	29	09
Row 18	0	•	0	.	•	•	•	•0		°	•	•	•	•	•	•	•	•		•
	0	0	°	0	0	•	•	· 0	0	•	· •	110.	110.	° 0	•	0	0	•		• 0
	• •	• •	• • •	• •	8	130 •	8	140.		170 .	190•	2 00 .	2 00 .	°	•	°	•	•	•	•
Row 19	0	0	0	°	0	•	°	•	•			•		•	•	•	•	•		•
	•	•	•	•	•	•	•	•	•			110.		130.	•	•	•	•		•
	• •	• •	° c	• •	•	150.	160.	170.	180.	190.	210.	230.	250.	260.	•	•	•	•		•0
	•	•	•	•																
Row 20	•	•	•	•	·	•	°				•	•	•		•	•	•	•		•0
	•	•	•	•	•	•	•				•	110.	120.		•	•	•	•		•
	•	•	•	160.	170.	180.	190.	210.	220.	230.	250.	290	300		260.	•	•	•		•0
	•	•	°	•																
Row 21	•	•	•	°	•	0	• •	•	•			•	•		•			·°		•
	•	•	•	°	•	•	•	•	•			110.	130.			°		190•		210.
	210.	220.	230.	240.	240.	250.	250.	270.	290.	310.	320.	340	350.				•	•		•0
	•	• •	• •	•																
Row 22	0	•	•	0	0	0	•	•	•	•		•	•	•					•	•
	•	•	•	•	•	•	.	•	•	•		120.	140.							250.
	270 . 0.	280 . 0.	280 .	290 .	300.	300.	300.	310.	330.	350.	370.	390.	390.			200.	150.			.
Row 23	0	•	0	0	0	0	•	•	•		•	•	•	0					•	•
	•	•	•	°	•	°	•	°	•	50.	140.	140.	160.	170.	180.	190	200	220.	240.	270.
	290 .	310.	330 . 0.	0 0	¥0.	350.	350.	360.	370.		400	400	400	350.					• •	•06

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	J colu	CE!									
	-	2	3	4	5	9	7	8	6	2	=		13	1	ļ	19	17	18	19	20
	21	22	23	24	25	56	27	28	53	8	31	32	33	¥	35	32	37	83	39	40
Model	41	42	43	44	45	46	47	48	49	20	51		53			26	23	28	29	09
MOJ	61	62	63	64																
					ł						l	l								
Row 24	•	•	•	•						•			•				•	•	•	•
	°	•	•	o	o	•	•	•	.	0.	160	150	170.	190		230.	280.	310.	330	360
	370. 0.	• 0 086	.0° 0°	,00 0,						400•			4 00 •		300		• 001	08	4 0	ċ
Row 25	ô	•	0	0	°			ò	•				o				0			•
	•	0	•	0	0			•	20.				200				240.			240.
	240.	240.	240.	240.	240.	240.	400	400	400	400	400	400	400	390	310.	280.	150.	120.	•	•
	.	•	.	.																
Row 26	•	°	°	•	°			•	°				•					•		•
	•	•	°	•	•			•	40.											220.
	230.	240.	240.	240.	240.	240.	240.	240.	240.	240.	240.	240.		400	350.	300	230.			•
	ô	.	°	°																
Row 27	°	•	•	°				•		•		•	•				°	°	°	•
	°	°	·	•				°												200
	220.	230.	230.	230.	230.	230.	230.	230.	230.		230.			200.	390	•				.
	°	ċ	ċ	ô																
Row 28	°	0	°	°	•			•	°	•			°				°			·°
	°		•	•	•			°	•09	130.			220.				200			150.
	170.	=	180.	180.	180.	180	180	180.	180	180.	180	180.	150.	120.	•	ċ	°			°
	o		o	.																
Row 29	°	°	•	ċ	•	•	•	•	•	0			•		•	•	°	•	•	°
	°	°	ċ	°	°	.	•	•	90	130	200	210.	200.	210.	.	· ·	o	o ·	0	o ·
	.	120.	120	120.	120.	120.	120.	120.	120.	10.			°		ċ	o	°	ċ	°	o
	• •	•	•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet-Continued

									Mode	column	E									
	12	22	3 23	24	5 25	6 26	7 27	88	9 29		31	12	13	⊉ ₩	15 35	2 %	17 37	∞ %	39 29	20
Mode 1 row	14 19	42	63	64 4	45	46	47	48	49	20	51	52	53	54	55	26	23	88	29	09
Row 30	0000	0000	0000	0000	000	000	•••	0 %	0000	000.0	• • •	•••	000	• • •	• • •	• • •	000	000	000	• • • •
Row 31	0000	0000	0000	• • • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 32	0000	0000	0000	• • • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 33	• • • •	• • • •	0000	• • • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
20 X	0000	0000	0000	• • • •	• • •	• • •	• • •		• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 35	0000	0000	• • • •	• • • •	• • •	• • •	· · ·	° ° °	• • •	• • •	• • •	• • •	000	• • •	• • •	· · ·	• • •	• • •	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	ol column	C III									
	-	2	2	4	5	9	7	80	6		=	12	13	14	15	16	1-	<u>®</u>	62	50
	21	22	23	24	25	56	27	28	53	ନ୍ଧ	31	32	33	¥	35	Я	37	88	33	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	57	28	23	09
MOL	19	62	63	64																
Row 5	400	400	400	400	400.	400	400	400	400	400	400		400				400.	400.	400.	400.
	400	4	400	400	400	400.	400	400	400.	400	400	400	400.	400.	400•	•	•	•	•	•
	°	0	•	°	°	•0	•	°	°	•	•		°				•	•	•	•
	o [*]		•	°																
Row 6		400	400	400	400	400	400	400	400	400	400					400.	400.		400	400.
	400	400	400	400	400	400.	400	400.	400.	400.	400	400	400	400.	400	400	•	•	•	•
	•	•	•	•	•	•	•	•	•	°	•					•	•		•	•
	o [*]	•	.	ં																
Row 7		400	400	400	400	400		400	400	400	400	400		400			400.	400		400
	400	400	400	400	400	400	400	400.	400.	400	400.	400	400	400	400	400	400	400	400	•
	0	0	•	•	•	•		•	•	•	•	•		•			•	•		•
	°	•	°	°																
Row 8	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400		400.	400	400	400
	400	400	400	400	400	400	400.	400.	400.	400.	400	400	400.	400	400	400	400	400	400	400.
	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•		•	•	•	•
	o [*]	•	.	.																
Row 9		400	400	400	400	400	400.	400	400	400	400	400	400	400	400.	400	400.	400	400	400.
	400		400	400.	4	400	400	400	400.	400.	400	400	400	400.	400	400	400	400	400	400
	400	350.	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•
	°		•	.																
Row 10		400	400	400	400	400	400	400	400	400	400	400		400.			400.	400	400	400.
	400		400.	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400
	400	400	400	•	•	°	•	•	•	•	•	•		•			•	•	•	•
	°		•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	el column	UMI									
	-	2	3	4	5	9	7	8	6	10	=	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	କ୍ଷ	31	32	33	¥	35	Ж	37	88	ଛ	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	56	57	28	29	09
LOW.	61	62	63	64																
Row 11	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400.	400	400	400.	400
	400	400	400	400•	400•	400	400•	400•	400	400	400	400.	400.	400	400•	400•		400•	400	400
	360.	350.	× 60.	330.	°	•	°	°	°	°	.	°	°	.	.	•	.	•	•	•
	•	•	•	•																
Row 12	400	400	400	400	400	400	400	400	400	400	400	400	400	_	400.	400.	400	400.	400	400.
	400	400	400•	400•	400•	400•	400•	400•	400	400	400	400.	400.	400.	400•	400•	400	400•	350.	320.
	310.	300	290•	270.	270.	270.	•	•	•	•	•	•	•		•	•	•	•	ਂ	•
	°	°	°	•																
Row 13	400	400	400	400	400	400	400	400	400	400	400	400	400	400	400.	400•	400	400	400	400
	400	400	400	400	400	400	400	400	400	400	400		400•	400	400	400	400	370.	330.	310.
	300	290.	280.	270.	270.	270.	270.	•	•	•	•	•	•	•	°	.	•	•	.	•
	°	•	°	°																
Row 14	400	400	400	400	400	400	400	400	400	400	400		400	400.	400.	400.			400	400
	400	400	400	400	400*	400•	400	400•	400.	400.	400.	400.	400•	400.	400•	400.		400.	400	300
	300.	300.	290.	280.	260.	260.	270.	270.	260.	•	°		•		•				•	•
	•	•	•	•																
Row 15		0	•	20.	400•	400	400	400	400	400		400•	400	400•	400•		400•	400	400	400
	400	400	400	400	400	400	400	400	400.	400		400	400.		400	400.		400	400	400.
	300	300	290.	290.	280.	280.	290•	280.	270.	260.	•	•	•		o°			o•	ċ	°
	°	°	°	•																
Row 16	•	°	•	10.	8	130.	400	400	400	400		400		400.	400•		400	400.	400	400.
	400	400	400	400	400	400	400	400	400	400	_	400	400	400	400.	_	400	400	400	400.
	400	, 60 0	290 .	290 . 0	280.	280.	3 00°	290.	2 80 .	. 260 .	250.	•		•	• o	•	•	° 0	• •	• •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	E CEL									
	-	2	2	4	5	9	7	8	6	10	=	12	13	14	15	92	17	2	19	20
	21	22	23	24	25	56	27	28	53	8	31	32	33	¥	35	Ж	37	88	33	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	21	82	29	9
LOW	61	62	63	64			i							ı						
					:									1						
Row 17	•	.	•	°	<u>.</u>	20.	290•	400	400	400	400•	400.					400	400	400	400•
	400	400 .	400 .	400 . 280.	400.	400 .	400°	400 . 290.	400 . 280.	400 . 250.	400 . 240.	370 . 230.	400	400 0	400	400 0	400 0	400 .	400 0	400 0
	°	0	•	•													3	•	:	;
Row 18	•	•	•	•	0	10.	150.	400	400.	400.	400	400.	400.	400			400	400	400	400
	400	400	400	400	400	400	400	400	400	400	400	290.	290•	400	400	400	400•	400	400	400
	400	400.	400 •	400.	270.	270.	270.	260.	250.	230.	210.	200.	200.				.	°.	•	•
Row 19	ċ	ċ	•	•	•	•	<u>.</u>	280.	400	400	400•	400	400			400.	400•	400	400	400
	400	400	400	400	400	400	400	400•	400	400	400	290.			400	400	400.	400.	400	400.
	400	400	400	400	400	250.	240.	230.	220.	210.	190.	170.	150.	140.		•	•	ċ	•	•
	°	°	•	•																
Row 20	°	°	0	•	•	•	•	•09	300	400	400	400					400	400	400	400
	400	400	400	400	400	400	400	400	400	400	400	290.	280.		400	400	400.	400	400	400
	400	400	400.	240.	230.	220.	210.	190•	180.	170.	150.	110.	100	. 06			•	•	0	.
	.	.	•	.																
Row 21	•	•	0	•	•	•	•	•	110.	290.	400	400			400		400	400.	400	400.
	400	400	400	400	400	400	400	400	400	400	350.	290.	270.	260.		400	400•	210.	210.	190.
	190.	180.	170.	160	160	150.	150.	ا گ	110.	•06	80	• 09		40•	160.		.	°	.	•
	°	°	• •	.																
Row 22	•	•	•	•	•	•	•	•	80.	190.	280.	400	400		400	400.	400.	400.	400	400
	400	400	400	400	400	400	400	400	400	400	350.	280.	260.	250.		230	230.	200.	190.	150.
	130	120 . 0	120.	• • •	00	100	100	. 06	. 0	50.	R R	<u>.</u>	. 0	0	120.	. 200	100.	°	°	• 0

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l colu	E									
	-	2	3	4	5	9	7	8	6	9 10 11	Ξ	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	58	53	R 8	<u>ب</u> د	22 23	33	ጃ ፤	ង	አ አ	37	8 8	R (40
Model	61	47 62	63	44 64	დ	40	4	4 8	4 V	20	[2	75	3	2 C	çç	8	/6	8	χ Σ	9
Do:: 03										ļ	l	l	į.	į.		į.				
	400	400	400	400	400	400	400	400.	400	350 .	260.	260.	240.	230.	220.	210.	200.	180.	160.	130.
	110.	06	0.	09	•09	50.	50.	40.									150.	50.	•	•
Row 24	·	°	•	°	•	•	•							310.	400•	400•	400.	400	400.	400
	400. 30.	400. 20. 0.	400. 10. 0.	400.	400.	400.	400.	400.	400 . 0.	330 . 0.	240 . 0.	250.	230.	210 . 30.	190 . 100 .	170 . 140 .	120 . 200 .	90 . 50 .	0.	0.0
Row 25	°	°	°	°		°		°					190.	240.	300	340.	400	400	400	400
	400	400	400	400	400.	400	400•	400	380.	310.	210.	210.	200.	190.	180.	140.	160.	160.	160.	160.
	160 . 0.	160.	160 . 0.	160.		160.		•					•	10.	•06	120.	150.	50•	.	•0
Row 26	•	°	•	•	•	•	•	° 0	•	20.	50.	70.	110.	170.	240.	290.	350.	400	400.	400
	400	400	400	400	400	400	400•	400.	360.	300	200•	.061	180.	160.	160.	160.	160.	160.	170.	180.
	170 . 0.	160 . 0.	160 . 0.	160 .	160.	160.	160.	160.	160.	160.	160.	160.	°	•	50.	100.	170.	•	•	•
Row 27	•	°	o	•	•	•	•	•		0	•	•	80.	110.	170.		260.	320.	400	400
	400	400	400	400.	400	400	400	400	350.	280.	170.	160.	160.	160.	170.	170.	200.	200.	190	200.
	0.	170 . 0.	170 . 0.	170 . 0.	170.	170.	170.	170.		170.	170.	170.	200.	200.	• 0		.	•	•	•
Row 28	°	ô	•	o	°	0	°	•		0	°	•	ċ	•09	. 06		210.	300	400	400.
	400. 150. 0.	400. 140. 0.	400 . 140 . 0.	400. 140. 0.	400.	400.	400 . 140 .	400 . 140.	340 . 170 .	270 . 170 .	160 . 170 .	170 . 170.	180 . 210.	170 . 280.	190 . 210 .	200 . 260 .	200.	170 . 0.	160.	170.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l colle	Ē									
	-	2	3	4	5	9	7	}	6	10 11	=	12	13	14	15	16	11	18	19	20
	21	22	23	24	25	56	27	28	59	8	31	32	33	¥	35	Я	37	æ	33	40
Model	41	42	43	44	45	46	47		49	20	51	52	53	54	55	56	22	28	26	9
NO.	61	62	63	64																
Row 29	•	•	•	•	•	•	•	•	•	•	•0	•	•	•	°	110.	200.	150.	200	400
	400	400	400	400	400	400	400	390	340.	270.	200	190.	200°	190.	180.	130.	80.	80.	80.	70•
	80°	0 0	°°°	°°°°	8 0	80.	80	80•	80•	70 •	70•	80	100.	° 06	•09	.	.	.	°	•
Row 30	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	•	0	8	09	200.
	400	400	400	400	400	400	400	370.	340	300	400	300	200.	150.	130.	40.	40.	40.	40.	40.
	40.	40.	40.	.0	40•	40•	40.	40.	40•	40•	40.	•	•	°	•	.	•	0	0	•0
Row 31	•	0	0	•	°	•	•	0	•	0	°	•	•	•	•	•	•	0	•	70.
	210.	300	300	310.	310.	390	350.	340.	340	320.	150.	80.	•06	•	•0	•0	•	•	•	•
	•	•	•	•	•	•	•	°	•	•	•	•	•	•	•0	•	•	•	•	•
	° 0	°	°	°																
Row 32	0	°	0	•	0	•	• 0	0	•	•	0	0	•	•	•0	•	•	• 0	•	•
	•	•	200•	200	200	200•	200	200	200	•0	•	•	•	•	•	•	•	•	•	•
	• •	· ·	• •	• •	•	°	.	.	•	°	•	•	.	.	•	.
Row 33	°	°	°	•	ċ	•	°	•	•	•	<i>.</i>	.	•	.	•	° 0	.	0	0	.
	°	•	·°	180.	180.	180.	180.	180.	•	•	°	•	•	•	•	•	•	•	•	•
	• •	• •	• •	• •	•	°	°	°	°	°	•	•	• •	°	•	•	°	•	•	°
Row 34	0	0	Ó	0	Ó	ď	ď	ď	ď	ď		Ö	ď	Ó	ď	ď	Ö	ď	Ó	Ö
	•				180.	180.	°		•				•			•		0	0	
	• •	• •	•	°	0	°	°	°		•	°	°	•	•	°	°	0	.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l column	٤									
	- ;	2 2	3	4 5	5 4	9 %	1- 2	8 8	6 6	01 05	= ;	12	52 2	4 2	15	92 %	12 13	<u>8</u> 28	5 5	200
Model	61	42 62	43 63	4 4 4 4 4 4 4 4	45	76 46	47	8 78	4 49	S 55	2 12	52 52	53	* *	55	2	57	28 28	59	09
Row 35	0000	0000	0000	0000	0. 180. 0.	0. 180. 0.	000	000	000	000	• • • •	000	•••	• • •	000	000	000	000	000	• • • •
Row 36	0000	0000	0000		.0 .0 .0	0. 180. 0.	000	• 0	000	000	0 0 0	000	000	• • • • • • • • • • • • • • • • • • • •	000	000	000	• • •	• • •	• • •
Row 1	0000		0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0			0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0		. 0 . 0						0 0		0.0.00.0		000	000	000
Row 2	0000	0000	0000	0000	• • •	000	• • •	000	000	000		• • •			• • •	• • •	• • •	• • •	000	• • •
Row 3	0000	0000			• • •	• • •	• • •	000	000	000	000	000	000	000		• • •	000	000	• • •	000

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

ı	i	i					
	29 9 9	000	000	000	000	000	• • •
	23 23 24	000	• • •	• • •	• • •	• • •	• • •
	18 38 58		• • •	• • •	• • •	• • •	• • •
	17 37 57	•••	• • •	• • •	• • •	• • •	• • •
	5 % % %	•••	• • •	• • •	• • •	• • •	• • •
	15 35 55	•••	• • •	• • •	• • •	• • •	• • •
	7 % %	• • •	• • •	• • •	• • •	•••	• • •
	13 33 53	000	• • •	• • •	• • •	• • •	· · ·
	12 32 52	•••	• • •	• • •	• • •	· · ·	· · ·
٤	31	•••	• • •	• • •	• • •	• • •	• • •
Model column	5 5 5 5 7 8	• • • •	• • •	• • •	• • •	• • •	· · ·
Mode	6 6 6 7 8 9	000	• • •	• • •	• • •	• • •	• • •
	8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	•••	• • •	• • •	• • •	• • •	• • •
	27	• • • •	• • •	• • •	• • •	• • •	• • •
	5 5 4 6	000	• • •	• • •	• • •	• • •	• • •
	5 25 45	000	• • •	• • •	• • •	• • •	• • •
	4 4 4 4 9	0000	• • • •	• • • •	• • • •	• • • •	0000
	23 43 63	0000	• • • •	0000	• • • •	• • • •	0000
	22 42 62	0000	• • • •	• • • •	• • • •	• • • •	0000
	121	0000	0000		• • • •	• • • •	0000
	= -	4	'n	9	7	ω	0,
	Model	R _O ₩	Row	Row	Row	Row	Row

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	l column	٤										
	1 21	22	3	24	5 25	92	7 27	8 28	9	5 5	31	12	13	<u> </u>	15	<u>م</u> م	17	38	19 39	20 40	
Model	61	42	63	4 4 6 4	45	46	47	48	6	20	51	52	53	54	55	26	57	58	59	09	
Row 10	0000	0000	0000	0000	• • •	000	• • •	000	• • •	• • •	• • •	• 0	000	• • •	• • •	• • • • • • • • • • • • • • • • • • • •	• • •	• • • •	• • •	• • •	
Row 11	0000		0000	0000	• • •	• • •	· · ·	• • •	· · ·	• • •		· · ·	• • •	• • •	• • •	• • •		• • •	• • •	• • •	
Row 12	0000	• • • •		0000		• • •	• • •			• • •			• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	
Row 13	0000	• • • •				• • •	• • •	• • •	• • •	· · ·	• • •	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	
Row 14	0000	• • • •			• • •	• • •	• • •	• • •	• • •	• • •		• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	
Row 15	0000	• • • •	0000	0000	000	• • •	• • •	• • •	· · ·		• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

-	2	3	4	70	٥	7	8	Mode!		LI T	12	13	4-	15	16	12	8	62	70
21 41 61	22 42 62	23 43 63	24 44 64	25 45	26 46	27	28 48	29	30	31	32 52	33 53	¥ ¥	35 55	36 56	37 57	28 28	39 59	4 0
0000	0000	0000	0000	000	000	• • • •	000	• • •	• • •	000	• • • •		• • • •	•••	•••	000	•••		
0000	• • • •	• • • •	0000	000	000	000	000	• • •	000	•••	• • •	• • •	• • •	• • •	• • •	· · ·	•••	• • •	•••
0000	• • • •	0000	0000	· · ·	000	20°.	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •
0000	• • • • • • • • • • • • • • • • • • • •		0000	• • •	000	0.00	120. 0.	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •
0000	0000		0000	• • •	000	• • •	0.00	00.	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
0000	0000		0000	000	000		20.	230. 0.	0.00	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	•••

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	100	CIMIS									
	- 5	2 2	3 %	4 2	5 5	9 %	1- 2	8 8	6 %	01	= =	12	55 55	4 2	15	25 %	71	8 8 8	62 02	8 8
Model	61	45 62	63	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	4 5	46	47		49	S 20	5 15	52	55	7 7	55	3 %	52	28 8	20 6	9 09
Row 22	0000	0000	0000	0000	000	•••	•••	000	00.	210. 0.	120. 0.	000	000	000	000	000	000	000	000	•••
Row 23	0000	0000		0000	• • •	• • •	• • •	000	0. 0.	200.	220. 0. 0.	0.0.0.	80°. 0°.	• • •	000	• • •	000	· • •	• • •	• • •
Row 24	0000	0000	• • • •	• • • •	• • •	• • •	• • •	• • •	• • •	00.	300. 0. 0.	220. 0.	130. 0.	0000	• • •	• • •	• • •	• • •	• • •	• • •
Row 25	0000	0000	• • • •	• • • • •	• • •	• • •	• • •	• • •	• • •	00.00.0	250. 0.	300. 0. 0.	210. 0.	160. 0.	00.	0000	• • •	· · ·	• • •	• • •
Row 26	0000	0000	0000	• • • •	• • •	· · ·	• • •	000	• • •	50.	00.	330. 0. 0.	290 . 0.	230. 0. 0.	160. 0.	0.0	50.	• • •	• • •	• • •
Row 27	0000	0000	0000	0000	• • •	000	• • •	• • •	• • •	000	• • •	• • •	320. 0.	290. 0.	230. 0. 0.	0.00	140. 0.	0000	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	dami 100	1									
Model	- 12 14 10	22 42 62	23 43 63	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	5 25 45	6 26 46	27	8 28 48	9 29 49	.1	31	12 32 52	13	14 34 54	15 35 55	16 36 56	17 37 57	18 38 58	19 39 59	20 40 60
Row 28	0000	0000	0000	0000	000		• • •	000	• • •		000	000	• • •	40. 0.	09	00.	.00 .0	.001 .0	000	• • •
Row 29	0000	• • • •	0000	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	0 0 0	20. 0.	°0 °0	100. 0.	50.	• • •
Row 30	0000	• • • •	0000	0000	· · ·	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	0000	0 0 0	20.
Row 31	0000	• • • •	0000	• • • •	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 32	• • • •		0000	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •
Row 33	0000		0000	0000	• • •	• • •	000	• • •	· · ·	• • •	• • •	• • •	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	el column	cwn									
	-	2	3	4	5	9	7		6	01	Ξ	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	30	31	32	33	¥	35	36	37	38	39	40
Model	41	4 2 62	43 63	4 4 64	45	46	47	i	46	50	51	52	53	54	55	56	57	58	59	09
М	260. 190. 0.	260. 0. 0.	260. 0. 0.	70. 0. 0.	70. 0. 0.	70. 0. 0.	70. 0. 0.	150. 0.	150. 0.	150. 0. 0.	150. 0. 0.	150. 0. 0.	150. 0. 0.	150. 0. 0.	160. 0.	160. 0.	170. 0. 0.	170. 0. 0.	180. 0. 0.	180. 0. 0.
4	260. 180. 0.	260. 190. 0.	260. 200. 0.	260. 200. 0.	260 . 200 . 0.	70. 210. 0.	70. 220. 0.	70. 230. 0.	70. 110. 0.	70. 110. 0.	150. 110. 0.	150. 110. 0.	150. 0. 0.	150. 0. 0.	150. 0. 0.	150. 0. 0.	160. 0. 0.	170. 0. 0.	170. 0. 0.	180. 0. 0.
r.	260. 170. 0.	260. 180. 0.	260. 200. 0.	260. 200. 0.	260. 200. 0.	260. 200. 0.	90. 200. 0.	90. 200. 0.	80. 200. 0.	70. 230. 0.	70 . 230 . 0.	120 . 110 . 0.	150. 110. 0.	150. 110. 0.	150. 110. 0.	150. 0. 0.	160. 0. 0.	160. 0.	160. 0. 0.	170. 0. 0.
9	260. 160. 0.	260. 160. 0.	260. 160. 0.	260. 160. 0.	260 . 160 . 0.	260 . 160 . 0.	260 . 180 . 0.	100. 190. 0.	100. 220. 0.	100. 340. 0.	110. 380. 0.	130 . 400 . 0.	140. 420. 0.	140. 420. 0.	140. 460. 0.	150. 110. 0.	160. 0.	160. 0.	160. 0.	160. 0.
7	260. 140. 0.	260. 140. 0.	260. 160. 0.	260. 90. 0.	260. 90. 0.	260 . 110 . 0.	260. 200. 0.	260 . 390 . 0.	260. 420. 0.	260. 420. 0.	250. 420. 0.	250. 460. 0.	250 . 460 . 0.	250. 460. 0.	250. 450. 0.	160. 450. 0.	170. 450. 0.	170. 450. 0.	200. 470. 0.	130. 0.
80	260. 260. 0.	260. 130. 0.	260. 120. 0.	260. 130. 0.	260 . 160 . 0.	260. 230. 0.	260 . 280 . 0.	260 . 300 . 0.	260 . 340 . 0	260 . 420 . 0.	250. 460. 0.	250. 470. 0.	250 . 460 . 0.	250. 450. 0.	240. 450. 0.	240. 450. 0.	240 . 460 . 0.	240. 470. 0.	250. 480. 0.	260 . 480 . 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	o le										
	-	2	3	4	5	9	7	®	6	_	=	12	13	4	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	30	31	32	33	ጃ	35	%	37	38	39	04
Model	4	42	43	4	45	46	47	4 8	49	20	51	52	53	<u>7</u> 4	22	26	27	28	29	09
MOL	61	62	63	64																
Row 9	260. 260. 500. 0.	260. 330. 500. 0.	260. 340. 0.	260. 150. 0. 0.	260 . 180 . 0.	260. 240. 0.	260. 250. 0.	260. 330. 0.	260. 350. 0.	260 . 380 . 0.	250. 470. 0.	250. 460. 0.	240. 460. 0.	240. 450. 0.	240. 450. 0.	220. 460. 0.	240. 470. 0.	250. 480. 0.	250. 480. 0.	250. 490. 0.
Row 10	260. 270. 500. 0.	260. 320. 500.	260° 330° 530° 0	260. 350. 0.	260 . 360 . 0.	260 . 260 . 0.	200 . 290. 0.	220. 340. 0.	250. 370. 0.	260 . 380. 0.	260 . 390. 0.	270. 450. 0.	280. 450. 0.	290. 450. 0.	300 . 460 . 0.	220. 470. 0.	230. 480. 0.	240. 480. 0.	240. 490. 0.	250 . 500 . 0.
Ro 1.1 1.1	260. 600. 510. 0.	260. 600. 520. 0.	260. 600. 520. 0.	260. 600. 540.	210. 380. 0.	210. 400. 0.	210. 410. 0.	280. 430. 0.	300. 360. 0.	340. 370. 0.	360 . 370 . 0.	260. 370. 0.	270. 450. 0.	280. 450. 0.	290. 470. 0.	300. 480. 0.	320. 480. 0.	390. 490. 0.	460. 500. 0.	600. 500. 0.
Row 12	250. 600. 510. 0.	250. 600. 520. 0.	250. 600. 530. 0.	250. 600. 570. 0.	230 . 600 . 580 .	230 . 600 . 600.	220. 430. 0.	260. 450. 0.	310. 470. 0.	360 . 490. 0.	380 . 370. 0.	400. 370. 0.	440. 370. 0.	480 . 370. 0.	530. 480. 0.	380. 480. 0.	410. 490. 0.	480. 500. 0.	510. 500. 0.	600. 500. 0.
Row 13	250. 600. 520. 0.	250. 600. 530. 0.	250. 600. 550. 0.	250 . 600 . 570 . 0.	250 . 600 . 600.	250 . 600 . 600 .	270 . 600. 600.	300. 600. 0.	330. 490. 0.	360. 500. 0.	390. 510. 0.	430. 360. 0.	470. 360. 0.	490. 360. 0.	550. 370. 0.	580. 370. 0.	600. 380. 0.	600. 490. 0.	600. 500. 0.	600. 500. 0.
Row 14	250. 600. 530. 0.	250. 600. 550. 0.	250. 600. 570. 0.	250. 600. 590. 0.	250. 600. 600.	250 . 600 . 600.	280 . 600. 600.	320 . 600 . 600.	340. 600. 600.	370 . 520 . 0.	420. 530. 0.	470. 550. 0.	500. 350. 0.	520 . 360. 0.	570. 360. 0.	590. 370. 0.	600. 380. 0.	600. 380. 0.	600. 390. 0.	600 . 520. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	col	umn									
	-	2	3	4	5	9	7	8	6	10	11	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	53	30	31	32	33	34	35	36	37	38	39	40
Model	41	42	43	44	45	46	47	48	49	22	51	52	53	%	55	26	27	28	23	09
MOL	61	62	63	29																
Row 15	0	0	6	0	ď	250.	300	340-	370-	390.	440.	480-	510.	530.	580.	600	600	600	600	600
.	600	•009	. 009	•009	•009	. 009	. 009	•009	009	.009	550.	570.	590.	350.	360.	360.	370.	380.	380.	400.
	550.	570 . 0.	.065 0.	•009 0•	•009	•009	•009	•009	•009	•009	°	.	o	•	.
Row 16	°	°	°	ò	•	•	320.	360.	380.	400	470.	500	520.	550.	590.	•009	•009	•009	•009	•009
	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	2 90°	•009	•009	2 90°	360.	370.	370.	380.	390.
	400 0	400 0	•009 0•	•009 •	•009	•009	•009	•009	•009	•009	•009	°	.	•	•	• 0	•	• 0	•	°
Row 17	°	°	0	•	0	•	0	290.	310.	430.	490.	520.	540.	570.	•009	•009	•009	•009	•009	•009
	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	360	370.	370.	380.
	390.	390.	400	•009	•009	•009	•009	•009	•009	•009	•009	•009	.	.	•	.	ċ	°	.	°
	;	•	•	•																
Row 18	°	.	•	°	•	•	•	70•	70•	450.	500	530.	550.	570.	•009	•009	•009	•009	•009	•009
	•009 •	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	360.	360.	370.
	380°	380°	390	4 00 •	•009	•009	•009	•009	•009	•009	•009	•009	520.	°	• •	• •	• •	• o	o	•
Row 19	°	0	•	°	•	•	.	•	70•	80.	100	570.	2 80°	•009	•009	•009	•009	•009	•009	•009
	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	360.
	370.	370.	380	390.	400	•009	•009	•009	•009	•009	•009	•009	•009	440.	°	o [*]	.	o	°	°
	.	•	.	•																
Row 20	•	0	.	°	.	•	•	•	°	80°	80°	150.	200•	•009	•009	•009	•009	•009	•009	•009
	•009	•009	•009 600	009	•009 •	•009 •	009	•009 •	•009 •	•009 •	•009 •	•009 •	•009 •	600	600	•009	•009 •	•009	•009	•009
	000	000	000	000	•000	•000	•000	•000	•000	•000	•000	•000	•000	•000	000	•	•	•	•	•

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

-	,				۲	۰	7	a	yow o	Model column	umo 1-1	5	<u> </u> <u> </u> <u>-</u>		7	4		ğ	2	
23 24 25	23 24 25	24 25	, ₂		, 8		27	78 78	, 62	2 8	2 5	32	. E	34	35.5	% <u>5</u>	37	2 82	39	5 04
42 43 44 62 63 64	43 44 45 63 64	44 45	45		4		47	48	49	20	51	52	53	54	55	56	57	28	29	09
0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	0°00°00°00°00°00°00°00°00°00°00°00°00°0	°009		009 600	1	0°• 600°•	009 600.	0° 600° 600°	009 600	80° 600° 600°	80. 600.	80. 600. 600.	80. 600. 550.	350. 600. 440.	.009 .000	600°. 000°.	600°. 000°.	.009 .000 0.	600. 000.
0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	009 •000		009 0009		009 •000	009 •000•	009 •000	009 600	009 600	80. 600. 600.	80. 600. 600.	70 . 600 . 600.	80. 600. 500.	370 . 600 . 360 .	400. 600. 0.	•009 •009	•009 •009	•009 •009
0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	°009		009 •000		009°	009 •000	009 •000	009 600	009 •000	009 •000	009 •000	50°. 600°. 600°.	70. 600. 500.	70 . 600 . 440 .	70 . 600 . 0	460. 600. 0.	•009 •009 •	.009 .000 0.
0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0 0 0 0 0 600, 600, 600, 600, 600, 600,	0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0° 0	°009		009°		°009	009 •000	009 •000	009 •000	°009	009 •000	009 •000	0° 600° 550°	20 . 600 . 440 .	60 °. 600 °. 220 °.	40 . 600 . 0	290 . 600 . 0.	200. 600. 0.	.009 .000
0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 600. 600. 470. 460.	0. 600. 460.		0. 600. 460.		009 600	009 600	009 600	.009 600.	009°	0°• 000°•	0°• 000°•	0° 600° 500°	0. 600. 400.	0. 600. 180.	40. 430. 0.	150. 440. 0.	150. 450. 0.	140. 460. 0.
0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	000 000 000 000 000 000 000 000	000 000	000	_	°000		°000 °000	°000 °000	°000	°000	°000	°000	°000	°000 °000	°000 °000	0. 400.	0. 410. 0.	30 . 350 . 0.	30. 350. 0.	30. 000. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

	-	2	3	4	5	9	1 1	-	11	8	Mode 8 9	Model colum 8 9 10	Model column 8 9 10 11	Model column 8 9 10 11 12	Model column 8 9 10 11 12 13	Model column 8 9 10 11 12 13 14	Model column 8 9 10 11 12 13 14 15	Model column 8 9 10 11 12 13 14 15 16	Model column 8 9 10 11 12 13 14 15 16 17	Model column 8 9 10 11 12 13 14 15 16 17 18
Model	21 4 4 1 6 1 9 1 9 1 9 1	22 42 62	23 43 63	24 44 64	25 45	26 46	27	28	29		20 30	30 31 50 51	20	30 31 50 51	30 31 32 50 51 52	30 31 32 33 50 51 52 53	30 31 32 33 34 50 51 52 53 54	30 31 32 33 34 35 50 51 52 53 54 55	30 31 32 33 34 35 36 50 51 52 53 54 55 56	30 31 32 33 34 35 36 37 50 51 52 53 54 55 56 57
30w 27	0.000 0000	000 0000	000 0000	0000	0° 0000 0000	.000	000°	0°. 600.	°000 °000	000°	1	0000	1	0000	0° 0° 600° 600°	0° 0° 0° 000 0000 000° 0	0° 0° 0° 0° 000 000° 000° 000° 000° 00	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
Row 28	0 000	0000	0009	.009 .000	°0 °0 °0	°009	.009 0.00	0° 000°	0° 600°	0° 600° 0°	<u>.</u>	0000	w w	• 0 0 0 0	0. 0. 600. 2	0. 0. 0. 600. 600. 230. 2 0. 0. 0.	0. 0. 0. 0. 600. 600. 230. 250. 1 0. 0. 0. 0.	0. 0. 0. 0. 0. 600. 600. 230. 250. 160. 1 0. 0. 0. 0. 0.	0. 0. 0. 0. 0. 0. 0. 0. 600. 600. 230. 250. 160. 150. 0. 0. 0. 0. 0. 0.	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
Row 29	0000	0 0 0	.009 .000	000	009	°009	°009	°009	°009	• 009		0. 200. 0.	(4	2000.	0. 0. 200. 200. 1 0. 0.	0. 0. 0. 200. 200. 110. 1	0. 0. 0. 0. 200. 200. 110. 0. 0. 0. 0. 0. 0. 0. 0.	0. 0. 0. 0. 0. 0. 0. 200. 200. 110. 110.	0. 0. 0. 0. 0. 200. 200. 110. 110. 0. 0. 0. 0. 0. 0. 0. 0. 0.	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 200. 200.
Row 30	0000	0000	0 0 0 0	000.00	.001 .00.	120.	0. 120. 0.	0. 120. 0.	0. 150. 0.	0. 150. 0.		•••	• • • •		• • •	• • • •		.00.00.00.00.00.00.00.00.00.00.00.00.00	0.	0.
Row 31	0000	0000	0000	0000	000		• • •	• • •	• • •	• • •		• • •	• • • •		000			.0 .0 .0 .0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
Row 32	0000	0000	0000	0000	000	000	•••	000	• • •	• • •		• • •	• • • • • • • • • • • • • • • • • • • •		• • •	• • • •		0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									P C C	amir 100 1										
	- =	2 2	3.3	4 %	5 25	9 %	7	8 %	6 6	-1 - "	1 - 5	12	13	4 4	15	2 %	17 %	8 %	6 %	20
Model	4 1 6 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	4 2 62	43 63	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	7 4	46	47	8 4	49	20 2	15	52	23 7	54	55	26 2	57	₹ %	59	Ç 09
Row 33	0000	0000	0000	0000	000	• • •	•••	000	000	• • •	000	000	000	000		000	000	000	000	• • • •
Row 34	• • • •	0000	• • • •		• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •		• • •	• • •	• • •
Row 35	0000	0000	• • • •		• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 36	• • • •	0000	0000	• • • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·
		_	Thickness of	iss of	lower	hydros	hydrostratigraphic	raphic	t in	of the	Santa	Fe Gro	Group assi	glgned	to mod	model layer	/er 3			
Row 1	0000	0000	• • • •		· · ·	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	· · ·	•••	· · ·	· · ·	· · ·	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	100 16	Ē									
	1_	2	3	4	5	9	7	8	6		=	12	13	14	15	92	17	182	19	20
	21		23	24	25	56	27	28	53	30	31	32	33	34	35	36	37	38	39	40
Model	4	42	43	4 ;	45	46	41	48	46	20	51	52	53	3 2	55	26	27	28	59	09
Mo.	ō		69	2																
Row 2			•	.	•	•	.	.	•009	•009								•009	009	•009
	909		°	°	•	•	•	•	•	•	•		•			•		•	•	•
	ં (• •	.	.	•	•	.	•	°	°		•		•	.		•	•	•	.
	٠		•	.																
Row	3 340			530.	530.	530.	530.	450.	450.	450.	450.	450.		450.				430.	420.	420.
	410			•009	•009	•009	•009	•009	•	°	•	•	•	•	•	•	•	•	•	•
		• •	• •		•	•	.	°	°°	.	.	•		°				•	°	•
100	245		240	9	072	630	23	630	620									02.4	024	0
	•		2 6			200	200	. 000	,									2 0	· ·	• 0
	4		9	9	3 9	• 0 • 0	• 000	• •	• •	• • •	5 c	• •	• c	• c	. c	• c	• •	. c	• •	• c
				; °	•	•	•	•	•									•	•	•
Row	5 340		340.	340	340	340.		510.	520.	530.								440.	440.	430.
	4 20		400	400	400	400	400	400	400	370.	370.	400		400•	400	•		°	°	•
		• •	°	•		•	•	•			ċ				°	.	•	•
Row	6 340.	340.	340.	340.	340.	340.	340.	500.	500.	500.	490.	470.	460.	460.	460.	450.	440.	440.	440.	400
	, G				•	0	0		• •	•	0		•							
	5		°	•																
Row	7 340		340.	340.	340.	340.	340.	340.	340.	340.	350.	350.	350.	350.	350.	440.	430.	430.	400	400.
	400		400	400	400	410.	400	210.	180	180.	180.	140.	140.	140.	150.	150.	150.	150.	130.	•
	ં (• •	ં ત	ં લ	°	ċ	•	o°	°	°	°	°	°	°	•	•	°	°	°	•
	٠		•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	C WIT									
	_	2	m	4	Ŋ	9	7	ω	σ	0	=	12	13	4	12	9	17	<u>8</u>	16	20
	21	55	23	24	25	26	27	28	62 9	30 00 00 00 00 00 00 00 00 00 00 00 00 0	<u> </u>	32	33	¥ 2	35 55	36	37	38	39	04 0
MOL MOL	19	62	63	49	-	9	÷	ç	ĵ.	3	5	77	3	ţ	3	3	5	۹ ا	5	8
ω	340. 340. 0.	340. 370. 0.	340. 400. 0.	340. 400. 0.	340. 400. 0.	340. 370. 0.	340. 320. 0.	340. 300. 0.	340. 260. 0.	340. 180. 0.	350. 140. 0.	350. 130. 0.	350. 140. 0.	350. 150. 0.	360 . 150 . 0.	360 . 150 . 0.	360 . 140 . 0.	360 . 130 . 0.	350. 120. 0.	340. 120. 0.
0	340. 340. 100.	340. 270. 100.	340. 260. 0.	340. 400. 0.	340. 400. 0.	340. 360. 0.	340. 350. 0.	340. 270. 0.	340. 250. 0.	340. 220. 0.	350. 130. 0.	350. 140. 0.	360. 140. 0.	360. 150. 0.	360. 150. 0.	380 . 140. 0.	360. 130. 0.	350. 120. 0.	350. 120. 0.	350. 110. 0.
2	340. 330. 100. 0.	340. 280. 100.	340. 270. 70. 0.	340. 250. 0.	340. 240. 0.	340. 340. 0.	140. 310. 0.	300 . 260 . 0.	300 . 230 . 0.	340. 220. 0.	340. 210. 0.	330. 150. 0.	320. 150. 0.	310. 150. 0.	300. 140. 0.	380. 130. 0.	370. 120. 0.	360. 120. 0.	360. 110. 0.	350. 100. 0.
=	340. 0. 90.	340. 0. 80.	340. 0. 0.	340. 0. 0.	140. 220. 0.	140. 200. 0.	140. 190. 0.	320 . 170 . 0.	300. 240. 0.	260. 230. 0.	240 . 230 . 0.	340. 230. 0.	330. 150. 0.	320. 150. 0.	310. 130. 0.	300 . 120 . 0.	280. 120. 0.	210. 110. 0.	140.	0. 00.
12	150. 0. 90.	150. 0. 80.	150. 0. 70.	150. 0. 30.	150. 0. 20.	150. 0. 0.	150. 170. 0.	340. 150. 0.	290. 130. 0.	240. 110. 0.	220. 230. 0.	200. 230. 0.	160. 230. 0.	120. 230. 0.	70. 120. 0.	220. 120. 0.	190. 110. 0.	120. 100. 0.	90. 100. 0.	.001
Row 13	150. 0. 80.	150. 0. 70.	150. 0. 50.	150. 0. 30.	150. 0.	150. 0.	260. 0. 0.	300. 0. 0.	270. 110. 0.	240 . 100. 0.	210. 90. 0.	170. 240. 0.	130. 240. 0.	110. 240. 0.	50. 230. 0.	20. 230. 0.	0. 220. 0.	. 0 10 0	.000.	100.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	n co la	CME									
	-	2	٣	4	5	٥	-	8	6	2	=	ı	l	l	l	ı	17	18	19	20
	21	22	23	24	25	56	23	28	53	30	31	32	33	34	35	36	37	88	39	40
Model	61	42	43 63	44 64	45	46	47	48	49	20	51						57	28	29	09
Row 14	150.	150.	150.	150.	150.	150.	190			230.							•	°	•	•
	•	°	•	•	•	•°	°	•	•	80.	70.	50.	250.	240.	240.	230.	220.	220.	210.	80•
	70.	20.	30	10.	.	.	•			.							°	.	.	°
	•	.	°	°																
Row 15	•	°	•	°	0	150.	200			210.	160	120.						•	•	•0
	•	°	°	•	•	°	°	•	°	•	50.	30.	10.	250.	240	240.	230.	220.	220.	200.
	50.	30.	10.	0	°	°	°			°	•	°						•	o	•
	°	o	o	o																
Row 16	0	•	•	°	0	0	280.	240.	220.	200		100	80.						°	•
	•	•	•	•	•	•	•	•	•	•	•	.01	•	•	0.	240.	230.	230.	220.	210.
	200	200	•	•	°	°	°	•	°	•		•	°						°	°
	ċ	.	o°	.																
Row 17	°	°	.	·°	0	0	°	310.	290•	170.	110.	80.	•09				•	•	•	•
	°	°	.	°	°	.	.	.	o	o	°	•	o	.	• •	•	240.	230.	230.	220.
	210.	210.	200.	.	•	°	.	°	.	°	°	•	°				•	•	.	•
	• •	• •	.	.																
Row 18	°	•	0	•	°	°	•	100	260.	150.	100	70.	50.		•	0	•	0	°	•0
	°	•	°	°	.	•	°	•	.	.	.	°	.	.	•	°	•	240.	240.	230.
	220.	220.	210.	200•	.	.	°	°	°	°	•	.	80		•	o	•	•	•	•
	o	°	°	o																
Row 19	o	•	•	•	•	•	0	50.	280.	300	500.	30.	0.	•	•	°	•	•	•	•0
	°	.	o	•	•	•	°	•	•	°	°	•	•	°	•	°	o	o	.	240.
	230 .	230.	220 . 0 .	210 .	200•	o	o	.	o [*]	.	.	•	•	160.	o	.	•	•	o	.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									1	30										
	- 5	2 2	۳ ٪	4 4	د در	9 %	ر د د	8 8	6 00	10 05	= =	12	55	4 2	5 5	2 %	77 2	85 85	65	20
Model	41	42 62	63 43	64 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	45	46	47		64	2 2	2 72	52	53	7. 7.	52	2, 25	52	28 28) 6 <u>5</u>	2
Row 20	0000	0000	0000	0000	000	• • •	• • •	000	100. 0.	310.	360.	450. 0.	400. 0.	00.	0. 0. 240.	000	000	000	000	• • •
Row 21			0000	0000	000		• • •		• • •	200.	380. 0.	440. 0.	510. 0.	520. 0. 50.	250. 0. 160.	0. 0. 150.	• • •	000		
Row 22			0000	0000	•••		• • •	• • •	• • •	50.	200.	460. 0.	490. 0.	530. 0. 0.	520. 0. 100.	230. 0. 240.	200.	• • •	• • •	
Row 23			0000	0000			• • •	• • •	• • •	• • •	00.	300. 0. 0.	420. 0.	530. 0. 0.	530. 0. 100.	530. 0. 160.	530. 0. 0.	0.0	• • •	• • •
Row 24	• • • •		0000	0000	• • •	•••	• • •	• • •	• • •	• • •	0.00	230. 0. 0.	360 . 0.	430. 0. 50.	460. 0. 160.	530. 0. 380.	560. 0. 0.	310. 0.	400. 0.	
Row 25	0.0120.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0	0000	0. 120. 0.	0.130.00	0.0140.	0.0140.		• • •	• • •	•••	• • •	150.	230. 0.	230. 0. 100.	250. 0. 200.	280. 0.	560. 170. 0.	450. 160. 0.	450. 150. 0.	460. 140. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	100 [CEE									
	-	2	8	4	5	9	7		6	10	=	12	13	4	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	59	30 31	31	32	33	አ	35	፠	37	38	33	40
Mode!	61	4 2 62	63	44 44	45	46	47		49	S.	51	52	53	₹	55	26	22	28	23	09
30w 26	0000	0000	0000	0000	0000	000	000	0000	000	0000	000	0000	150. 0.	160. 0.	180. 0. 0000.	200 . 200 . 000.	510 . 190 . 000.	570. 250. 0.	570. 250. 0.	570. 0000. 0.
Row 27	0000	.000	0000	000	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	0000	• • • • • • • • • • • • • • • • • • • •	• • • • • • • • • • • • • • • • • • • •	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	000	• • • • • • • • • • • • • • • • • • • •	• • • • • • • • • • • • • • • • • • • •	°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°°	0. 310. 000.	70. 400. 0.	210. 500. 0.	400. 000.	400. 000.	570. 0000. 0.
Row 28	0. 570. 0.	0000		0000	000	000	000	• • •	000	000	000	000	0. 370. 0.	0. 350. 0.	0. 0.	0. 450. 0.	• • • • • • • • • • • • • • • • • • • •	150. 0.	200.	400. 0.
Row 29	400°.	0. 560. 0.	• • • •	0000	• • •	• • •	• • •	• • •	000	• • •	0. 400.	0. 400.	0. 470. 0.	0. 490. 0.	• • •	• • •	• • •	• • •	• • •	200.
Row 30	200. 0.	.004	0. 520. 0.	500°.	0. 500. 0.	0. 480. 0.	480. 0.	0. 480. 0.	0. 450. 0.	0. 450. 0.	000	000	000	• • •	000	000	000	000	000	
Row 31	0000	• • • •	• • • •	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	•••

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Model	ဝ၁	umn									
	- 12	2	3	4 4	5	6 26	7 72	8 88	9		11	12	13 33	7 42	15	5 36 36	17	18 38	59 39	20
Model	14 9	42	63	4 4 4 4	45	94	47	8	46	20	15	52	53	54	55	26	57	28	29	09
Row 32	0000	0000	0000	0000	000		000	000	000	000	000	000	000	000	000	000	000	000	000	
Row 33	0000	• • • •	• • • •	0000	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	· · ·	· · ·	· · ·	• • •	· · ·	· · ·	• • •
Row 34	0000		• • • •	0000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	· · ·	· · ·	
Row 35	• • • •	• • • •		• • • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •		· · ·	• • •			
Row 36	0000				• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •			• • •			• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

17 18 19	6 37 38 39 40	57 58 59			model layer 4	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0			
14 15	3 34 35 36	54 55			assigned to	assigned to	assigned to 0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.	assigned to 0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.0.	assigned to the period of the
1 12	31 32 33	52			Santa Fe Group	Santa Fe Group 0. 0. 0. 0. 0. 0.	Santa Fe Group 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	Santa Fe Group 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	Santa Fe Group 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000
Mode 6	28 29 30	49			aphic unit of the	0°.00°.00°.00°.00°.00°.00°.00°.00°.00°.	to + 1 co 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	+ · · · · · · · · · · · · · · · · · · ·	to + co
6 7	26 27	46 47			le hydrostratigraphic				
٦ 4	23 24 25	44		65 64	ness	1	T	1	1
1	21 22 2	42	Ç		79	0000			
		Model	3	E > -	E)	Row -	Row 1		_

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	Ę									
Model	21 41 61	22 42 62	23 43 63	24 44 64	5 25 4 5	6 26 46	7 27 47	8 28 48	9 29 4 9	10 30 50	11 31 51	12 32 52	13 33 53	14 34 54	15 35 55	16 36 56	17 57 57	18 38 58	19 39 59	20 40 60
Row 6	0000	0000	0000	0000	000	• • •	000	• • •	• • •	000	• • •	• • •	000	• • •	• • •	000	000	000	000	• • •
Row 7			0000	• • • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	· · ·	• • •	• • •	• • •	• • •
% ₩			0000		• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	· · ·	000	• • •	• • •	• • •	• • •
80 ▼	0000	• • • •	0000	0000	• • •	• • •	• • •	000	000	000	• • •	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 10			0000		000	• • •	• • •	• • •	000	000	000	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •
Row 11		0. 220. 0.	0. 230. 0.	0000	• • •	• • •	000	000	· · ·	000	• • •	000	• • •	000	• • •	000	· • •	000	• • •	150. 0. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	n loo le	 									
	-	2	~	4	5	9	7	•	6	2	=		l		ļ	92		28	6	20
	21	22	23	24	25	56	27	28	53	30	31	32	33	75	35	36	37	38	39	40
Model	41	42	43	44	45	46	47		49	20	51					26		58	23	09
NO.	19	62	63	64																
Row 12	°	°	°	•	•	•	•	•	0	°	•	•	•	•		•		•	•	200.
	210.	220.	230.	240.	250.	270.	•	•	•	•	•	•	•	•	•	•	•	•	•	•
	•	•	•	•	°	•	•	•	•	•	•	•	•	•		•		•	•	•
	°	°	•	°																
Row 13	0	o	0	°	°	0	o	•	0	•	0	0	•	•					200	210.
	210.	220.	230.	250.	260	270.	280.	300	•	•	•	•	•	•	•	•	•		•	•
	° 0	.	· •	.	•	.0	•09	•	•	•	°	•	•	•				•	°	•
	o O	• •	.	• •																
Row 14	o	•	•	°	•	•	°	°	°	°	•	•	•	•				220.		210.
	210.	220.	230.	250.	270.	290•	320.	320.	340.	°	•	•	•	•0	•	•	•			•
	•	°	•	0	10.	50.	50.	50.	50.	°	•0	•	•0	•					•	•
	ċ	°	•	.																
Row 15	o	0	0	°	•	0	•	0	•	°	•0	•	•						210.	200
	220.	240.	260.	280	300	320.	340.	350.	360.	370.	°	•	•	•	•	•	•		•	•
	°	•	•	•	10.	40•	50.	50.	50.	40*	•	•	•					•	•	•
	ċ	°	°	°																
Row 16	°	0	0	•	•	•0	•	0	0	•	0	•							210.	200
	220.	240.	260.	280.	300	320.	340.	360.	380.	390	400	•	20.	30.	•	•	•		•	•0
	.	ં (.	.	10.	40•	50.	50.	40.	40.	30.	•						°	•	•
	• •	.	.	• •																
Row 17	0	•	0	•	°	•	•	0	•	0	•	•							210.	200.
	220.	240	260.	280.	300	320.	340.	360.	380.	400	420.	430.	10.	20.	30.	•	•	•	•	•
	.	o o	• o	° 0	•01	30.	40•	40•	40•	30.	20•	20.							•	•
	•	•	•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

					- 1		ı		Mode	8	ı	- 1	- 1	- 1	- 1	ı	- 1	- 1	- [
	- ;	2 5	س ز	₹ ;				ω (ο (2 9										50
MO40	5 7	55 7 2	23	7 7	52 ¥	56 26	27	78	50	0 6		32	33	¥ 2	35 55	ኢ አ	75	æ 2	39	40
MOL.	F 59	62	63	4				2	}	₹										3
9							1		1		1		1	ı	1			1	1	
¥0¥	2000	• 6	0 0	• 0	0 0															•002
	0,77	• 0	007	• 0 0	• • •	30.	40,	6	30.	30.	20.	, 20, 1	• • •	2 0	• 0	, 0				
	•	•	•	•	•															•
Row 19	•	•	0											80.		230.	220.	210.	200	210.
	230.	250.	270.		310.															•
	•	•	· •	· 0		20.	30.	30.	30.	20.	10.	.0	10.		•					•0
	°	.	•	°																
Row 20	°	•	°	•														210.	200	210.
	230.	250.	270.	290.	310.															•
	•	•	•	•		20.	30.	30.	20.	20.	10.	10.	•	•0	•	•	•			•0
	°	°	•	•																
Row 21	•	0	•	•																210.
	230.	250.	270.	290.	310.				380.					420.	410.	410.				•
	° c	° 0	ં લ	° c		20•	20°	20.		•	0.	°	•				°	•	•	•
	•	•	,	•																
Row 22	°	•	•	•	°				•	•	•		•	•	•	•				220.
	240.	260.	280.	300	320.															30.
	.	.	ં ં	ં ં	0	0	•	•				•					• •	•	°	•
	•	•	•	•																
Row 23	•	0	•	•																190.
	210.	230.	250.	270.	290.	310.	330.	350.	370.	390.	420.	420.	420.	420.	410.	410.	410.	410.	0	30.
	ž <	. c	. c	. c																•
	•	•	•	•																

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									No.											
	- ;	2 5	2 2	4 5	2 2	9 %	1- 5	ω ε	0 6	2 5	= =	12	5 :	14	55	16	17	<u>8</u> 8	62 8	20
Model	61	42 42 62	43 63	64 4 4 4 4	45	46	47	48	49	20 20	51.	52 52	53	7 7	55	2 %	57	28 28	20 70	60
Row 24	0. 150. 10.	0. 180. 20. 0.	210. 30.	0. 240. 0.	0. 270. 10.	0. 300. 10.	0. 330. 10.	0. 360. 10.	0. 390. 10.	410. 10.	0. 420. 0.	420. 0.	0. 420. 0.	0. 420. 0.	0. 4 10. 0.	0.4	0. 410. 0.	0. 0.	410. 0.	120. 4 10. 0.
Row 25	50.00	000	0. 130. 0.	.0 .0 .0	0. 200. 0.	0. 240. 0.	0. 270. 10.	0. 300. 10.	0. 340. 10.	0. 380. 10.	0. 420. 0.	0. 420. 0.	0. 420. 0.	0. 420. 0.	4 10. 0.	4 10. 0.	• • •	• • •	· · ·	• • •
Row 26	• • • •	50.	100.00.00	.0 .0 .0	.0 .0 .0	0. 220. 0.	0. 260. 0.	0. 300. 0.	0. 340. 0.	0. 380. 0.	0. 420. 0.	0. 420. 0.	0. 420. 0.	0. 420. 0.	000	000	000	· · ·	000	• • •
Row 27	0000	0000	50°.	001	0. 0. 0.	0. 180. 0.	0. 220. 0.	0. 260. 0.	300.	360.	0. 420. 0.	0. 420. 0.	0. 420. 0.	0. 420. 0.	• • •	• • •	• • •	• • •	• • •	• • •
Row 28	0000	0000	0 0 0	70.	200. 0.	0. 260. 0.	0. 310. 0.	0. 360. 0.	390 .	0. 400. 0.	0. 420. 0.	0. 420. 0.	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 29	0000	0000	0000	0 000	.0 .0 .0	0. 250. 0.	300.	360.	0. 370. 0.	400°.	• • •		· · ·	000	000	000	• • •	• • •	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	l column	٤									
	1 21	2 22	3 23	4 24	5 25	6 26	7 27	8 28	9 29	50 30	31	12 32	13 33	4 2	15 35	5 35	17 37	18 38	19 39	20 4 0
Model	19	42	63	44 64	45	94	41	84	49	20	51	52	53	54	55	56	57	28	59	09
Row 30	0000	0000	0000	0000	000	• • •	000	000	000	000	000	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	
Row 31	0000	••••	0000		• • •	• • •	• • •				• • •	• • •	• • •		• • •		• • •	• • •	• • •	· · ·
Row 32	0000	0000	0000		• • •	• • •	• • •	000	• • •	000	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •
Row 33	0000	0000	• • • •	0000	• • •	• • •	• • •	• • •	• • •	000	• • •	• • •	• • •	· · ·	• • •		• • •	• • •	• • •	• • •
Row 34	0000	• • • •	0000	0000	• • •	· · ·	000	•••		• • •	• • •		· · ·		000		· · ·	• • •	• • •	000
Row 35	0000	0000	• • • •		• • •	• • •	• • •	· · ·	• • •	• • •	• • •	• • •	• • •	· · ·	· · ·	· · ·	· · ·	• • •	• • •	• • •

Table 3.—"Inicknesses of hydrostratigraphic units assigned to each model layer, in feet—Continued

	1 21	22	23	4 4 5		55	96		6 7 8 26 27 28	6 7 8 26 27 28	Model col. 6 7 8 9 10 26 27 28 29 30	6 7 26 27	Model col. 6 7 8 9 10 26 27 28 29 30	6 7 8 9 10 11 26 27 28 29 30 31	6 7 8 9 10 11 12 26 27 28 29 30 31 32	6 7 8 9 10 11 12 13 14 15 26 27 28 29 30 31 32 33 34 35	6 7 8 9 10 11 12 13 14 26 27 28 29 30 31 32 33 34	6 7 8 9 10 11 12 13 14 15 26 27 28 29 30 31 32 33 34 35	6 7 8 9 10 11 12 13 14 15 16 26 27 28 29 30 31 32 33 34 35 36	6 7 8 9 10 11 12 13 14 15 16 17 26 27 28 29 30 31 32 33 34 35 36 37
Model row	41 61	45	63	3.2	45		4	_	84	84	48 49 50	48 49 50	48 49 50 51	48 49 50 51 52	48 49 50 51 52 53	48 49 50 51 52 53 54 55	48 49 50 51 52 53 54 55	48 49 50 51 52 53 54 55 56	48 49 50 51 52 53 54 55 56 57	48 49 50 51 52 53 54 55 56 57 58
Row 36	0000	0000	0000	0000	000	000	000		000		000		0.0000	0.0000000000000000000000000000000000000	0.0000000000000000000000000000000000000	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
			Thickm	Thickness of	lower	hydros	stratigraphi		raphic	raphic unit	c unit	c unit of the	c unit of the Santa	c unit of the Santa	c unit of the	c unit of the Santa Fe Group assigned	c unit of the Santa	c unit of the Santa Fe Group assigned	c unit of the Santa Fe Group assigned	c unit of the Santa Fe Group assigned
Row 1	0000	0000	0000	0000	000	000	000		000	000		000		.0000		0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
Row 2	0000	0000	0000	0000	000	000	000	_	000	0. 390. 0. 0.		390. 0.	390. 390. 0. 0. 0. 0.	390, 390, 390, 390, 390, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	390. 390. 390. 390. 390. 390. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	390, 390, 390, 390, 390, 340, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	390. 390. 390. 390. 390. 390. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	390, 390, 390, 390, 390, 340, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	390, 390, 390, 390, 390, 340, 290, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0	390, 390, 390, 390, 390, 340, 290, 250, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0, 0
Row 3	120. 50. 0.	120. 0. 0.	120. 0. 0.	0.000	0.00	0.00	0.00	520 . 0.	000	0. 520. 0. 0. 0. 0.		520. 0.	520. 520. 0. 0. 0. 0.	520. 520. 520. 520. 520. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	520. 520. 520. 520. 520. 460. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	520. 520. 520. 520. 520. 460. 420. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	520. 520. 520. 520. 520. 460. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	520. 520. 520. 520. 520. 460. 420. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	520 . 520 . 520 . 520 . 520 . 460 . 420 . 380 . 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	520. 520. 520. 520. 460. 420. 380. 340. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
Row 4	120. 70. 0.	120. 0.	120. 0.	120. 0.	120. 0.	0.00	0. 0.	(7)	0.00	.0. 10. 0. 0.		0.00	10. 0. 0.	10. 0. 510. 510. 460. 0. 0. 0. 0. 0. 0. 0. 0. 0.	10. 0. 510. 510. 460. 420. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	10. 0. 510. 510. 460. 420. 390. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	10. 0. 510. 510. 460. 420. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	10. 0. 510. 510. 460. 420. 390. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0	10. 0. 510. 510. 460. 420. 390. 350. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.	10. 0. 510. 510. 460. 420. 390. 350. 310. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	CIMIC									
	-	2	3	4	5	9	7	8	6	2	=	12	13	4	5	91	17	18	6	20
	21	22	23	54	25	56	27	28	53	30	31	32	33	¥	35	36	37	38	39	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	23	28	29	09
MO.	61	62	63	40																
		5	5	9	5	5	ć	ć	9	c	c	c						010	5	0,
8		071	•071	071	971	• 07 1	• 0,7	•07	<u>.</u>	.	.	.						• •	• 0	.
	• o	.	0 0
		• • • •								• •	.	•
Row 6	120	120.	120.	120.	120.	120.	120.	20.	20.	20.	20.	30.					240.	150.	80.	•
	0	•	°	•	•	•°	•	•	120.	200	300	350.	370.	380.	370.		•	°	•	•
	00	• •	• •	. .	°	•	°	.	°	°	•0	•0				.	•	•	.	•0
Row 7	120.	120.	120.	120.	120.	120.	130.	130.	140.	140.	140.	50.	50.						•06	•
	°	o	°	•	•	•	130.	250.	360.	380.	400	410.	400*	390.	370.	380	390.	400	400	•
	0	°	•	•	•	•	•	•	•	•	°	•	°						•	•
	o°	o°	°	.																
Row 8		120.	120.	120.	120.	120.	130.	130.	140.	150.	170.							250.	260.	270.
	290	°	°	•	•	30.	80.	190•	290.	400•	420.	420.	400•	390.	370.	390°	400.	410.	410.	410.
	ં લ	ં લ	.	.	•	°	•	•	°	°	°							°	•	•
	•	•	.	•																
Row 9		120	120.	120.	120.	120.		130.	130.	150.		170.						250.	280.	270.
	300	460.	460.	°	•	40.	140.	210.	290.	330.	440.	430.	410.	400	390	410.	420.	420.	420.	410.
	400.	400 • •		. .	°	o		.	.	o	
Row 10		120	120	120.	120.	120.	•	0	°	0	230.	220.	250.					280.	300.	330.
	360.	460.	460.	460.	460.	70.	160.	220.	300	350.	360.	480.	470	400	420.	430.	430.	430.	420.	410.
	•00 •00	400 0	410 .	. .	°	°	.	.	° 0					•	.	.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	Ē									
	-	2	2	4	5	9	7	8	6	2	=	12	13	4	2	16	12	82	62	20
	21	22	23	24	25	56	27	28	53	30	31	32	33	%	35	36	37	38	33	40
Model	4	42	43	44	45	46	47	48	49	20	51	25	53	72	55	26	57	28	29	09
MOL	19	62	63	64																
Row 11	120.	120.	120.	120.	•	•	•	•09	260.	300	340.	220.	250.				510.	590	.089	.061
	800	800	800	190	460.	470.	470.	480.	300	340.	350.	350.	460.	450.	450.	440.	440.	430.	420.	410.
	400	400•	420.	430.	°	•	.	°	.	•	°	°	°				°	°	•	•
	°	°	•	•																
Row 12	0	•	•	•	•		•	80	270.	320.	360	390	450.	500	550.	440.	510.		680	780.
	790		.064	780.	770.	750.	480.	480.	490	500.	350.	350.	350.	340.	440.	440.	430.	420.	420.	410.
	400.	400	450.	490.	°		•	°	•	•	•	•	•	•	•	•	•	•	•	•0
	• o		• •	.																
Row 13	•	•	•	°	•	•	•	•06	280.	330.	370.	400	460.	510.	560	620.	700	590	680	780.
	790•	. 06 <i>t</i>	.077	770	760•	740.	730.	710.	200	500	510.	350	350.	340.	340.	360.	370.	420.	420.	410.
	400	400•	440.	490	540.	570.	290	•	•	•	•	•	•	•	•	•	•	•	•	•
	°	.	.	°																
Row 14	0	•	•	•	•	•	°	120.	310.	360.	400	430.	480.	520.	580.	640.	720.	750.	.077	780.
	780.	780.	760.	160	750.	730.	710.	•069	6 70 •	540.	530.	2 80°	340.	340.	350.	350.	360.	380.	390	410.
	400	400•	440.	490.	550.	550.	2 80°	290	290	°	•	•	.	.	°	.	•	•	•	.
	•	•	.	.																
Row 15	•	•	•	°	•	•	°	260.	340.	390.	430.	460.	510.	550.	610.	670.	750.	760.	770.	780.
	780		760•	750.	730.	710.	•069	. 029	6 50 °	630.	550.	290	•009	350.	500	350.	360	370.	380	400•
	4 10.	400	440.	490.	520.	540.	2 80°	591.	2 80°	580.	•	°	•	•	•	•	•	•	•	•
	•		• •	•																
Row 16	•	•	•	•	•	•	210.	240.	330.	390	440.	470.	520.	570.	630.	.089	740.	750.	770.	780.
	780	760	750.	740.	720.	700	. 089	•099	640.	620	610.	•009	•009	•009	•009	290	2 80°	590	590	390.
	• 00 • 0	400 . 0	490 . 0	550 . 0.	570.	590.	290	290	•260°	280 °	570.	•	•	•	•	•	•	.	°	•

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	Model column	Ē									
	-	2	3	4	5	9	7	8	6	2	=	12	13	14	15	16	11	18	19	20
	21	22	23	24	25	56	27	28	53	ଛ	31	32	33	¥	35	፠	37	ጽ	39	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	22	28	59	09
₩ OJ	19	62	63	64																
Row 17	.	•	•	•	•	•	•	220.	330.	400	470.	500.					730.	750.	770.	780.
	780.	-	750.	730.	710.	•069	6 70 °	6 50 .	630.	620.	610.	•009	•009	•009	•009	•009	2 80°	590.	2 80°	2 80°
	290°	580.	550.	570.	2 80°	2 80°	2 80°	2 80°	2 90°	580.	570.	570.					°	•	°	°
	.		•	•																
Row 18	•	°	•	•	•	•	•	•	°	410.	500						740.	760.	.077	770.
	770.	.09/		720.	700.	. 089	•099	640.	620.	610.	. 009	•009	•009	•009	•009	•009	•009	2 90°	2 80°	590.
	2 80°	2 80°		590.	590	2 80°	2 80°	2 80°	590	580	570.						•	•	•	•
	°	°		.																
Row 19	°	•	°	•	•	0	0	°	°	•	500	550.					750.	760.	770.	770.
	760	750.	730.	710.	•069	670.	6 50 .	630.	620.	610.	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009
	590	590	590	2 80°	590	2 90°	2 80°	590.	590	580.	570.	570.					•	•	•	•
	o'	o	•	°																
Row 20	•	0	°	•	•	0	•	•	•	•	•						750.	760.	.077	770.
	760.	740.	720.	•069	•089	•099	640.	620.	610.	•009	•009	•009	•009		•009	•009	•009	•009	•009	•009
	•009	•009	590	590.	2 80°	590	580	580	580	580.	570.						•	•	°	•
	.	•	.	•																
Row 21	°	•	•	•	•	•	•	°	•	•	•							.097	.077	.077
	760.		720.	680.	670.	650.	630.	610 . 580	600° 580	600°.	600.	600°.	600 . 560	600 . 550	540 540	• 009	•009 •	• 009	• 009	•009
	• •	• •	0.0	• 0	•	•	•	•	•	•	•							•	•	:
Row 22	•	°	•	•	•	•	0	•	•	•	•						550.	760.	760.	760.
	750.	740.	720.	. 029	•099	640.	620.	610.	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009	•009
	•009	• 009	.065 0.	.065 0.	580 °	580 .	580.	580.	280°	280°	570.						o°	• •	• •	• •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

									Mode	el column	UMI									
	-	2	~	4	5	9	7	8	6	2	=	12	13	14	15	16	17	18	19	20
	21	22	23	24	25	56	27	28	59	8	31	32	33	¥	35	Ж	37	82	39	40
Model	14 4	42	43	44	45	46	47	48	49	20	51	52	53	54	55	26	22	58	29	09
ě O	ō	70		40																
Row 23	0	o	°	°	0	°	•	°	° 0	°	•	•	°	•	30.	30.	200.	560.	760.	760.
	750 . 590.	740.	720 . 590.	660 . 580 .	650 . 580 .	630 . 580 .	610 . 580 .	600 . 580 .	600 . 580 .	600 . 580 .	600 . 570 .	600 . 570 .	600 . 560 .	600 . 550.	600 . 540 .	600 . 530 .	.009	.009	.009	.009 0.
	•	•	•	.																
Row 24	°	•	•	•	•	0	•	•	•	•	0	•	•	•	•	•	38	330.	120.	750.
	740.	730.	720.	650 .	640 .	620 .	610 . 580	600°.	600 .	580	570	600.	500.	550	610 .	610.	610.	. 009	•009	•009
	• •	0	0	· °	•	•	•	•	•	•	•	•	•	•	•	2	•	•	•	•
Row 25	0	0	0	•	°	0	•	°	•	0	°	0		•		0	20.	190.	40.	20.
	730	7.30.	720.	650.	630.	610.	•009	•009	•009	.009	•009	•009	•009	610.	610.	610.	440.	440	450.	450.
	000	000	000	000	000	•000	•000	•000	580.	580.	570.	570.		550.		•	•	°	•	•
	o°	o ʻ	o [*]	o ʻ																
Row 26	°	°	°	0	°	•	°	•	•	•	•			•				•	•	•
	720.	720.	720.	6 50 °	620.	•009	•009	•009	•009	•009	•009	•009	610.	610.	610.	•000	•000	•000	•000	•000
	• 000	• • •	• • •	• • •	000	•000	• 000	• 000	•000	•000	•000			•000				•	•	•
Row 27	°	°	.	.	°	•	•	0	°	·°	.	.	.	•	• •	·°		•	•	°
	°	710.	720.	650.	610.	•009	590	590.	590.	•009	•009	•009	•009	610.	340.	•000	•000	•000	•000	•000
	• • • •	•°°	• • •	000	• 000	• 000	• 000	• 000	•000	•000	•000	•000	•000	• 000	°	•		•	°	°
Row 28	0	0	ó	ó	0	ó	0	0	o	0	ó	0	ó	o	ó	0	0	o	0	ó
	· • •	710.	720.	650.	•009	590.	580.	580.	580.	590.	•009 •	•009 •	220.	280.	000	•000	· • •	· • •	· • •	· • •
	. ·			· •	·	o ʻ	.	• •	.	.	•	.	.	• •	·	•	•	• •	;	.

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Continued

2		3	4	5	9	7	8	Mode I	10 1	IMD 1	12	13	4-	15	91	17	8	19	50
22 42 62		23	24 44 54	25 45	26 46	27	28	29 49	8 8	31	32 52	33 53	X X	35 55	26 28	37 57	8 88	39	40
000	1	0.710.	0.	000	0. 580.	0. 580.	0. 580.	0.	0.000	0.000.00	0.	• • • •	•••	000	•••	• • • •		•••	000
0 000		0 000	0 000	000	• • •	· · ·	• • •	· · ·	.0 .0 .0	• • •	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	000	000
		0 000	0 000	• • •	· · ·	· · ·	· · ·	• • •	• • •	· · ·	• • •	000	• • •	• • •		000	• • •	000	• • •
				000	• • •	• • •	• • •	• • •	• • •	· · ·	• • •	• • •	000		• • •	· · ·		• • •	· · ·
0000			0000		000	• • •	• • •	• • •	• • •	· · ·	• • •	000	• • •	• • •	• • •	000	• • •	• • •	000
0000	0000	0000	0000	000	000	· · ·	· · ·	· · ·	· · ·	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •	• • •

Table 3.--Thicknesses of hydrostratigraphic units assigned to each model layer, in feet--Concluded

									Mode	8	uwn									
	-	2	~	4	5	9	7	-	6	2	=	12	13	4	15	16	12	18	19	20
	21	22	23	24	25	56	27	28	59	જ્	31	32	33	¥	35	8	37	82	39	40
Model	41	42	43	44	45	46	47	48	49	20	51	52	53	54	55	99	57	28	59	09
MOL	19	62	63	99																
Row 35	°	0	•	•	•	•	•	•	•	•	•	•0	•	•	•0	•	•	0	0	•
	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•
	°	.	.	.	•	•	.	.	°	.	.	.	°	°	.	.
	•	•	•	•																
Row 36	°	.	.	•	•	•	•	•	•	•	•	•	•	•	•	•	.	•	•	.
	o	.	ċ	.	.	•	°	•	.	.	.	•	•	•	.
	•	.	°	.	.	.	•	•	•	•	°	•	.	•	•	•	•	•	•	•
	•	°	•	•																

Table 4.--Specified inflow to river reaches by stress period

	Specified inf	low, in cubic f	eet per second
Stress period	Reach 1, Rio Grande at Leasburg diversion	Reach 5, Rio Grande at Las Cruces wastewater outfall	Reach 45, drain at Anthony wastewater outfall
1	1,082.000	0.151	0.000
2	1,003.000	0.170	0.000
3	515.000	0.260	0.000
4	655.000	0.490	0.000
5	375.000	1.330	0.000
6	303.000	2.430	0.000
7	200.000	3.710	0.000
8	452.000	4.850	0.000
9	274.000	4.790	0.000
10	287.000	5.370	0.000
11	111.000	6.160	0.000
12	320.000	5.970	0.000
13	317.000	5.370	0.000
14	331.000	6.070	0.000
15	189.000	6.420	0.000
16	373.000	6.760	0.000
17	346.370	7.067	1.286
18	252.200	7.072	1.685
19	416.660	7.343	1.267
20	406.200	7.851	1.018
21	365.280	9.145	0.894
22	456.310	9.181	0.968

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model

[Units are cubic feet per second]

C L o a 1 y R u e o m	1915 —	1920	1927	1941	1948 	1951	1954	1958		1962	
r w n	1919	1926	1940	1947	1950	1953	1957	1 96 0	1961	1963	1964
1 13 8 1 14 7 1 15 7 1 16 4 1 16 8	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0
1 18 4 1 19 3 1 19 4 1 20 2 1 22 11	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 -0.007	-0.021	0 -2. 102			0 -1.071 -0.021	0 -0.981 -0.021
1 22 12 1 22 13 1 23 9 1 23 11 1 23 12	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0		0 -0.791 0 -0.605 -1.782	0 -0•974	0 -2.085 0 -0.958 -2.944	0 -0.945
1 23 13 1 23 42 1 23 51 1 23 53 1 24 11	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	-0.633 0 0 0 0	-1.613 0 0 0 0		-0.202 0.006 0 0	-1.058 0.014 0 0 0	-1.062 0.019 0.009 0
1 24 13 1 24 14 1 24 16 1 24 47 1 25 7	0 0 0 0	0 0 0 -0.046	0 0 0 -0•043	0 0 0 0	0 0 0 0	-0.633 0 0 0 0	-0.860 0 0 0	-0.638 0 0 0 0	-0.312 0 0 0 0	-0.441 0 0 0 0	-0.271 0 0 0 0
1 25 39 1 25 42 1 26 13 1 26 31 1 26 42	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0
1 26 52 1 26 54 1 26 55 1 27 39 1 28 11	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0	0 0 0 0

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

	C											
L	o 1											
•	u	1915	1 92 0	1927	1941	1948	1951	1954	1958		1962	
e o r w	m n	1919	1926	1940	<u> </u>	— 1950	1953	1957	1960	1961	1963	1964
1 28		0	0	0	0	0	-0.372					-0.820
1 28 1 29		-0.302 0	-0.3 02	-0.479 0	-1.285 0	-2.968 0	-0.372	-2.578 -1.137		-2.175 -0.935	-1.953 -0.824	-1.944 -0.820
2 11	-	Ö	ő	ő	Ö	ő	0	0	0	0	0	0
2 12		0	0	0	0	0	0	0	0	0	0	0
2 12	9	0	0	0	0	0	0	0	0	0	0	0
2 12		0	0	0	0	0	0	0	0	0	0	0
2 13		0	0	0	0	0	0	0	0	0	0	0
2 13 2 13		0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0
2 13	47	U	U	U	U	U	U	U	U	U	U	U
2 14		0	0	0	0	0	0	0	0	0	0	0
2 14		0	0	0	0	0	0	0	0	0	0	0
2 15		0	0	0	0	0	0	0	0	0	0	0
2 15 2 16		0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0
2 10	4	U	U	U	U	U	U	U	U	U	U	O
2 16		0	0	0	0	0	0	0	0	0	0	0
2 17		0	0	0	0	0	0	0	0	0	0	0
2 19 2 19		0 -0.047	0	0 -0.081	0	0	0	0 -0.052	0 -0.052	0 -0.052	0 -0.052	0 -0.053
2 19		0.047	0.000		-0.002						-0.029	-0.052 -0.029
2 21	12	0	0	0	0	0	0	0	0	0	0	-0.554
2 21		0	-	-0.053								
2 22		0	0	0	0	0		0	0	0	0	0
2 22		0	0	-0.003	-0.052	-0.058	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059
2 22	43	-0.072	-0.094	-0.129	-0.107	-0.119	-0.122	-0.122	-0.122	-0.122	-0.122	-0.122
2 23	9	0	0	0	0	0	0	0	0	0	0	0
2 23		0	0	0	0	0	0	0	0	0	0	0
2 23		0	0	0	0	0	0	0	0	0	0	-2.025
2 23		-0.039	-0.054	-0.061								
2 23	42	0	0	0	0	0	0	0	0	-0.013	-0.030	-0.039
2 23		0	0	0	0	0	0	0	0	0	0	0
2 23		0	0	0	0	0	0	0	0	0	0	0
2 23 2 23		0	0	0	0 -0.029	0 -0.029	0 -0.029	0 -0.029	0 -0.029	0 -0.029	0 -0.029	-0.018 -0.029
2 23		0 0	0 0	-0.023 0	0.029	0.029	0.029	0.029	0.029	0.029	0	0.029
2 23	53	U	U	U	U	U	U	U	U	U	U	U

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

C											
L o											
a 1											
y R u		1920	1927	1941	1948	1951	1954	1958		1962	
eom				_	-						
rwn		1926	1940	1947	1950	1953	1957	1960	1961	1963	1964
								1700			
2 22 50		0	0.000	-0.022	0.000	0.000	0.000	-0.022	0.000	0.000	-0.022
2 23 58 2 24 14		0 0	-0.002 0	0.022	-0.022 0	-0.022 0		0.022		-0.022 0	0.022
2 24 15		0	0	0	0	0	0			-4 .190	
2 24 16		Ö	0	0	0	0	0	0	0	0	0
2 24 17		0	0	0	0	0	0	0	0	0	0
2 24 17	O	Ü	Ū	Ū	J	Ū	Ü	Ü	U	O	O
2 24 41		-0.050	-0.080		-0.168			-0.178			
2 25 39		0	0	0	0	0	0	0	0	0	-0.001
2 25 41		0	0	0	0	0	0	0	0	-	0
2 25 42		0	0	0	0	0				-0.826	
2 25 45	0	0	0	0	0	0 ,	0	0	0	0	0
2 25 50	0	0	0	0	0	0	0	0	0	0	0
2 26 10		0	0	0	0	0	0	0	0	0	0
2 26 16		0	0	0	0	0	0	0	0	0	0
2 26 17		0	0	0	0	0	0	-0.200			-0.116
2 26 31		0	-0.002	-0.022	-0.022	-0.022	-0.022				
2 26 41	0	0	-0 016	-0.174	-0.239	-0.299	-0.421	-0.575	-0.723	-0.806	-0.949
2 26 41		0	0	0.174		0.299	0.421	0.5/5	0.723		
2 26 47		0	0	0		-	-1.138				•
2 26 47		0	0	0		0					
2 26 52		0	0	0	0	0	0	0	0.930	0.023	0.621
2 20 32	U	U	U	U	U	U	U	U	U	U	O
2 26 54		0	0	0	0	0	0	0	0	0	0
2 26 55		-0.037		-0.033							
2 27 14		0	0	0	0	0	0		0		-
2 27 15		0	0	0	0	0				-0.403	
2 27 16	0	0	0	0	0	0	0	0	0	0	0
2 27 17	0	0	0	0	0	-0.135	-0.919	-0.514	-1.283	-1.732	-2.470
2 27 18	0	0	0	0	0	0	-0.169				
2 27 19	-0.038	-0.042	-0.053	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059
2 27 36	0	0		-0.023		-0.025	-0.025	-0.025	-0.025	-0.025	-0.025
2 27 39	0	0	0	0	0	0	0	0	0	0	-0.001
2 27 50	0	0	0	0	0	0	0	-1.054	-0.936	-0.825	-0.821
2 28 14		0	0	0	0	0	0	0	0.930	-0.049	
2 28 15		0	0	0	0	0	0	0	0	0	0.070
2 28 17		0	-0.034	-0.283	-0.290	-0.496	-0.530				-0. 481
2 28 44		0	0	0.203	0.230	0.430	0	0.550	0.330	0.551	0
	~	~	~	-	-	-	-	•	~	~	~

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

C L o a 1 y R u e o m	1915 —	1920	1927	1941	1948 	1951	1954 —	1958		1962	
r w n	1919	1926	1940	1947	1950	1953	1957	1960	1961	1963	1964
2 28 45	^						0	^			
2 28 46	0	0	0	0	0	0	0	0	0 -0.936	0 -0.825	0 -0.821
2 28 48	0	0	0	0	0	0	0	-0.300			
2 28 49	0	0	0	0	0	0	0	0	0	0	0
	0	0	0	0	0	0	0	0	0	0	0
2 28 50	0	0	0	0	0	0	0	0	0	0	-0.821
2 29 23	0	-0.033	-0.042	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021
2 29 43	0	0	0	0	0	0	-0.581	-1.054	-0.936	-0.825	-0.821
2 29 44	0	0	0	0	0	0	0	-0.300	-0.936	-0.825	-0.821
2 29 45	0	0	0	0	0	0	0	0	0	-0.826	-0.822
2 29 46	0	0	0	0	0	0	0	0	0	-0.826	-1.643
2 30 44	0	0	0	0	0	0	0	0	0	0	0
3 24 15	0	0	0	0	0	0	-0.94 0	-6. 794	-6. 693	-9.118	-7.8 05
3 24 16	0	0	0	0	0	0	-0.147	-3.149	-2.184	-2.321	-0.597
3 24 17	0	0	0	0	0	0	0	0	-0.334	-2.040	-3.746
3 25 16	0	0	0	0	0	0	0	-1.455	-3. 175	-2.249	-2.396
3 26 14	0	0	0	0	0	0	0	0	0	0	0
3 26 16	0	Ö	Ö	Ö	Ö	Ö	Ö	0	0	ő	Ö
3 26 17	0	ő	ő	0	ő	ő	Ö	0	ő	Ö	Ö
3 27 17	0	Ő	0	0	0	Ő	ő	0	0	0	Ŏ
3 27 17	0	ő	0	Ő	ő	Ö	Ö	Ő	ő	o O	Ö
J 27 10											
Totals	-0.54	-0.76	-1.14	-2.40	-4.23	-12.76	-25.36	-35.63	-39.86	-47.80	-52.21

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

C											
L o											
a 1											
y R u	1965	1967	1969		1973		1977	1979	1981	1983	
eom					-						
rwn	1966	1968	1971	1972	1975	1976	1978	1 98 0	1982	1984	1985
1 13 8	0	0	0	0	0	-0.387	-0 664	-0.226	0	0	0
1 14 7	0	0	0	0	0	-0.387	-0.420			-0.462	
1 15 7	Ö	Ö	0	0	0	0.765		1.285	1.279	1.676	1.715
1 16 4	Ö	Ö	Ö	Ö	0	-2.817	-2.305			-1.405	-1.708
1 16 8	Ö	Ö	Ö	0	Ö	0		-0.417	0	0	0
1 18 4	-2.236	-2.307		-3.238	-3.336	-0.377				-0.073	
1 19 3	0	0	0	0	0	-0.9 70	-0.794			-0.255	
1 19 4	-0.9 11		-0.961	-1.321		0	0	0	0	0	0
1 20 2		-0.021				-0.021	-0.021		-0.021		
1 22 11	-0.422	-0.694	-0.650	-0.662	-0.493	-0.136	-0.432	-0.038	-0.001	0	0
1 22 12	0	0	0	0	0	-0.056	-0.056	-0.056	-0.056	-0.056	-0.056
1 22 13	-2.204	-2.881				-0.340	-1.696		0	-0.015	0
1 23 9	0	0	0	0	0	0	-0.014		-0.028		-0.028
1 23 11		-1.485				-0.256	-1.099				-1.366
1 23 12		-2.149				-0.101	-0.461		0	-0.005	0
1 23 13	-1.040	-1.646	-1.085		-0.844	-0.386	-1.924		-0.021	-0.020	0
1 23 42	0.033	0.039	0.037		0.060	0.029	0.076	0.092	0.101	0.124	0.116
1 23 51	0.015	0.018	0.017	_	0.032	0.053	0.067	0.073	0.076	0.079	0.078
1 23 53	0	0	0	0	0	0	0	0.042	0.176	0.135	0.109
1 24 11	0	0	-0.001	-0.063	-0.023	-0.030	-0.030	-0.034	-0.029	-0.023	-0.058
1 24 13	-0.379	-0.959	-0.643	-0.608	-0.520	-0.094	-0.662	-0.095	0	-0.008	0
1 24 14	-0.341	-0.964	-1.234	-1.684	-1.085	-0.315	-1.078	-0.160	-0.165	-0.344	-0.510
1 24 16	0	0	0	0	0	-0.011	-0.028	-0.029	-0.033	-0.031	-0.025
1 24 47	0	0	0	0	0	0	0	0	0	0	0
1 25 7	0	0	0	0	0	0	0	0	-0.008	-0.036	-0.023
1 25 39	0.004	0.007	0.007	0,008	0-012	0.022	0-027	0-028	0-022	0-027	0.026
1 25 42	0	0	0	0	0	0.022	0	0.020	-0.064		-0.111
1 26 13	0	0	0	0	0	-0.015		-0.022		-0.022	-0.020
1 26 31	0	0	0	0	0	0.019	0.106	0.165		0.286	0.332
1 26 42	0	0	0	0	0	-0.657		-0.325		0.200	0.552
· · -	-	-	-	=	-		5.505		55225	-	ū
1 26 52	0	0	0	0	0	0	0.094	0 . 069	0.067	0.073	0.087
1 26 54	0	0	0	0	0	0	0	0	0	0.057	0.172
1 26 55	0	0	0	0	0	0	0.062	0.179	0.084	0.116	0.106
1 27 39	0.009	0.018	0.023	0.031		0.065		0.082	0.092	0.092	0.114
1 28 11	0	0	0	0	0	-0.046	-0.046	-0.046	-0.046	-0.046	-0.046

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

C											
L o a 1											
y Ru	1 9 65	1967	1969		1973		1977	1979	1981	1983	
eom	1905	1907	1909		1973		17//	17/7	1901	1903	
rwn	1966	1968	1971	1972	1975	1976	1978	1980	1982	1984	1985
	1900	1700	17/1	1972	1973	1770	1970	1700	1702	1704	
1 28 43	-0.802	- ∩-714	-0.216	0-002	0.002	0.002	0.002	0.002	0.002	0.002	0.002
1 28 46	0	0	0	0	0	0	0	0	0	0	0
1 29 43	-0.802		-0.701		-0.710	-0.742		-0.401	-	-0.375	-
2 11 7	0	0	0	0	0	0	0	-0.345		-0.873	
2 12 8	0	0	0	0	0	0	0	0	0	-0.747	
	-	, and the second					-				
2 12 9	0	0	0	0	0	0	0	0	-0.008	-0.062	-0.023
2 12 10	0	0	0	0	0	-0.039	-0.039	-0.039	-0.039	-0.039	-0.039
2 13 8	0	0	0	0	-0.787	-0.387	-0.420	-0.640	-1.080	-0.973	-1.036
2 13 9	0	0	0	0	-0.380	-0.387	-0.420	-0.226	-0.329	-0.945	-1.019
2 13 47	0	0	0	0	0	0	0	0	0	-0.136	-0.156
2 14 7	0	0	0	0	- 0.787	-0.441	-0.479	-0.558	- 0.561	-1.374	-1.505
2 14 8	0	0	0	0	0	-0.053	-0.166	-0.27 0	-0.293	-0.294	
2 15 6	0	0	0	0	0	-0.595	-0.627	-0.586		-0.628	
2 15 7	0	0	0	0	- 0.787	-0.387	-0.438	-0.417	-0.263	-0.193	-0.254
2 16 4	0	0	0	0	0	0	0	-0.313	-0.254	0	0
2 16 7	0	0	0	0	0	0	0.490	0 /17	0.410	0 112	0.007
2 16 7	0	0	0	0	0	0 0	-0.489 -0.489	-0.417 -0.417	-0.410		
2 17 7	0	0 0	0	0	0				0 -0.002	0 -0.004	0 -0.006
2 19 11	0	•	0 -0.052	0 053	0 -0.052	0 - 0 . 052	0 - 0•052	0 -0.052		-0.052	
2 19 16	-0. 052	-0. 052			-0.032	-0.032 -0.029	-0.032		-0.029	-0.029	
2 19 49	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029
2 21 13	-1.847	-0.559	-0.126	-0.133	-0.066	0	-0.482	-1.047	-0.386	-0.912	-1.657
2 21 23	-0.030	-0.030	-0.030		-0.030	-0.030	-0.030	-0.030	-0.030	-0.030	-0.030
2 22 13	0	0	-0.079	-0.811	-0.210	-0.164	-2.286	-1.462	-1.789	-1.828	-0.093
2 22 31	-0.059	-0.059			-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059
2 22 43		-0.122			-0.122	-0.122	-0.122	-0.122	-0.122	-0.122	-0.122
2 23 9	0	0	0	0	0	0	0	0	0	-0.017	-0.017
2 23 10	0	0	0	0	0	-0.028	-0.038	-0.079	-0.078	-0.109	-0.117
2 23 12	-0.865	-0.790	-1.738	-1.968	- 0.768	-1.101	-1.727	-1.756	-2.463	-2.401	-1.629
2 23 28	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059	-0.059
2 23 42	-0.068	-0.081	-0.078	-0.087	- 0.126	-0.057	-0.151	-0. 183	-0.201	- 0.249	-0.231
2 23 47	0	0	0	0	0	-0.569	-1.650	-2.165	-2.373	-2.207	-1.095
2 23 50	0	0	0	0	0	0.302	-0.671	-0.307	-0.736	-0.625	-0.685
2 23 51	-0.029	-0.034	-	-0.039		-0. 107	-0.134	-0.147	-0. 151	-0.158	-0.155
2 23 52	-0.029	-0.029	-0.029		-0.029	-0.029	-0.029	-0.029	-0.029	-0.029	-0.029
2 23 53	0	0.023	0	0	0.025	0	0	-0.085			
~ ~ ~ ~	J	U	J	0	9	J	•	0.003	01000	0-200	0.0

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Continued

C						· · · · · · · · · · · · · · · · · · ·					
L o											
a 1											
y R u	1965	1967	1 9 69		1973		1977	1979	1981	1983	
e o m		_								-	
r w n	1966	1968	1971	1972	1975	1976	1978	1980	1982	1984	1985
2 23 58	-0.022	-0.022	-0.022	-0.022	-0-022	-0.022	-0.022	-0.022	-0.022	-0.022	-0.022
2 24 14		-0.711	-0.745			-0.113	-2.632				-1.718
2 24 15	-3.164	-3.023				-2.884	-5.021			-3.242	-2.654
2 24 16	0	0	-0.014	-0.096	-0.376	-0.578	-1.955	-1.114	-0.660	-0.488	-0.897
2 24 17	-0.386	-0.799	-0.384	-2.156	0	-0.421	-2.986	-2.397	-0.895	-0.965	-0.805
2 24 41		-0.178						-0.116			
2 25 39	-0.008		-0.013			-0.044	-0.053			-0.054	
2 25 41	0	0	0		0	0			-1.047		
2 25 42	-0.804		-0.703					0	0	0	0
2 25 45	0	0	0	U	0	U	0	-0.640	-1.314	-0.902	-0.643
2 25 50	0	0	0	0	0	0	0	-0.238	-0.403	-0.584	-0.571
2 26 10	0		-0.009			-0.028	-0.028		-0.028	-0.028	-0.028
2 26 16	0		0			-0.048	-0.021		0	0	0
2 26 17		- 0.296					-0.988				
2 26 31	-0.022	-0.022	-0.015	-0.001	-0.001	-0.119	-0.212	-0.330	-0.453	- 0.572	-0.665
2 26 41	-1.162	-1.256	-1.639	-1.969	-2.020	-2.188	-2.253	-2.313	-2.394	-2.584	-2.727
2 26 43	0	0	0	0	0	0	0	0	- 0 .9 41	-1.159	-0.830
2 26 47		-0.715				-0.246		-0.522		-0.246	-0.175
2 26 49	-0.803	-0.715				-0.036		-0.162		0	0
2 26 52	0	0	0	0	0	0	-0.188	-0.139	-0.135	-0.146	-0.174
2 26 54	0	0	0	0	0		0	0	0		-0.344
2 26 55		-0.033						-0.358			
2 27 14	0	0		0			0			0	0
2 27 15						-0.634		-0.373			
2 27 16	U	0	U	Ü	Ü	0	0	0	U	-0.002	-0.010
2 27 17		-3. 101						-1.465			0
2 27 18		-0.290				0	0	0 (01	0	0	0
2 27 19				-0.139		-0.267		-0.491			
2 27 36					-0.025	-0.025		-0.025 -0.165			
2 27 39	-0•018	-0.035	-0.046	-0.001	~ ∪•∪ 9 ∠	-0.130	70•144	~0•100	~0•104	~0•100	~∪•220
2 27 50		-0.715				0	0	0	0	0	0
2 28 14		-0.087		-0.102		-0.135		-0.165			
2 28 15	0	0	0	0	0	0	0	0		-0.017	
2 28 17		-0.457				0	0	0	0	0	0
2 28 44	-0.395	-0.716	-0.703	-0.660	-0.712	-0.476	-0.551	-0.693	-1.146	-1.222	-1.438

Table 5.—Schedule of nonagricultural ground-water withdrawal specified in the model—Concluded

L o o a 1 1		C											
a 1 y R u 1965 1967 1969 1973 1977 1979 1981 1983 e o m </td <td>L</td> <td></td>	L												
y R u 1965 1967 1969 1973 1977 1979 1981 1983 e o m													
r w n 1966 1968 1971 1972 1975 1976 1978 1980 1982 1984 1985 2 28 45 0 0 -0.703 -0.660 -0.712 -0.710 -0.948 -0.610 -0.485 -0.749 -0.398 2 28 46 -0.803 -0.715 -0.932 -1.318 -1.421 -1.673 -1.064 -1.060 -1.029 -1.576 -1.482 2 28 48 0 0 -0.230 -0.659 -0.711 -0.618 -0.574 -1.022 -0.705 -0.871 -0.615 2 28 49 0 0 -0.703 -0.660 -0.712 -1.420 -0.867 -0.588 -1.116 -1.459 -2.165 2 28 50 -1.198 -1.431 -1.405 -1.318 -1.421 -1.715 -1.468 -1.800 -1.838 -2.051 -2.203 2 29 <td>y R</td> <td></td> <td>1965</td> <td>1967</td> <td>1969</td> <td></td> <td>1973</td> <td></td> <td>1977</td> <td>1979</td> <td>1981</td> <td>1983</td> <td></td>	y R		1965	1967	1969		1973		1977	1979	1 9 81	1983	
2 28 45 0 0 -0.703 -0.660 -0.712 -0.710 -0.948 -0.610 -0.485 -0.749 -0.398 28 46 -0.803 -0.715 -0.932 -1.318 -1.421 -1.673 -1.064 -1.060 -1.029 -1.576 -1.482 28 48 0 0 -0.230 -0.659 -0.711 -0.618 -0.574 -1.022 -0.705 -0.871 -0.615 28 49 0 0 -0.703 -0.660 -0.712 -1.420 -0.867 -0.588 -1.116 -1.459 -2.165 28 50 -1.198 -1.431 -1.405 -1.318 -1.421 -1.715 -1.468 -1.800 -1.838 -2.051 -2.203 29 23 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.021 -0.887 -0.857 -0.857 -0.857 29 44 -0.803 -0.715 -0.702 -0.659 -0.711 -1.133 -1.171 -1.138 -0.603 -0.875 -0.857 29 44 -0.803 -0.715 -0.702 -0.659 -0.711 -1.299 -1.221 -1.121 -0.347 -0.347 -0.347 29 45 -0.804 -0.716 -0.703 -0.660 -0.712 -1.973 -1.501 -1.432 -0.943 -1.403 -1.419 2 29 46 -1.607 -1.432 -1.406 -1.319 -1.422 -3.094 -2.519 -2.142 -1.825 -2.104 -1.975 2 30 44 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	e o	m											
2 28 46	r w	n	1966	1968	1971	1972	1975	1976	1 9 78	1980	1 9 82	1984	1985
2 28 46													
2 28 46	2 28	45	0	0	-0.703	-0.660	-0.712	-0.710	-0.948	-0.610	-0.485	-0.749	-0.398
2 28 48 0 0 0 -0.230 -0.659 -0.711 -0.618 -0.574 -1.022 -0.705 -0.871 -0.615 2 28 49 0 0 0 -0.703 -0.660 -0.712 -1.420 -0.867 -0.588 -1.116 -1.459 -2.165 2 28 50 -1.198 -1.431 -1.405 -1.318 -1.421 -1.715 -1.468 -1.800 -1.838 -2.051 -2.203 2 29 23 -0.021													
2 28 49 0 0 0 -0.703 -0.660 -0.712 -1.420 -0.867 -0.588 -1.116 -1.459 -2.165 2 28 50 -1.198 -1.431 -1.405 -1.318 -1.421 -1.715 -1.468 -1.800 -1.838 -2.051 -2.203 2 29 23 -0.021													
2 28 50			_										
2 29 23			-	_									
2 29 43		-	14170	1. 131	1.103	14310	10121	10,15	14 100	1,000	1000	20031	20203
2 29 44	2 29	23	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021	-0.021
2 29 44	2 29	43	-0.803	-0.715	-0.702	-0.659	-0.711	-1.133	-1.171	-1.138	-0.603	-0.875	-0.857
2 29 46 -1.607 -1.432 -1.406 -1.319 -1.422 -3.094 -2.519 -2.142 -1.825 -2.104 -1.975 2 30 44 0 0 0 0 0 0 0 -0.333 -1.374 3 24 15 -5.566 -4.974 -6.494 -5.247 -6.429 -7.053 -5.076 -6.549 -8.116 -5.255 -7.551 3 24 16 -2.174 -1.343 -1.610 -1.888 -3.460 -4.195 -2.970 -2.652 -2.938 -2.535 -2.172 3 24 17 -4.480 -3.657 -1.626 -1.367 -2.494 -3.211 -1.506 -3.131 -2.432 -1.398 -0.866 3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 17 0 0 0 0 0 0 0 0 0 <td>2 29</td> <td>44</td> <td>-0.803</td> <td>-0.715</td> <td>-0.702</td> <td>-0.659</td> <td>-0.711</td> <td></td> <td>-1.221</td> <td>-1.121</td> <td>-0.347</td> <td>-0.347</td> <td>-0.347</td>	2 29	44	-0.803	-0.715	-0.702	-0.659	-0.711		-1.221	-1.121	-0.347	-0.347	-0.347
2 30 44 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	2 29	45	-0.804	-0.716	-0.703	-0.660	-0.712	-1.973	-1.501	-1.432	-0.943	-1.403	-1.419
3 24 15 -5.566 -4.974 -6.494 -5.247 -6.429 -7.053 -5.076 -6.549 -8.116 -5.255 -7.551 3 24 16 -2.174 -1.343 -1.610 -1.888 -3.460 -4.195 -2.970 -2.652 -2.938 -2.535 -2.172 3 24 17 -4.480 -3.657 -1.626 -1.367 -2.494 -3.211 -1.506 -3.131 -2.432 -1.398 -0.866 3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 0 0 0 0 -0.055 -0.070 -0.054 -0.060 -0.053 3 27 17 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 18 0 0 0 0 -0.502 -0.578 -0.631 <	2 29	46	-1.607	-1.432	-1.406	-1.319	-1.422				-1.825		-1.975
3 24 15 -5.566 -4.974 -6.494 -5.247 -6.429 -7.053 -5.076 -6.549 -8.116 -5.255 -7.551 3 24 16 -2.174 -1.343 -1.610 -1.888 -3.460 -4.195 -2.970 -2.652 -2.938 -2.535 -2.172 3 24 17 -4.480 -3.657 -1.626 -1.367 -2.494 -3.211 -1.506 -3.131 -2.432 -1.398 -0.866 3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 0 0 0 0 -0.055 -0.070 -0.054 -0.060 -0.053 3 27 17 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 18 0 0 0 0 -0.502 -0.578 -0.631 <													
3 24 16 -2.174 -1.343 -1.610 -1.888 -3.460 -4.195 -2.970 -2.652 -2.938 -2.535 -2.172 3 24 17 -4.480 -3.657 -1.626 -1.367 -2.494 -3.211 -1.506 -3.131 -2.432 -1.398 -0.866 3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 0 0 0 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 17 0 0 0 0 0 0 0 -0.578 -0.631 -0.610 -0.692	2 30	44	0		0	0	0	0	0	0	0	-0.333	-1.374
3 24 17 -4.480 -3.657 -1.626 -1.367 -2.494 -3.211 -1.506 -3.131 -2.432 -1.398 -0.866 3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 -0.048 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 17 0 0 0 0 0 0 0 -0.631 -0.610 -0.692	3 24	15	-5.566	- 4 . 974	-6.494	-5.247	-6. 429	-7. 053	-5.076	-6. 549	- 8.116	-5. 255	-7.5 51
3 25 16 -2.588 -1.858 -2.595 -1.654 -3.153 -3.520 -2.178 -2.033 -2.690 -2.484 -1.117 3 26 14 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 -0.258 3 27 17 0 0 0 0 0 0 -0.602 -0.435 -0.408 3 27 18 0 0 0 -0.252 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692	3 24	16	- 2.174	-1.343	-1.610	-1.888	-3.460	-4.195	-2.97 0	- 2.652	-2.938	- 2•535	-2.172
3 26 14 0 0 0 0 0 0 0 -0.022 -0.022 -0.015 3 26 16 0 0 0 0 -0.015 -0.048 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 -0.258 3 27 17 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 18 0 0 0 -0.252 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692	3 24	17	-4.48 0	-3.657	-1.626	-1.367	-2.494	-3.211	-1.506	-3.131	-2.432	-1.398	-0.866
3 26 16 0 0 0 -0.015 -0.048 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 0 0 -0.258 3 27 17 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 18 0 0 0 -0.252 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692	3 25	16	-2.588	-1.858	- 2 . 595	-1.654	-3. 153	-3.520	-2.178	-2.033	-2.69 0	-2.484	-1.117
3 26 16 0 0 0 -0.015 -0.048 -0.055 -0.070 -0.054 -0.060 -0.053 3 26 17 0 0 0 0 0 0 0 0 0 0 -0.258 3 27 17 0 0 0 0 0 0 0 -0.023 -0.435 -0.408 3 27 18 0 0 0 -0.252 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692				_	_	_			_				
3 26 17 0 </td <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>•</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td>							•						
3 27 17 0 0 0 0 0 0 0 -0.435 -0.408 3 27 18 0 0 0 0 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692													
3 27 18 0 0 0 0 -0.252 -0.402 -0.502 -0.578 -0.631 -0.610 -0.692											-	-	
							-	-	_				
Totals -47.63 -48.99 -49.74 -55.99 -53.48 -54.58 -68.52 -62.56 -59.58 -60.69 -59.86	3 27	18	0	0	0	0	-0.252	-0.402	-0.502	- 0.578	-0.631	-0.610	-0.692
Totals -47.63 -48.99 -49.74 -55.99 -53.48 -54.58 -68.52 -62.56 -59.58 -60.69 -59.86													
100°6C_ 60°00, 00°4C_ 00°77 -00°77 -00°77 - 22°00, 14 -24- 24°00, 00°74 -00°77 -00°77	Tata	10	_/17 62	_/,8_00	_/,0 7/.	_55 00	_53 //9	_5/, 50	_68 52	-6 2 56	_50 50	60 60	_50 86
	Iota	ц	-4/.03	~ 40∙33	-47. /4		-JJ•40	- 24•20	-00 •J2	-02.30	صر∙در−	-00.09	-25.00

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>

[* indicates sites used for sensitivity tests]

Latitude	Longitude	Well number	Moo Row	del Col- umn	Measured hydraulic head (feet)	Model- derived hydraulic head (feet)	Dif- fer- ence (feet)
		Comparison i	for 1	947, la	yer 1		
* 31°49'14"	106°53'13"	29S.1E.6.110	4	12	3,865.00	3,798.70	-66.30
* 31°55'21"	106°54'27"	27S.1W.26.430	5	21	3,803.60	3,808.20	4.60
* 32°02'19"	106°57'04"	26S.1W.16.330	5	29	3,804.00	3,824.10	20.10
* 32°08'53"	106°59'18"	25S.2W.12.240	6	36	3,842.00	3,833.20	-8.80
* 31°49'15"	106°46'34"	28S.2E.31.340	7	9	3,811.50	3,787.60	-23.90
* 31°57'45"	106°51'27"	27S.1E.17.210	7	22	3,780.00	3,807.20	27.20
* 32°00'47"	106°53'26"	26S.1W.25.410	7	26	3,760.00	3,816.20	56.20
* 32°02'58"	106°52'12"	26S.1E.18.220	8	28	3,820.00	3,820.30	0.30
* 32°07'09"	106°52'16"	25S.1E.19.240	9	32	3,785.00	3,830.20	45.20
* 32°09'23"	106°53'03"	25S.1E.6.330	10	35	3,910.00	3,839.20	-70.80
* 31°59′57"	106°46′18"	26S.2E.31.410	11	22	3,808.00	3,803.00	-5.00
* 31°51′38"	106°41′43"	28S.2E.24.110	12	10	3,782.00	3,774.80	-7.20
* 32°05′25"	106°46′52"	25S.2E.31.130	13	28	3,820.00	3,819.40	-0.60
* 32°02′47"	106°45′02"	26S.2E.17.240	14	25	3,804.00	3,809.70	5.70
* 32°17′02"	106°52′36"	23S.1E.30.210	14	48	3,882.30	3,875.00	-7.30
31°54'10"	106°39'05"	28S.3E.5.140	16	11	3,767.60	3,766.10	-1.50
31°47'47"	106°35'40"	29S.3E.12.300	17	3	3,747.20	3,743.60	-3.60
32°17'25"	106°51'01"	23S.1E.21.314	17	48	3,887.10	3,882.50	-4.60
31°47'10"	106°34'22"	29S.4E.18.233	18	2	3,750.00	3,740.10	-9.90
31°47'39"	106°34'02"	29S.4E.7.440	19	2	3,745.70	3,735.70	-10.00
31°59'35"	106°40'26"	27S.3E.6.213	19	20	3,768.00	3,790.10	22.10
32°00'28"	106°40'40"	26S.3E.31.123	19	21	3,790.00	3,792.00	2.00
32°06'30"	106°44'03"	25S.2E.28.220	19	28	3,815.30	3,819.00	3.70
32°01'20"	106°40'48"	26S.3E.30.114	20	22	3,791.70	3,794.70	3.00
32°15'25"	106°48'07"	23S.1E.36.333	20	41	3,863.70	3,870.80	7.10
32°15′59"	106°48'35"	23S.1E.35.231	20	43	3,866.50	3,874.40	7.90
32°16′39"	106°48'45"	23S.1E.26.311	20	45	3,870.00	3,877.90	7.90
31°48′54"	106°34'01"	JL-49-12-501	21	4	3,729.50	3,734.00	4.50
32°01′40"	106°40'19"	26S.3E.19.432	21	22	3,789.80	3,793.50	3.70
32°15′03"	106°47'37"	24S.1E.1.144	21	40	3,862.50	3,868.60	6.10

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>--Continued

						Model-	
					Measured	derived	Dif-
			Ma	1.1			
			MO	del	hydraulic	hydraulic	fer-
T 1	· · · ·	77 11 1		Col-	head	head	ence
Latitude	Longitude	Well number	Row	umn	(feet)	(feet)	(feet)
22272422	1050/0/10	222 - 12 //2			0.005.10	0.000.60	0.50
32°19′02"	106°49′13"	23S.1E.10.442	21	50	3,885.10	3,888.60	3.50
32°22′04"	106°50′45"	22S.1E.28.142	21	54	3,907.60	3,908.10	0.50
31°57′37"	106°37′36"	27S.3E.15.143	22	16	3,771.60	3,775.60	4.00
32°09′44"	106°44′03"	25S.2E.4.422	22	32	3,824.40	3,836.30	11.90
32°24′29"	106°51′44"	22S.1E.8.421	22	57	3,916.50	3,923.20	6.70
32°26′18"	106°52′27"	21S.1E.31.412	22	60	3,935.40	3,937.60	2.20
32°28′42"	106°53′54"	21S.1W.13.323	22	62	3,947.60	3,945.30	-2.30
32°29′34"	106°54'40"	21S.1W.11.431	22	63	3,956.00	3,950.80	-5.20
32°16′06"	106°46′35"	23S.2E.31.213	23	42	3,865.90	3,872.60	6.70
32°23′10"	106°50′37"	22S.1E.21.211	23	5 6	3,909.60	3,918.00	8.40
32°06′56"	106°41′28"	25S.2E.24.413	24	28	3,813.80	3,819.20	5.40
32°16′45"	106°46′20"	23S.2E.30.412	24	43	3,867.40	3,873.30	5.90
32°06′52"	106°40′58"	25S.3E.19.331	25	27	3,814.10	3,814.90	0.80
32°14′43"	106°45′06"	24S.2E.5.422	25	38	3,858.40	3,858.60	0.20
32°19′48"	106°47'19"	23S.1E.1.443	25	50	3,887.50	3,894.70	7.20
32 19 48	100 47 19	233.16.1.443	23	50	5,887.50	3,874.70	7.20
32°23′23"	106°49′37"	22S.1E.15.431	25	55	3,908.40	3,917.20	8.80
32°10′11"	106°42′14"	25S.2E.2.221	26	31	3,829.00	3,827.70	-1.30
32°12′00"	106°43′05"	24S.2E.22.444	26	34	3,839.00	3,842.90	3.90
32°16′46"	106°45′18"	23S.2E.29.243	26	42	3,867.60	3,869.90	2.30
32°18′46"	106°45′19"	23S.2E.17.210	28	46	3,863.10	3,870.60	7.50
32°18′59"	106°45′19"	23S.2E.8.434	28	47	3,870.30	3,882.50	12.20
32°20′45"	106°46′37"	22S.2E.31.340	28	51	3,896.00	3,898.00	2.00
32°10′10"	106°40′16"	24S.3E.31.430	29	31	3,817.00	3,832.20	15.20
* 32°07′19"	106°37′25"	25S.3E.22.120	32	26	3,825.00	3,815.70	-9.30
52 07 15	100 37 23	238.36.22.120	32	20	3,023.00	3,013.70	7.50
		Comparison	of 19	75, lay	ver 1		
* 31°53′55"	106°59′06"	28S.1W.6.323	3	21	3,813.37	3,810.60	-2.77
* 31°59′08"	107°00′50"	27S.2W.2.411	3	27	3,823.60	3,825.50	1.90
* 32°03′06"	107°03′36"	26S.2W.17.214	3	32	3,828.80	3,844.30	15.50
* 32°02'30"	107°01′31"	26S.2W.15.443	4	31	3,833.70	3,829.90	-3.80
* 31°55′36"	106°54′46"	27S.1W.26.413	5	21	3,810.69	3,807.70	-2.99
* 32°00′54"	106°53′39"	26S.1W.25.414	7	26	3,822.00	3,815.70	-6.30
* 32°07'46"	106°57′20"	25S.1W.16.331	7	35	3,835.00	3,833.70	-1.30
* 32°03′09"	106°52′16"	26S.1E.18.222	8	28	3,814.47	3,819.80	5.33
* 31°59′55"	106°49′03"	26S.1E.35.333	9	23	3,795.50	3,806.10	10.60
* 32°12′39"	106°56′01"	24S.1W.22.123	9	41	3,872.53	3,847.70	-24.83
32 12 37	200 00	2.0.14.22.123	_	• •	0,012.55	2,2	

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>--Continued

Latitude	Longitude	Well number	Mod	del Col- umn	Measured hydraulic head (feet)	Model- derived hydraulic head (feet)	Dif- fer- ence (feet)
* 32°06′51"	106°51′11"	25 S .1E.21.331	10	31	3,836.17	3,827.40	-8.77
* 32°11'23"	106°53′23"	24S.1W.25.422	10	37	3,848.00	3,843.90	-4.10
* 31°49′52"	106°41′35"	28S.2E.36.142	11	8	3,776.35	3,772.90	-3.45
* 31°51′18"	106°42′26"	28S.2E.23.324	11	10	3,782.47	3,775.90	-6.57
* 32°08'26"	106°51'12"	25S.1E.16.114	11	33	3,836.69	3,833.70	-2.99
* 32°05′28"	106947400#	050 0E 21 122	13	28	2 01/ 20	3 010 70	4.10
	106°47′02"	25S.2E.31.133			3,814.30	3,818.40	
* 31°50′50"	106°38′49"	28S.3E.28.114	14	8	3,754.95	3,757.90	2.95
31°52′12"	106°38′19"	28S.3E.16.414	16	9	3,749.45	3,750.20	0.75
32°16′07"	106°50′26"	23S.1E.33.214	16	45 11	3,876.70	3,873.30	-3.40
31°54′27"	106°38′57"	28S.3E.4.111	17	11	3,759.36	3,760.00	0.64
31°54′57"	106°39′05"	27S.3E.32.244	17	12	3,762.65	3,763.60	0.95
32°17′36"	106°50′59"	23S.1E.20.134	17	48	3,877.15	3,880.30	3.15
31°51′52"	106°37′19"	JL-49-12-108	18	8	3,746.61	3,744.70	-1.91
31°58′37"	106°40′25"	27S.3E.7.231	18	18	3,779.71	3,778.80	-0.91
32°15′39"	106°49'22"	23S.1E.34.423	18	43	3,874.85	3,872.00	-2.85
31°48′25"	106°34′50"	29S.4E.7.111	19	4	3,731.53	3,733.60	2.07
31°55′30"	106°38′33"	27S.3E.28.341	19	13	3,765.59	3,764.70	-0.89
31°59′28"	106°40′23"	27S.3E.6.231	19	19	3,782.88	3,783.50	0.62
31°59′53"	106°40′39"	26S.3E.31.341	19	20	3,784.04	3,787.30	3.26
32°18′53"	106°50′40"	23S.1E.9.433	19	51	3,890.87	3,889.50	-1.37
010/0/00	106004400#	TT /0 10 F00	0.0	,	2 752 15	2 721 20	00.05
31°49′20"	106°34′38"	JL-49-12-502	20	4	3,752.15	3,731.30	-20.85
31°51′52"	106°35′39"	JL-49-12-103	20	7	3,747.55	3,743.40	-4.15
32°00′44"	106°40′41"	26S.3E.30.343	20	21	3,789.23	3,789.00	-0.23
32°15′44"	106°48′08"	23S.1E.35.424	20	42	3,861.49	3,870.60	9.11
32°19′21"	106°50′01"	23S.1E.10.134	20	51	3,893.93	3,889.50	-4.43
31°48′54"	106°34′01"	JL-49-12-501	21	4	3,729.80	3,732.30	2.50
31°56′39"	106°38′04"	27S.3E.21.421	21	15	3,765.74	3,769.90	4.16
31°59′10"	106°39'13"	27S.3E.5.414	21	19	3,776.12	3,783.90	7.78
32°00'41"	106°39'45"	26S.3E.29.334	21	20	3,786.85	3,786.50	-0.35
32°03′36"	106°41′11"	26S.2E.12.421	21	24	3,805.12	3,801.10	-4.02
32°26′11"	106°53′09"	21s.1W.36.221	21	60	3,932.88	3,937.20	4.32
32°27'17"	106°53′31"	21S.1W.25.232	21	61	3,942.00	3,940.40	-1.60
31°55′51"	106°37′22"	JL-49-04-407	22	13	3,763.27	3,760.30	-2.97
31°55′52"	106°37′22	JL-49-04-408	22	13	3,762.56	3,760.30	-2.26
31°57′48"	106 37 11" 106°37'58"	27S.3E.16.224	22	13 17	3,762.36	3,777.00	5.24
31 37 48"	TOD 21.28.	213.3E.10.224	22	Ι/	3,111.70	3,777.00	J. 24

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>--Continued

							· · · · · · · · · · · · · · · · · · ·
			Mod	del	Measured hydraulic head	Model- derived hydraulic	Dif- fer-
Latitude	Longitude	Well number	Row	Col- umn	nead (feet)	head (feet)	ence (feet)
	201161 2440	WOII HOURDOI			(1000)	(1000)	
01050401	10400040	077 07 0 0/0			0 775 (4	0 777 00	1 00
31°58′31"	106°38′10"	27S.3E.9.243	22	17	3,775.61	3,777.00	1.39
32°03′41"	106°40′34"	26S.3E.7.144	22	24	3,800.59	3,802.40	1.81
32°20′10"	106°49′14"	23S.1E.3.422	22	52	3,890.22	3,894.90	4.68
31°55′56"	106°36′31"	JL-49-04-417	23	12	3,758.51	3,762.00	3.49
31°56′19"	106°36′43"	JL-49-04-405	23	13	3,758.88	3,762.60	3.72
31°59′28"	106°38′10"	27S.3E.4.241	23	18	3,780.18	3,779.40	-0.78
31°57′33"	106°36′44"	JL-49-04-115	24	16	3,767.24	3,773.50	6.26
32°09′06"	106°42′36"	25S.2E.11.142	24	30	3,824.37	3,826.00	1.63
32°18′18"	106°47′04"	23S.2E.18.313	24	47	3,867.40	3,880.00	12.60
32°18′27"	106°47'35"	23S.1E.13.411	24	47	3,868.35	3,880.00	11.65
32 10 27	100 47 55	230.10.13.411	2-4	7,	3,000.33	3,000.00	11.03
31°56′37"	106°35′43"	JL-49-04-436	25	13	3,751.45	3,764.30	12.85
31°57′11"	106°35′42"	JL-49-04-439	25	14	3,747.97	3,767.40	19.43
31°59′06"	106°36′51"	JL-49-04-122	25	18	3,778.33	3,778.40	0.07
31°58′04"	106°35′43"	JL-49-04-138	26	16	3,748.42	3,771.40	22.98
32°16′53"	106°45′18"	23S.2E.29.234	26	42	3,858.00	3,863.90	5.90
32 10 33	100 10 10	200.22,27,29			•,	- ,	
32°08′46"	106°40′48"	25S.3E.7.312	27	29	3,818.09	3,822.20	4.11
32°12′05"	106°42′43"	24S.2E.23.341	27	33	3,836.87	3,838.10	1.23
31°59′53"	106°35′40"	JL-49-04-102	28	18	3,767.45	3,778.40	10.95
31°59′53"	106°35′31"	JL-49-04-101	28	18	3,790.50	3,778.40	-12.10
32°04′05"	106°37′31"	26S.3E.3.344	28	23	3,801.54	3,797.50	-4.04
200054/0#	1060204271	050 28 22 110	20	25	2 902 11	2 906 90	2 (0
32°05′48"	106°38′37"	25S.3E.33.112	28	25	3,803.11	3,806.80	3.69
32°22′10"	106°46′40"	22S.2E.30.123	29	53	3,890.98	3,902.40	11.42
* 32°08′24"	106°35′50"	25S.3E.13.111	34	27	3,812.00	3,821.40	9.40
		Comparison i	for 19	975, la	ayer 2		
* 31°59′18"	106°39′13"	27S.3E.5.412	21	19	3,776.76	3,781.70	4.94
* 31°55′54"	106°36′57"	JL-49-04-418	22	12	3,758.84	3,756.90	-1.94
* 31°56′19"	106°37′05"	JL-49-04-404	22	13	3,758.23	3,756.80	-1.43
* 32°14′10"	106°46′27"	24S.2E.7.231	22	37	3,854.20	3,853.80	-0.40
* 31°55′56"	106°36′45"	JL-49-04-410	23	12	3,760.62	3,753.90	-6.72
<u> </u>					• –	•	
* 31°56′52"	106°36′23"	JL-49-04-425	24	14	3,746.07	3,751.10	5.03
* 31°57′03"	106°36′43"	JL-49-04-427	24	15	3,752.63	3,747.10	-5.53
* 31°57′20"	106°36′22"	JL-49-04-422	24	15	3,754.35	3,747.10	-7.25
* 31°58′19"	106°37′07"	JL-49-04-110	24	17	3,762.88	3,763.40	0.52
* 32°04′14"	106°39′58"	26S.3E.6.442	24	24	3,799.09	3,803.00	3.91

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>--Continued

Model hydraulic hydraulic head head (feet) Latitude Longitude Well number Row umn (feet) (feet) * 31°58′03" 106°35′43" JL-49-04-165 26 16 3,737.46 3,753.80 3 32°16′23" 106°44′56" 23S.2E.28.333 26 41 3,838.00 3,839.20 32°19′14" 106°46′25" 23S.2E.7.411 27 48 3,856.72 3,853.70 32°18′32" 106°45′13" 23S.2E.17.243 28 46 3,834.68 3,826.30	Dif- fer- ence feet) 16.34 1.20 -3.02 -8.38 25.19
* 32°16'23" 106°44'56" 23S.2E.28.333 26 41 3,838.00 3,839.20 * 32°19'14" 106°46'25" 23S.2E.7.411 27 48 3,856.72 3,853.70 * 32°18'32" 106°45'13" 23S.2E.17.243 28 46 3,834.68 3,826.30	1.20 -3.02 -8.38 25.19
* 32°16'23" 106°44'56" 23S.2E.28.333 26 41 3,838.00 3,839.20 * 32°19'14" 106°46'25" 23S.2E.7.411 27 48 3,856.72 3,853.70 * 32°18'32" 106°45'13" 23S.2E.17.243 28 46 3,834.68 3,826.30	1.20 -3.02 -8.38 25.19
* 32°19′14" 106°46′25" 23S.2E.7.411 27 48 3,856.72 3,853.70 4 32°18′32" 106°45′13" 23S.2E.17.243 28 46 3,834.68 3,826.30	-3.02 -8.38 25.19
* 32°18'32" 106°45'13" 23S.2E.17.243 28 46 3,834.68 3,826.30	-8.38 25.19
	25.19
" 32 10 33 100 43 21 233.2E.0.443 20 47 3,013.31 3,041.10 A	
	0.50
* 32°22'20" 106°47'10" 22S.1E.25.222 28 53 3,897.70 3,897.20	
	-0.41
,	17.61
32 17 30 100 43 31 230,24,3,342 27 47 3,030,77 3,037,70	.,
Comparison for 1975, layer 3	
* 31°56′27" 106°36′37" JL-49-04-416 23 13 3,738.59 3,749.40 1	0 0 1
	10.81
· · · · · · · · · · · · · · · · · · ·	L5.80
	3.60
·	-0.14 L9.76
* 31°58'17" 106°37'06" JL-49-04-113 24 17 3,724.64 3,744.40 1	.9.70
* 32°16'40" 106°46'12" 23S.2E.30.243 25 42 3,849.00 3,859.80	10.80
	-0.20
	0.20
Comparison for 1984-85, layer 1	
* 32°01'12" 107°04'57" 26S.2W.30.233 1 31 3,844.00 3,862.70 1	18.70
* 31°53′29" 106°59′33" 28S.1W.7.113B 2 20 3,801.00 3,810.50	9.50
* 31°56′10" 107°00′20" 27S.2W.25.111 3 23 3,805.00 3,814.90	9.90
* 32°02'30" 107°01'35" 26S.2W.15.434 3 31 3,841.00 3,841.80	0.80
	9.70
* 31°48'12" 106°51'45" 29S.1E.8.124 5 10 3,797.00 3,793.30	-3.70
* 32°04′28" 106°56′39" 26S.1W.4.412 6 31 3,826.00 3,826.20	0.20
	14.10
	0.50
* 31°59′55" 106°49′00" 26S.1E.35.332 9 23 3,800.00 3,805.50	5.50
	17.60
	28.50 ·1.70
	3.40
, , , , , , , , , , , , , , , , , , ,	
* 31°57′03" 106°45′01" 27S.2E.21.111 11 18 3,788.00 3,788.90	0.90

Table 6.--<u>Differences between measured and model-derived</u>
<u>hydraulic heads</u>--Concluded

Latit	ude	Longitude	Well n	umber	<u>Mo</u> Row	del Col- umn	Measure hydraul head (feet)	ic	Mode derive hydrau head (feet	ed lic	Dif- fer- ence (feet)
* 31°5	1'01"	106°41'07"	28S.2E.	24 444	12	8	3,766.	00	3,76	4 30	-1.70
* 31°5		106°46'07"	26S.2E.		12	22	3,795.		3,79		4.40
* 31°4		106°39'41"	28S.3E.		13	7	3,766.		3,75		-7.70
* 32°1		106°53′15"	23S.1W.		13	47	3,870.		3,86		-2.70
* 31°5		106°38′32"	28S.3E.		15	8	3,748.		3,75		4.80
* 31°5	2′38"	106°39′23"	28S.3E.	17.214	15	9	3,758.	00	3,75	3.30	-4.70
	0'11"	106°37′18"	28S.3E.		16	6	3,750.		3,74		-1.20
	6′07"	106°44′05"	25S.2E.			28	3,818.		3,81		-0.30
32°0	4'05"	106°37′31"	26S.3E.	3.344D	29	23	3,801.	00	3,79	9.20	-1.80
			Compari	son for	198	4-85,	layer 2				
* 32°0	2155"	106°57′08"	26S.1W.	16.334	6	29	3,821.	00	3,82	2.60	1.60
* 32°0		106°51′12"	25S.1E.		11	33	3,837.		3,83		-3.80
* 32°0		106°46′57"	25S.2E.			28	3,821.		3,81		-3.00
* 32°1		106°47'21"	24S.1E.			37	3,853.		3,853		0.40
* 32°1	7′45"	106°49′21"	23S.1E.			47	3,872.		3,874		2.50
* 31°5	5′54"	106°36′57"	JL-49-0	4-418	22	12	3,732.	00	3,75	L.50	19.50
* 32°1		106°46′27"	24S.2E.		22	37	3,854.	00	3,85		-1.60
* 32°1		106°48′10"	23S.1E.	23.244A	22	46	3,867.	00	3,869	9.40	2.40
* 32°1 * 31°5		106°44'45" 106°36'12"	24S.2E. JL-49-0		24 25	35 14	3,842. 3,730.		3,844 3,745		2.20 15.60
				Summ			·				
		Number	Ext	reme		Ave	rage diff	erei	nces (f	eet)	
		of		rences					Root		
		differ-	(fe	et)		A	pproximat	e	mean	Мe	an
Layer	Perio	od ences	Small	Large	M	ean	median	5	square	abso.	lute
1	104	, E./.	70.00	F.C. 20	1	0/	2 60		18.61	10	7/.
1 1	1947 1975		-70.80 -24.83			.94 .43	3.60 1.09		7.72		.74 .46
2	1975		-8.38			.18	-0.40		9.11		.13
3	1975		-15.80			.09	-0.40		11.25		.73
1	1984-8		-28.50			.86	-1.45	•	9.47		.42
2	1984-8		-3.00				1.90		8.18		.26
			Al	l diffe	renc	e s					
							0.90		12.43	7	. 34
		Differences					:11 . **-11		/ .l. \		
		Differences	at site	s outsi	de ti	ne Mesi	illa vall	ey ((*)		

Table 7.--Average differences, in feet, between model-derived and measured hydraulic heads for the standard and each sensitivity test

[Averages included 90 differences for 1947, 1975, and 1984-85. Approximate median is the average of the two values nearest the median]

Abbreviated test name	Arith- metic mean	Mean absolute	Approx- imate median	Root mean square
S 0.3 1.5E-6	0.23	9.44	-0.25	16.09
STANDARD	-0.03	9.45	-0.46	16.06
S 0.40 2.E-6	0.41	9.45	-0.05	16.13
KMSFD7 3	-0.24	9.49	-0.45	16.03
KMSFD15 8	0.19	9.50	-0.57	16.14
S 0.13 .67E-6	-0.32	9.52	-0.60	16.06
KLSF20	0.57	9.66	-0.47	16.29
ANISOTROPY 150,750,750,150	-0.86	9.67	-1.70	15.98
ANISOTROPY 67,333,333, 67	0.61	9.78	0.00	16.39
KUSF45 18	1.00	9.79	0.51	16.37
KUSF20 8	-1.27	9.86	-1.47	16.26
KLSF9	-0.67	9.98	-0.31	16.53
KMSFS30 15	-0.96	10.00	-0.98	16.61
KMSFS13 7	1.17	10.25	0.35	16.74

Explanation of abbreviated test names:

S 0.3 1.5E-6

Specific yield (layers 1 and 2) set to 0.3 and specific storage (layers 2-4) set to 1.5 x 10^{-6} per foot.

STANDARD

The standard is the model described in this report with no changes. Hydraulic-conductivity values of the standard, using the abbreviations explained below, were as follows:

KUSF=30 and 12 feet per day, KMSFS=20 and 10 feet per day, KMSFD=10 and 5 feet per day, and KLSF=13 feet per day.

Horizontal to vertical hydraulic-conductivity anisotropy ratios were:

100:1 for the river alluvium, 500:1 for the upper and middle

hydrostratigraphic units of the Santa Fe Group, and

100:1 for the lower hydrostratigraphic units of the Santa Fe Group.

Specific yield for layers 1 and 2 was 0.2, and specific storage for layers 2-4 was 1 x 10^{-6} foot of thickness.

Table 7.--Average differences, in feet, between model-derived and measured hydraulic heads for the standard and each sensitivity test--Concluded

Explanation of abbreviated test names:

S 0.40 2.E-6	Specific yield (layers 1 and 2) set to 0.4 and specific storage (layers 2-4) set to 2 x 10^{-6} per foot.
KMSFD7 3	Hydraulic conductivity of the deep part of the middle hydrostratigraphic unit of the Santa Fe Group (KMSFD) set to 3 feet per day in the Las Cruces area (rows 25-36, columns 35-64) and 7 feet per day elsewhere.
KMSFD15 8	KMSFD set to 8 feet per day in the Las Cruces area and 15 feet per day elsewhere.
S 0.13 .67E-6	Specific yield (layers 1 and 2) set to 0.13 and specific storage (layers 2-4) set to $6.7~{\rm E}^{-7}$.
KLSF20	Hydraulic conductivity of the lower hydrostratigraphic unit of the Santa Fe Group set to 20 feet per day throughout the model.
ANISOTROPY 150,750,750,150	Horizontal to vertical anisotropy ratios set to 150 for the river alluvium, 750 for the upper and middle hydrostratigraphic units of the Santa Fe Group, and 150 for the lower hydrostratigraphic unit of the Santa Fe Group.
ANISOTROPY 67,333,333, 67	Horizontal to vertical anisotropy ratios set to 67 for the river alluvium, 333 for the upper and middle hydrostratigraphic units of the Santa Fe Group, and 67 for the lower hydrostratigraphic unit of the Santa Fe Group.
KUSF45 18 (KUSF)	Hydraulic conductivity of the upper hydrostratigraphic unit of the Santa Fe Group set to 18 feet per day in the Las Cruces area and 45 feet per day elsewhere.
KUSF20 8	KUSF set to 8 feet per day in the Las Cruces area and 20 feet per day elsewhere.
KLSF9	KLSF set to 9 feet per day throughout the model.
KMSFS30 15	Hydraulic conductivity of the shallow part of the middle hydrostratigraphic unit of the Santa Fe Group set to 15 feet per day in the Las Cruces area and 30 feet per day elsewhere.
KMSFS13 7	KMSFS set to 7 feet per day in the Las Cruces area and 13 feet per day elsewhere.

